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5 **Evaluating Landsat Thematic Mapper spectral indices for estimating burn severity of**
6 **the 2007 Peloponnese wildfires in Greece**

7 Running head: Evaluating spectral indices for estimating burn severity

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15 **Brief summary.** This paper evaluates the performance of three different spectral indices for
16 estimating burn severity. The indices were pre/post-fire differenced and correlated with field
17 data of severity. In addition the burned pixels' bi-temporal shifts in the corresponding bi-
18 spectral feature spaces were studied. Results reveal the importance of the short-wave and mid
19 infrared spectral region in complement to the near infrared spectral region for assessing post-
20 fire effects. Further research directions for estimating burn severity with remote sensing data
21 are given.

22 **Abstract.** A vast area (more than 100 000 ha) of forest, shrubs and agricultural land burned
23 down at the Peloponnese peninsula in Greece during the 2007 summer. Three pre/post-fire
24 differenced Landsat Thematic Mapper (TM) derived spectral indices were correlated with
25 field data of burn severity for these devastating fires. These spectral indices were the

26 Normalized Difference Vegetation Index (NDVI), the Normalized Difference Moisture Index
27 (NDMI) and the Normalized Burn Ratio (NBR). The field data consists of 160 Geo
28 Composite Burn Index (GeoCBI) plots. In addition, indices were evaluated in terms of
29 optimality. The optimality statistic is a measure for the index's sensitivity to fire-induced
30 vegetation depletion. Results show that the GeoCBI-dNBR (differenced NBR) approach
31 yields a moderate-high $R^2 = 0.65$ whereas the correlation between field data and the
32 differenced NDMI (dNDMI) and the differenced NDVI (dNDVI) was clearly lower
33 (respectively $R^2 = 0.50$ and $R^2 = 0.46$). The dNBR also outperformed the dNDMI and
34 dNDVI in terms of optimality. The resulting median dNBR optimality equalled 0.51 while
35 the median dNDMI and dNDVI optimality values were respectively 0.50 and 0.40
36 (differences significant for $p < 0.001$). However, inaccuracies observed in the spectral indices
37 approach indicate that there is room for improvement. This could imply improved
38 preprocessing, revised index design or alternative methods.

39 **Additional Keywords:** fire severity; Normalized Burn Ratio, Normalized Difference
40 Vegetation Index, spectral index, wildfires, Geo Composite Burn Index, optimality.

41 **Introduction**

42 Wildfires play a major role in Mediterranean Type Ecosystems (MTEs) (Vazquez and
43 Moreno 2001; Diaz-Delgado et al. 2004; Pausas 2004; Pausas et al. 2008) as they partially or
44 completely remove the vegetation layer and affect post-fire vegetation composition, water and
45 sediment regimes, and nutrient cycling (Kutiel and Inbar 1993). As such they act as a natural
46 component in vegetation succession cycles (Trabaud 1981; Capitaino and Carcaillet 2008;
47 Roder et al. 2008) but also potentially increase degradation processes, such as soil erosion
48 (Thomas et al. 1999; Chafer 2008; Fox et al. 2008). Assessment of the fire impact is thus a
49 major challenge to understand the potential degradation after fire (Kutiel and Inbar 1993; Fox

50 et al. 2008) and to comprehend ecosystem's post-fire resilience (Epting and Verbyla 2005;
51 Lentile et al. 2007).

52 The terms fire severity and burn severity are often interchangeably used (Keeley 2009)
53 describing the amount of damage (Hammill and Bradstock 2006; Gonzalez-Alonso et al.
54 2007; Chafer 2008) the physical, chemical and biological changes (Landmann 2003; Chafer et
55 al. 2004; Cocke et al. 2005; Stow et al. 2007; Lee et al. 2008) or the degree of alteration
56 (Brewer et al. 2005; Eidenshink et al. 2007) that fire causes to an ecosystem. Some authors,
57 however, suggest a clear distinction between both terms by considering the fire disturbance
58 continuum (Jain et al. 2004), which addresses three different temporal fire effects phases:
59 before, during and after the fire. In this context, fire severity quantifies the short-term fire
60 effects in the immediate post-fire environment (Lentile et al. 2006) and is usually measured in
61 an initial assessment scheme (Key and Benson 2005). As such, it mainly quantifies vegetation
62 consumption and soil alteration. Burn severity, on the other hand, quantifies both the short-
63 and long-term impact as it includes response processes (e.g. resprouting, delayed mortality),
64 which is evaluated in an extended assessment (EA) that incorporates both first- and second-
65 order effects (Lentile et al. 2006; Key 2006). In this study burn severity, defined as the
66 absolute magnitude of environmental change caused by a fire (Key and Benson 2005), is
67 estimated one year post-fire.

68 Several remote sensing studies have discussed the potential of satellite imagery as an
69 alternative for extensive field sampling to quantify burn severity over large areas. These
70 studies evaluated the use of spectral unmixing, simulation techniques and spectral indices to
71 assess burn severity (for a comprehensive review of remote sensing techniques for burn
72 severity assessment, see Kasischke et al. 2007; French et al. 2008). Spectral mixture analysis
73 (Rogan and Yool 2001; Lewis et al. 2007; Robichaud et al. 2007) and simulation models
74 (Chuvieco et al. 2006; De Santis and Chuvieco 2007; De Santis et al. 2009) have proven to

75 provide valuable information with regards to burn severity. Spectral indices, however, are a
76 more popular approach, mainly because of their computational and conceptual simplicity.
77 These spectral indices are typically based on Normalized Difference Spectral Indices
78 (NDSIs), such as the Normalized Difference Vegetation Index (NDVI) (Isaev et al. 2002;
79 Chafer et al. 2004; Diaz-Delgado et al. 2004; Ruiz-Gallardo et al. 2004; Hammill and
80 Bradstock 2006; Hudak et al. 2007) or the widely used Normalized Burn Ratio (NBR) (e.g.
81 Lopez-Garcia and Caselles; Epting et al. 2005; Key and Benson 2005; Miller and Thode
82 2007). The NDVI combines the reflectance in the R (red) and NIR (near infrared) spectral
83 region and is a measure for the amount of green vegetation, whereas the NBR relates to
84 vegetation moisture by combining the NIR with MIR (mid infrared) reflectance. Since fire
85 effects on vegetation produce a reflectance increase in the R and MIR spectral regions and a
86 NIR reflectance drop (Pereira et al. 1999), bi-temporal image differencing is frequently
87 applied on pre- and post-fire NDVI or NBR images. This results respectively in the
88 differenced Normalized Difference Vegetation Index (dNDVI) (Chafer et al. 2004; Hammill
89 and Bradstock 2006) and the differenced Normalized Burn Ratio (dNBR) (Key and Benson
90 2005). The advantage of these pre/post-fire differenced indices is that they permit a clear
91 discrimination between unburned sparsely vegetated areas and burned areas, which is difficult
92 in mono-temporal imagery (Key and Benson 2005).

93 A wide range of field data has been considered to validate the remotely sensed indices for
94 estimating burn severity: % live trees (Lopez-Garcia and Caselles 1991; Alleaume et al. 2005;
95 Smith et al. 2007) or % tree mortality (Kushla and Ripple 1998; Isaev et al. 2002), basal area
96 mortality (Chappell and Agee 1996), combustion completeness (Alleaume et al. 2005),
97 changes in Leaf Area Index (LAI) (Boer et al. 2007) and fractional cover of several
98 components (Kokaly et al. 2007; Lewis et al. 2007; Robichaud et al. 2007). However, by far
99 the most widely used field measurement is the Composite Burn Index (CBI) (Key and Benson

100 2005). The CBI is a semi-quantitative field sampling approach based on an expert judgement
101 procedure, developed as an operational methodology for validating remotely sensed
102 assessments of burn severity on a national scale in the USA as part of the FIREMON (Fire
103 Effects Monitoring and Inventory Protocol) project. The CBI is fundamentally different to the
104 above-mentioned field approaches because in the CBI the sample plot is considered in a
105 holistic way. Several attributes (e.g. char height, % LAI change...) of the plot are visually
106 examined and numerically rated per ecosystem stratum (substrates, low shrubs, tall shrubs,
107 intermediate trees and high trees). The total plot score, which is an average of the average
108 stratum ratings, expresses the plot's burn severity. Recently, GeoCBI, a modified version of
109 the CBI, has been developed (De Santis and Chuvieco 2009). The main modification of the
110 GeoCBI consists of the consideration of the fraction of coverage (FCOV, the percentage of
111 cover with respect to the total extension of the plot) of the different vegetation strata, which
112 resulted in a more consistent relation between the GeoCBI and the remotely sensed burn
113 severity measure (De Santis and Chuvieco 2009). The GeoCBI-dNBR relationship recently
114 experienced a knowledge gain for the North American boreal region (Epting et al. 2005, Allen
115 and Sorbel 2008; Hall et al. 2008; Hoy et al. 2008; Murphy et al. 2008). However, studies that
116 assessed the empirical relationship between vegetation indices and field data in the fire-prone
117 Mediterranean biome (De Santis and Chuvieco 2007) are underrepresented in literature.

118 The dNBR approach has been questioned (Roy et al. 2006) as it was initially developed for
119 detecting burned areas (Lopez-Garcia and Caselles 1991) rather than evaluating within-burn
120 differences in combustion completeness. To evaluate dNBR index performance, a pixel-based
121 optimality measure originating from the spectral index theory (Verstraete and Pinty 1996),
122 which varies between zero (not at all optimal) and one (fully optimal), has been developed
123 (Roy et al. 2006). An optimal burn severity spectral index needs to be very sensitive to fire-
124 induced vegetation changes and insensitive to perturbing factors such as atmospheric and

125 illumination effects. Very low mean optimality values were reported using in situ reflectance,
126 Landsat Enhanced Thematic Mapper plus (ETM+) and Moderate Resolution Imaging
127 Spectroradiometer (MODIS) sensed data, suggesting that the dNBR approach is incapable of
128 retrieving reliable information with regards to burn severity (Roy et al. 2006). However,
129 markedly higher mean optimality measures were found for six burns in Alaska, USA (Murphy
130 et al. 2008). Also, the dNBR optimality statistics were found to outperform the dNDVI
131 optimality measures (Escuin et al. 2008) suggesting that the dNBR remains the most optimal
132 NDSI for estimating burn severity.

133 Several authors highlight the need for an independent validation of burn severity
134 assessments based on spectral indices for specific regions and vegetation types (Cocke et al.
135 2005; Key et al. 2005; Lentile et al. 2006 ; Chuvieco and Kasischke 2007; Fox et al. 2008).
136 As the technique is conceptually and computationally easy, burn severity maps based on
137 spectral indices could form an important instrument for post-fire management practices in the
138 fire-prone Mediterranean ecoregion. It is therefore our objective to evaluate different spectral
139 indices derived from Landsat TM imagery for assessing burn severity of the large 2007
140 Peloponnese wildfires in Greece. This general objective is fulfilled (i) by evaluating the
141 relationship between field data and several pre/post-fire differenced vegetation indices and (ii)
142 by comparing optimality statistics of those indices.

143 **Study area**

144 The area of interest is located at the Peloponnese, Greece (36°30'-38°30' N, 21°-23° E) (see
145 figure 1). Elevations range between 0 and 2404 m above sea level. Hot, dry summers alternate
146 with mild, wet winters resulting in a typical Mediterranean climate. For the Kalamata
147 meteorological station (37°4' N, 22°1' E) the mean annual precipitation equals 780 mm and

148 the average annual temperature is 17.8 °C (Hellenic National Meteorological Service,
149 www.hnms.gr).

150 Large wildfires struck the area (Gitas et al. 2008) in the 2007 summer. The first large burn
151 initiated on 26/07/2007 and lasted until 01/09/2007. The fires devastated a large amount
152 (more than 100 000 ha) of coniferous forest, broadleaved forest, shrub lands (phrygana and
153 maquis communities) and olive groves. Black pine (*Pinus nigra*) is the dominant conifer
154 species. Phrygana is dwarf scrub vegetation (< 1 m), which prevails on dry landforms
155 (Polunin 1980). Maquis communities consist of sclerophyllous evergreen shrubs of 2-3 m
156 high. The shrub layer is characterised by e.g. Kermes oak (*Quercus coccifera*), Hungarian oak
157 (*Q. frainetto*), mastic tree (*Pistacia lentiscus*), sageleaf rockrose (*Cistus salvifolius*), hairy
158 rockrose (*C. incanus*), tree heath (*Erica arborea*), and thorny burnet (*Sarcopoterum*
159 *spinosum*). The olive groves consist of *Olea europaea* trees whereas oaks are the dominant
160 broadleaved species.

161 **Methods**

162 *Data and preprocessing*

163 For assessing burn severity of the summer 2007 Peloponnese fires two anniversary date
164 Landsat TM images (path/row 184/34) were used (23/07/2006 and 13/08/2008) (step 1 in
165 figure 2). The images were acquired in the summer, minimizing effects of vegetation
166 phenology and differing solar zenith angles. The images were subjected to geometric,
167 radiometric, atmospheric and topographic correction (step 2 in figure 2).

168 The 2008 image was geometrically corrected using 34 ground control points (GCPs),
169 recorded in the field with a Garmin eTrex Vista GPS (Global Positioning System) (15 m error
170 in x and y under ideal condition (Garmin 2005), but up to 35.5 m under closed canopy
171 (Chamberlain 2002)). The resulting Root Mean Squared Error (RMSE) was lower than 0.5

172 pixels. The 2006 and 2008 images were co-registered within 0.5 pixels accuracy. All images
 173 were registered in Universal Transverse Mercator (zone 34S), with ED 50 (European Datum
 174 1950) as geodetic datum.

175 Raw digital numbers (DNs) were scaled to at-sensor radiance values (L_s) (Chander et al.
 176 2007) but with band-specific parameters proposed for Landsat TM data processed and
 177 distributed by the ESA (European Space Agency) (Arino et al. s.d.). The radiance to
 178 reflectance conversion was performed using the COST method (Chavez 1996):

$$179 \quad \rho_a = \frac{\pi(L_s - L_d)}{(E_o / d^2)(\cos \theta_z)^2} \quad (1)$$

180 where ρ_a is the atmospherically corrected reflectance at the surface; L_s is the at-sensor
 181 radiance ($\text{Wm}^{-2}\text{sr}^{-1}$); L_d is the path radiance ($\text{Wm}^{-2}\text{sr}^{-1}$); E_o is the solar spectral irradiance
 182 (Wm^{-2}); d is the earth-sun distance (astronomical units); and θ_z is the solar zenith angle. The
 183 COST method is a dark object subtraction (DOS) approach that assumes 1 % surface
 184 reflectance for dark objects (e.g. deep water). After applying the COST atmospheric
 185 correction, pseudo-invariant features (PIFs) such as deep water and bare soil pixels, were
 186 examined in the images. No further relative normalization between the images was required.

187 It was necessary to correct for different illumination effects due to topography. This was
 188 done based on the C correction method, an empirical modification of the cosine correction
 189 approach (Teillet et al. 1982), using a digital elevation model (DEM) and knowledge of the
 190 solar zenith and azimuth angle at the moment of image acquisition. Topographical slope and
 191 aspect data were derived from 90 m SRTM (Shuttle Radar Topography Mission) elevation
 192 data (Jarvis et al. 2006) resampled and coregistered with the Landsat images. The illumination
 193 is modeled as:

$$194 \quad \cos \gamma_i = \cos \theta_p \cos \theta_z + \sin \theta_p \sin \theta_z \cos(\phi_a - \phi_o) \quad (2)$$

195 where γ_i is the incident angle (angle between the normal to the ground and the sun rays); θ_p
196 is the slope angle; θ_z is the solar zenith angle; ϕ_a is the solar azimuth angle; and ϕ_o is the
197 aspect angle. Then terrain corrected reflectance ρ_t is defined as:

$$198 \quad \rho_t = \rho_a \left(\frac{\cos \theta_z + c_k}{\cos \gamma_i + c_k} \right) \quad (3)$$

199 where c_k is a band specific parameter $c_k = b_k/m_k$ where b_k and m_k are the respective
200 intercept and slope of the regression equation $\rho_a = b_k + m_k \cos \gamma_i$. Since topographic
201 normalization works better when applied separately for specific land cover types (Bishop and
202 Colby 2002) burned area specific c-values were calculated by masking the unburned areas
203 using a two-phase threshold method (Veraverbeke et al. in press).

204 To assess burn severity in the field, 160 GeoCBI plots were collected one year post-fire, in
205 September 2008. The GeoCBI is a modified version of the Composite Burn Index (CBI) (De
206 Santis and Chuvieco 2009). The (Geo)CBI is an operational tool used in conjunction with the
207 Landsat dNBR approach to assess burn severity in the field (Key and Benson 2005). The
208 GeoCBI divides the ecosystem into five different strata, one for the substrates and four
209 vegetation layers. These strata are: (i) substrates, (ii) herbs, low shrubs and trees less than 1
210 m, (iii) tall shrubs and trees of 1 to 5 m, (iv) intermediate trees of 5 to 20 m and (v) big trees
211 higher than 20 m. The strata are grouped in the understory (i-iii) and the overstorey (iv-v). In
212 the field form, 20 different factors can be rated (e.g. soil and rock cover/colour change, %
213 LAI change, char height) (see table 1) but only those factors present and reliably rateable, are
214 considered. The rates are given on a continuous scale between zero and three and the resulting
215 factor ratings are averaged per stratum. Based on these stratum averages, the GeoCBI is
216 calculated in proportion to their corresponding fraction of cover, resulting in a weighted
217 average between zero and three that expresses burn severity.

218 The 160 sample points were selected based on a stratified sampling approach, taking into
219 account the constraints on mainly accessibility and time, which encompasses the whole range
220 of variation found within the burns. Contributing to this objective 10 out of the 160 plots were
221 measured in unburned land, with a consequent GeoCBI value of zero. The field plots consist
222 of 30 by 30 m squares, analogous to the Landsat pixel size. The pixel centre coordinates were
223 recorded based on one measurement with a handheld Garmin eTrex Vista GPS device. To
224 minimize the effect of potential misregistration plots were at least 90 m apart and chosen in
225 relatively homogeneous areas of at least 60 by 60 m, although preferably more (Key and
226 Benson 2005). This homogeneity refers both to the fuel type and the fire effects. Of the 160
227 field plots 67 plots were measured in shrub land, 58 in coniferous forest, 17 in broadleaved
228 forest and 18 in olive groves. Figure 3 shows example low, moderate and high severity plot
229 photographs for the coniferous forest fuel type.

230 *Spectral indices and optimality*

231 In this study the potential of three Normalized Difference Spectral Indices (NDSIs) for
232 assessing fire-induced vegetation change is evaluated using TM bands most sensitive to post-
233 fire reflectance changes: TM3 (630-690 nm), TM4 (760-900 nm), TM5 (1550-1750 nm) and
234 TM7 (2080-2350 nm). Reflectance in the visual (TM3) and mid infrared (TM5 and TM7)
235 regions increases after fire, while the NIR region (TM4) is characterised by a reflectance drop
236 (Pereira et al. 1999). To capture this information, The Normalized Difference Vegetation
237 Index (NDVI) combines R (TM3) band with NIR (TM4) band information whereas the
238 Normalized Difference Moisture Index (NDMI) (Wilson and Sader 2002) and the Normalized
239 Burn Ratio (NBR) combine the NIR (TM4) band with a MIR (TM5 and TM7, respectively)
240 band. The NBR has become the standard spectral index for assessing fire/burn severity,
241 especially in North American regions, whereas the NDMI has not been evaluated before for

242 fire/burn severity applications. Nevertheless, it has been suggested that TM5 is well suited for
 243 remote sensing of canopy water content (Tucker 1980). Consequently it might also reflect
 244 post-fire reflectance changes and was included in this study. These are the formulas of the
 245 spectral indices used (steps 3 and 4 in figure 2):

$$246 \quad NDVI = \frac{TM\ 4 - TM\ 3}{TM\ 4 + TM\ 3} \quad dNDVI = NDVI_{pre} - NDVI_{post} \quad (4)$$

$$247 \quad NDMI = \frac{TM\ 4 - TM\ 5}{TM\ 4 + TM\ 5} \quad dNDMI = NDMI_{pre} - NDMI_{post} \quad (5)$$

$$248 \quad NBR = \frac{TM\ 4 - TM\ 7}{TM\ 4 + TM\ 7} \quad dNBR = NBR_{pre} - NBR_{post} \quad (6)$$

249 For evaluating the optimality of the bi-temporal change detection, the TM4-TM3, TM4-
 250 TM5 and TM4-TM7 bi-spectral spaces were considered (see figure 4). If a spectral index is
 251 appropriate to the physical change of interest, in this case fire-induced vegetation depletion,
 252 there exists a clear relationship between the change and the direction of the displacement in
 253 the bi-spectral feature space (Verstraete and Pinty 2006). In an ideal scenario a pixel's bi-
 254 temporal trajectory is perpendicular to the first bisector of the Cartesian coordinate system.
 255 This is illustrated in figure 4 for the displacement from unburned (U) to optimally (O) sensed
 256 burned. However, in practice perturbing factors such as atmosphere and illumination decrease
 257 the index performance. For example, in figure 4, a pixel displaces from unburned (U) to
 258 burned (B) after fire. Here, the magnitude of change to which the index is insensitive is equal
 259 to the Euclidian distance $|OB|$. Thus the observed displacement vector UB can be
 260 decomposed in the sum of the vectors UO and OB, hence, the index optimality is defined as
 261 (Roy et al. 2006):

$$262 \quad optimality = 1 - \frac{|OB|}{|UB|} \quad (7)$$

263 As $|OB|$ can never be larger than $|UB|$, the optimality measure varies between zero and
264 one. If the optimality measure equals zero, then the index is completely insensitive to the
265 change of interest. An optimality score of one means that the index performs ideal for
266 monitoring the change of interest.

267 Due to the non-linearity of the relationship between field and spectral indices estimates of
268 burn severity (Zhu et al. 2006, Hall et al. 2008), second-degree polynomial regressions were
269 performed to correlate the spectral indices (independent variables) and GeoCBI field data of
270 burn severity (dependent variables). Regression model results were compared using two
271 goodness-of-fit measures: the coefficient of determination R^2 and the Root Mean Squared
272 Error (RMSE). The coefficient of determination is an estimate of the proportion of the total
273 variation in the data that is explained by the model. The RMSE is a measure of how much a
274 response variable varies from the model predictions, expressed in the same units as the
275 dependent data. The RMSE describes how far points diverge from the regression line. In
276 addition, optimality statistics of all burned pixels were compared for the different indices. The
277 median statistic was used for this purpose because of its robustness to outlier values and
278 because the optimality distribution functions appeared to be non-normal.

279 **Results**

280 *Correlation with field data*

281 The distribution plots and regression lines of the GeoCBI and pre/post-fire differenced
282 spectral indices are displayed in figures 5D, 6E and 6F. Comparison of the R^2 statistics shows
283 that the GeoCBI-dNBR relationship proved to be the strongest. This relationship yielded a
284 moderate-high $R^2 = 0.65$ for a polynomial fitting model. This was followed by the GeoCBI-
285 dNDMI correlation which had an $R^2 = 0.50$. The GeoCBI-dNDVI relationship was the
286 weakest ($R^2 = 0.46$). The decreasing trend in R^2 statistic is at the same time associated with an

287 increasing RMSE (0.35, 0.42 and 0.44 for the relationships between the GeoCBI and
288 respectively dNBR, dNDMI and dNDVI data). The spectral index values of the dNBR
289 approach clearly range more than those of the dNDMI and dNDVI approaches. The within-
290 burn dNBR range almost doubles the within-burn dNDVI range. Most field plots have dNBR
291 values ranging from 0 and 0.8 (see figure 5F) and dNDMI and dNDVI between 0 and 0.5 (see
292 figures 5D and 5E). Figures 5A, 5B and 5C depict respectively the dNDVI, dNDMI and
293 dNBR maps. The dNBR map clearly reveals more contrast in the burnt areas than the other
294 maps.

295 *Index optimality*

296 Figures 6A-C depict the dNDVI, dNDMI and dNBR optimality maps of the burned areas. The
297 dNBR index (median = 0.51) outperformed the dNDMI and dNDVI indices (medians of
298 respectively 0.50 and 0.40), whereas the dNDMI provided better results than the dNDVI. The
299 performance differences are also reflected when the respective histograms are inspected (see
300 figures 6D-F). A large number of pixels have a dNDVI optimality lower than 0.1 and the
301 number of pixels steadily decreased with increasing dNDVI optimality. The dNDMI
302 histogram is more equally distributed. Although many pixels have dNBR optimality scores
303 above between 0.2 and 0.4 we can observe a slightly increasing trend in terms of number of
304 pixels when dNBR optimality increases. According to the non-parametric Wilcoxon test
305 (Hollander and Wolfe 1999) differences in median optimality and distribution functions are
306 statistically significant ($p < 0.001$).

307 **Discussion**

308 The dNBR approach gave the overall best correlation with GeoCBI field data followed by the
309 dNDMI and the dNDVI approach. Indices with a mid infrared spectral band yielded better
310 results than indices lacking a MIR band. This corroborates with earlier research findings:

311 AVHRR (Advanced Very High Resolution Radiometer) spectral indices based on the NIR
312 and MIR channels had a higher discriminatory potential for burned surface mapping than
313 indices based on the NIR and red channels (Pereira 1999), the importance of the MIR region
314 for burned shrub-savannah discrimination with MODIS (Moderate Resolution Imaging
315 Spectroradiometer) data has been demonstrated (Trigg and Flasse 2001) and significant post-
316 fire spectral changes occurred in the 1500-2500 nm region using hyperspectral AVIRIS
317 (Airborne Visible and Infrared Imaging Spectroradiometer) data (van Wagtenonk et al.
318 2004). In previous studies assessing the correlation between several spectral indices and CBI
319 field data the NBR was ranked as the first index in pre/post-burn approaches (Epting et al.
320 2005). For fires in several regions in the USA dNBR yielded higher correlations than dNDVI
321 (Zhu et al. 2006). In this report the within-burn range of dNDVI values was about half the
322 within-burn range of dNBR values, which is similar to our results. They also concluded that
323 dNDVI was more influenced by hazy remote sensing conditions due to the elevated potential
324 of atmospheric scattering in the red spectral region. Overall results show a moderate-high
325 correlation between GeoCBI field data and dNBR for this case study in a Mediterranean
326 environment. Polynomial fitting models resulted in $R^2 = 0.65$. These outcome falls within the
327 range of results of previous studies (French et al. 2008).

328 In studies based on the spectral index theory the dNBR had a higher mean optimality
329 (0.49) than the dNDVI (0.18) based on Landsat TM/ETM+ images (Escuin et al. 2008). Our
330 results approximate to the values reported in similar studies of 0.49 (Escuin et al. 2008) and
331 ranging from 0.26 to 0.8 for six burns in Alaska, USA (Murphy et al. 2008). However, results
332 contrast with the very low mean dNBR optimality scores (0.1) based on Landsat ETM+
333 imagery for African savannah burns (Roy et al. 2006). These authors also report low dNBR
334 optimality values for MODIS sensed fires in other ecosystems (Russia, Australia and South
335 America). These results suggest that the dNBR index is to a high degree suboptimal for

336 assessing burn severity. These poor optimality results, however, can possibly be explained by
337 the fact that Roy et al. (2006) included unburned pixels in their optimality analysis.
338 Unaffected pixels are generally associated with low optimality scores as a pixel's
339 displacement in the bi-spectral space is only due to the noise (Escuin et al. 2008). This
340 explains the low optimality values reported (Roy et al. 2006).

341 The NDMI based approach, which had not been evaluated before for estimating burn
342 severity, performed better than the NDVI based approach. However, the NBR outperformed
343 the NDMI. This can be explained by the typically lower pre-fire reflectances in Landsat TM
344 band 7 (2080-2350 nm) than in Landsat TM band 5 (1550-1750 nm) due to a higher degree of
345 water absorption by vegetation at longer wavelengths. Therefore fire-induced reflectance
346 increase is likely to be more explicit in TM7 than in TM5. As a result, an index with TM7
347 instead of TM5 is able to capture a larger range of variation in post-fire effects.

348 Apart from the fact that the dNBR outperformed the dNDMI and dNDVI, use of the dNBR
349 to indicate burn severity is still problematic. When the GeoCBI-dNBR scatter plot and
350 regression line (see figure 5F) are examined, three points of defectiveness attract attention: (i)
351 the insensitivity of the regression model to unburned pixels, (ii) the saturation of the model
352 for GeoCBI values higher than approximately 2.5, and (iii) the moderately high dispersion of
353 the point cloud around the fitting line. First, the regression line crosses the x-axis at $dNBR = -$
354 0.23 while the unburned reference plots are situated closer to $dNBR = 0$. According to the
355 regression equation (see figure 5F) an unburned plot with a dNBR value of zero would be
356 associated with a GeoCBI value of 0.91, which is a clear overestimation of severity.
357 Secondly, the regression model reveals asymptotic behaviour for GeoCBI values higher than
358 2.5. As a consequence the empirical model potentially underestimates high severity plots and
359 is not able to differentiate between them. This phenomenon was also reported in previous
360 studies (e.g. van Wagtenonk et al. 2004; Epting et al. 2005). As a solution for the

361 insensitivity to unburned pixels and the saturation problem, a non-linear model based on a
362 saturated growth model was proposed (Hall et al. 2008). This model effectively handled the
363 insensitivity and saturation problems, however, at the expense of a lower R^2 and a higher
364 RMSE. Thirdly, the GeoCBI-dNBR model has a RMSE of 0.35, which is about one ninth of
365 the total GeoCBI variation. The observed GeoCBI values thus substantially diverge from the
366 model predictions.

367 Potential sources of inaccuracy arise from both the field and satellite levels. For example,
368 67 GeoCBI plots were measured in shrub land to fulfill the need for a stratified sampling
369 approach that requests that the number of plots of each fuel type is in proportion to the total
370 area burned of each pre-fire land cover type. However, as is known (e.g. van Wagendonk et
371 al. 2004; Epting et al. 2005), the CBI approach underperforms in non-forested areas. Part of
372 the observed inaccuracy can also be explained by the fact that that both field and satellite data
373 are imperfect proxies of burn severity. The CBI is based on semi-quantitative judgement
374 procedure and therefore possibly lacks absoluteness, while several noise factors hamper
375 satellite image analysis.

376 The amount of noise in the dNBR approach appeared to be fairly high as the median dNBR
377 optimality of 0.51 is considerably lower than the optimality of 1. An important part of the
378 spectral change in the TM4-TM7 bi-spectral space occurs parallel to the NBR isolines (confer
379 distance $|OB|$ in figure 4). Deficient preprocessing (no or unsatisfactory atmospheric
380 correction, topographic correction, image-to-image normalization...) can introduce noise in a
381 remote sensing analysis. The application of these procedures in burn severity applications is
382 sometimes blurred (French et al. 2008), although its importance has already been
383 demonstrated for example by revealing the effect of illumination on index values (Verbyla et
384 al. 2008).

385 These findings can direct the burn severity research in different directions. First, a
386 thorough review of the influence of preprocessing steps (especially atmospheric and
387 topographic correction) on dNBR performance is suggested. Secondly, it is desired to
388 improve the index design towards an index whose isolines are oriented to realize a higher
389 degree of sensitivity to burn severity while providing insensitivity to other sources of spectral
390 variation. These first two research directions retain the conceptual ease of the spectral indices
391 approach. A third alternative could focus on the further development of more advanced
392 remote sensing techniques into operational use. In this context, radiative transfer models
393 (Chuvieco et al. 2006; De Santis and Chuvieco 2007; De Santis et al. 2009) and spectral
394 mixture analysis (Lewis et al. 2007) have already proven to have big potential.

395 **Conclusions**

396 Results of the field data and optimality based analyses confirm one another, demonstrating
397 that the dNBR approach was the best index of the three spectral indices tested for estimating
398 burn severity in this case study in a Mediterranean environment. Results, however, also
399 indicate that the dNBR approach suffers from some striking inaccuracies. The empirical fit
400 between field and remotely sensed data is subject for improvement while the mean dNBR
401 optimality score was markedly lower than the ideal scenario with optimality values of one.
402 Further research in burn severity mapping should therefore focus on (i) noise removal (e.g. by
403 improved preprocessing), (ii) improved index design and (iii) alternative methods such as
404 radiative transfer models and spectral unmixing.

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611 **Fig. 1.** Location of the study area and distribution of the field plots (Landsat TM image
612 13/08/2008, UTM 34S ED50).

613 **Fig. 2.** Methodological workflow.

614 **Fig. 3.** Example photographs of a high, moderate and low severity plot in coniferous forest.

615 **Fig. 4.** Example pre/post-fire trajectory of a pixel in the TM4-TM3, TM4-TM5 or TM4-TM7
616 feature space. A pixel displaces from unburned (U) to burned (B). The index (NDVI, NDMI
617 or NBR) is sensitive to the displacement $|UO|$ and insensitive to the displacement $|OB|$.

618 **Fig. 5.** dNDVI, dNDMI and dNBR maps (a,b and c) and scatter plots and regression lines for
619 the GeoCBI-dNDVI (d), GeoCBI-dNDMI (e) and GeoCBI-dNBR (f) relationships.

620 **Fig. 6.** dNDVI (a and d), dNDMI (b and e) and dNBR (c and f) optimality maps and
621 histograms.

622 **Table 1.** GeoCBI criteria used to estimate fire/burn severity in the field (after De Santis and
623 Chuvieco 2009).