Earthquake stress drop estimates: What are they telling us?

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Lots of data for big earthquakes (rupture dimensions, slip history, etc.)

Small earthquakes are only observed from seismograms; no direct measurements of physical properties

## Two parameters

displacement  $= D$ 

 $\overline{\phantom{a}}$ 

 $area = A$ 



shear modulus



**Stress drop** 
$$
\Delta \sigma = \sigma_{\text{final}} - \sigma_{\text{initial}}
$$

average shear stress on fault

fault area average



displacement

## Circular crack model  $\Delta \sigma =$  $7 \pi \mu D = 7 M_0$  $16 r$   $16 r^3$ = average displacement *r D*

 $M_0 = \mu AD = \mu \pi r^2 D$ 

fault radius

Stress drop is proportional to displacement/radius ratio

(*Eshelby*, 1957; *Brune*, 1970)

# Seismology 101



In theory, far-field seismometer will record displacement pulse from small earthquake (can be either *P* or *S* wave), ignoring attenuation and other path effects

Area under displacement pulse *f*(*h*τ) is related to seismic moment  $M_0$  (one measure of event strength)

Pulse width  $\tau$  is related to physical dimension of fault, rise time, and rupture velocity



# Spectral Analysis 101

Time Series Spectrum







Original spectrum

$$
\Delta \sigma = \frac{7 M_0}{16 r^3}
$$

Assume circular crack model

cubed!

*r*

 $\cdot M_0$  Correct for geometrical spreading and radiation pattern

> Assume rupture velocity and source model (*Brune*, *Madariaga*, *Sato* & *Hirasawa*, *Kaneko & Shearer*, etc.)

## General Δσ results and issues

- $\Delta \sigma = 0.2$  to 20 MPa from corner frequency studies
- Much less than absolute shear stress levels predicted by Byerlee's law and rock friction experiments
- Little dependence of average  $\Delta \sigma$  on  $M_0$ , implying selfsimilar scaling of earthquakes, but possibility of small increase with  $M_0$  has been debated
- Some evidence that plate-boundary earthquakes have lower Δσ than mid-plate earthquakes
- Hard to compare  $\Delta\sigma$  results among studies because they often use different modeling assumptions and are based on small numbers of earthquakes





- Online database of seismograms, 1984–2003
- $\cdot$  > 300,000 earthquakes
- *P* and *S* multi-taper spectra computed for all records
- 60 GB in special binary format





Egill Hauksson

# Isolating Spectral Contributions



 $d_{ij} \approx e_i + s_j + x_{k(i,j)}$ 



- $\bullet$  > 60,000 earthquakes, >350 stations
- 1.38 million *P*-wave spectra (STN > 5, 5-20 Hz)
- Iterative least squares approach with outlier suppression

#### Assumed source model

• *Madariaga* (1976), *Abercrombie* (1995)



We fit data (solid lines) between 2 and 20 Hz, using:

$$
u(f) = \frac{\Omega_0}{1 + (f/f_c)^n}
$$

$$
f_c = \frac{0.42 \beta}{(M_0/\Delta \sigma)^{1/3}}
$$

(assumes rupture velocity =  $0.9 \beta$ )

Model prediction (dashed lines) is for  $\Delta\sigma = 1.60$  MPA (constant)

#### Calculated Earthquake Stress Drops



- $\cdot$  65,070 events
- $\bullet$  > 300,000 spectra
- 1989–2001
- $\bullet$  > 4 spectra/event
- 5 20 Hz band

 $Red = fewer high$ frequencies, lower stress drop or high near-source attenuation

 $Blue = more high$ frequencies, higher stress drop or low near-source attenuation

# Empirical Green's Function (EGF)



#### Subtract small event from big event to get estimate of true source spectrum for big event

## Source-specific EGF method

For each event, find 500 neighboring events:



Then subtract EGF from target event spectrum and compute  $\Delta\sigma$  for this event

#### Observed source Δσ using spatially varying EGF method



#### How variable are earthquake stress drops?



- Harder to resolve high Δσ events due to high corner frequencies
- Results are more reliable when more stations are stacked
- $\Delta \sigma$  = 0.2 to 20 Mpa
- $\sim$ 10x local scatter
- $\bullet$  ~10x regional variations

## Earthquake scaling



uniform scaling of all parameters (self similarity)

Constant Δσ



**or** 



Big Earthquake

**Small Earthquake** 

## Median stress drop does not vary with  $M_W$



## Stress drop versus depth



- Average  $Δσ$  increases from 0.6 to 2 MPa from 0 to 8 km
- But slower rupture velocities at shallow depths could also explain trend
- Nearly constant from 8 to 18 km
- Large scatter at all depths

#### Stress drop versus type of faulting

3895 high-quality focal mechanisms from J. Hardebeck (2005)



## Landers Aftershocks



- Along-strike changes in Δσ
- Related to mainshock slip?

Profiles for slip model of *Wald & Heaton* (1994)

## Comparison to Landers Slip Model



**Homestead Valley Fault** 

 $Red =$  low  $\Delta \sigma$ Blue = high  $\Delta\sigma$ 



Slip model from *Wald & Heaton* (1994)

#### Landers Slip Models



from www.seismo.ethz.ch/srcmod/

## Average Δσ (smoothed over 500 events)



- $\bullet$  0.5 to 5 MPa
- Coherent patterns
- What does it mean?
- Does this say anything about absolute stress?

## Conclusions for Southern California

- Stress drops range from 0.2 to 20 MPa for  $M_I = 1$  to 3.4 earthquakes, with no dependence on moment.
- Spatially coherent patterns in average stress drop  $(0.5)$ to 5 MPa), no consistent decrease near active faults.
- Shallow earthquakes radiate less high frequencies than deeper events, implying slower rupture velocities or lower stress drops.
- Landers aftershocks have strong along-strike variations in stress drop with possible correlation to slip models.
- Hard to resolve any temporal changes.

## 1989-2001 *b*-values



- Computed for each event and 500 nearest neighbors
- $M = 2$  to 4
- median  $b = 1.12$



#### not much correlation!

