

# Using local topography to constrain absolute stress (with some additional thoughts on heterogeneity)

Karen Luttrell  
Louisiana State University

## Acknowledgements

- Bridget Smith-Konter, University of Hawaii
- Jeanne Hardebeck, US Geological Survey
- All Contributors to the SCEC Community Stress Model

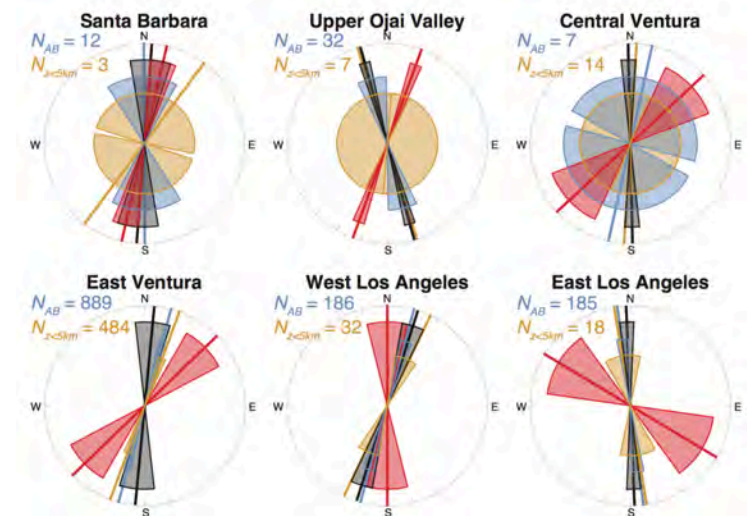
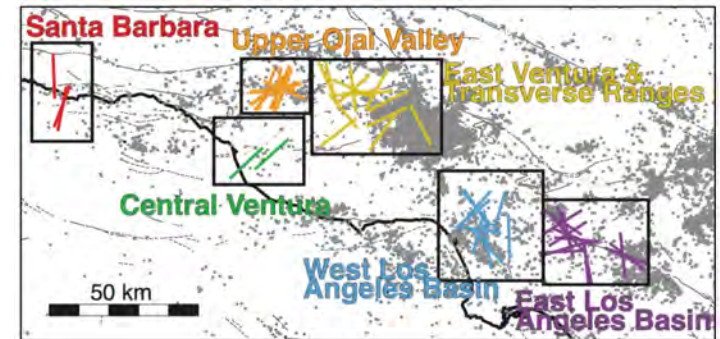
### LSU Students

Joel Spansel (BS)  
Erin Schwartz (BS)  
Phoenix Harris (BS)  
Elliott Helgans (MS)



# What we're not going to talk about

- Do borehole breakout SHmax agree with focal mechanism inversion SHmax?
  - (not really)
- Why not?
  - Depth heterogeneity?
  - Lateral heterogeneity?
  - Temporal heterogeneity?
- We're testing "depth" and "lateral" by doing new FM inversions with only local Eqs
  - Fits better in some places/ circumstances...
  - But in others, differences can't be explained by differences in spatial sampling...



Luttrell and Hardebeck, ongoing

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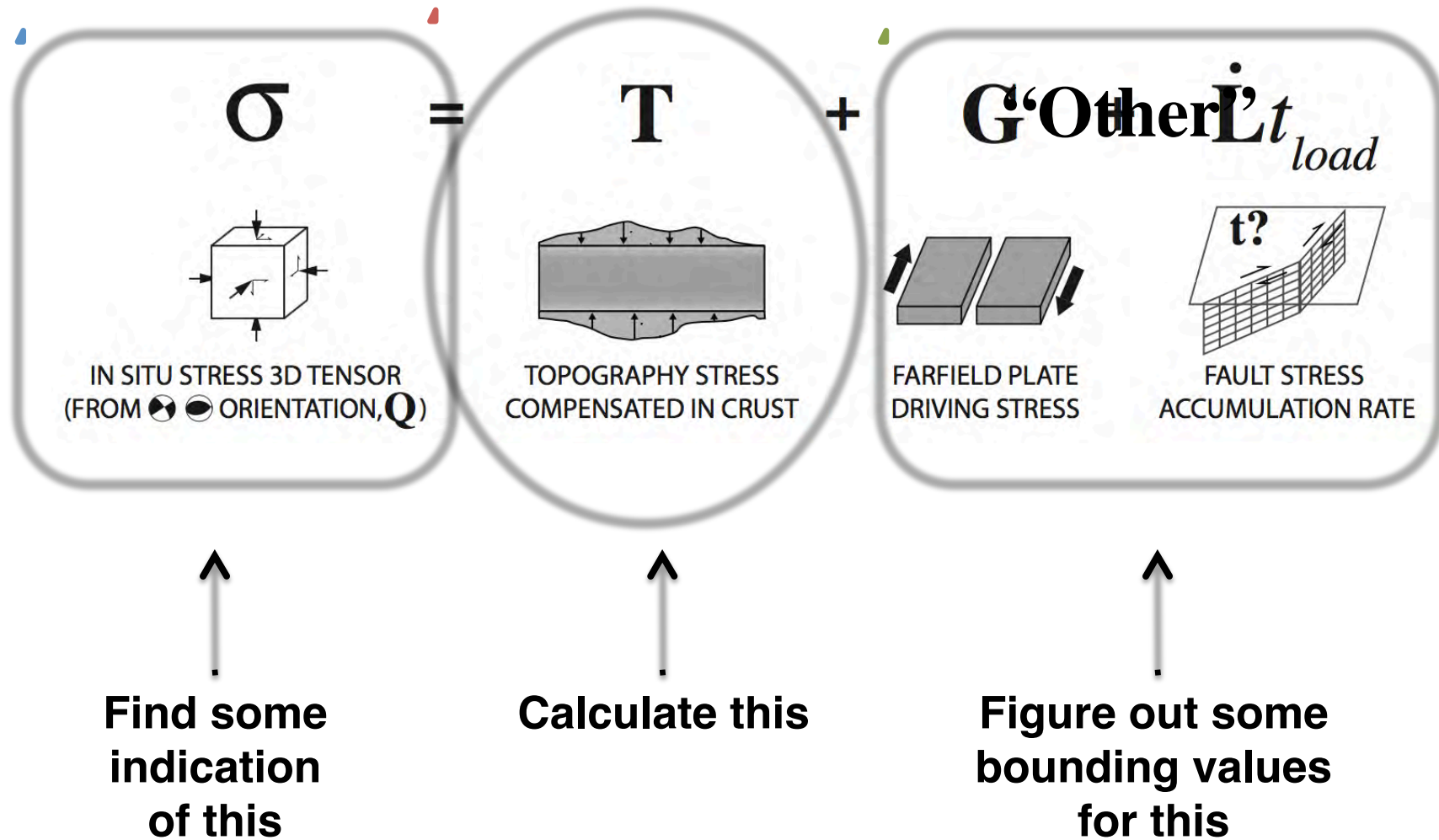
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# Simple forward model of stress field



# Previous Incarnations of Absolute Stress Constraint via Topography megathrust earthquake

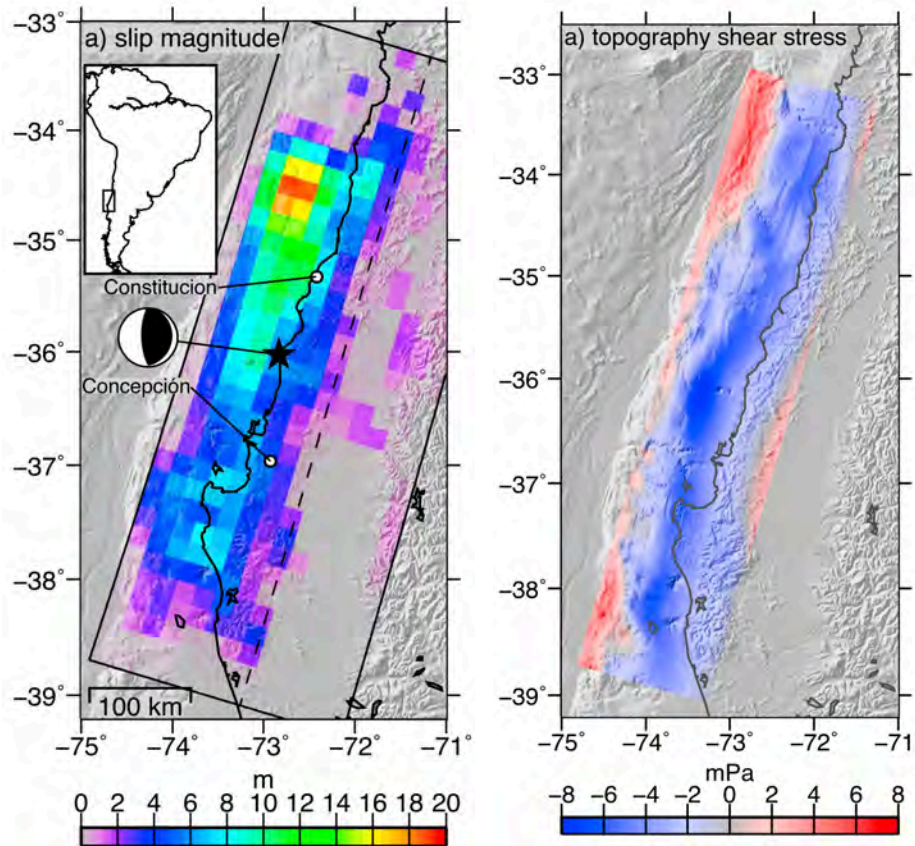
JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 116, B11401, doi:10.1029/2011JB008509, 2011

## Estimates of stress drop and crustal tectonic stress from the 27 February 2010 Maule, Chile, earthquake: Implications for fault strength

Karen M. Luttrell,<sup>1,2</sup> Xiaopeng Tong,<sup>1</sup> David T. Sandwell,<sup>1</sup> Benjamin A. Brooks,<sup>3</sup> and Michael G. Bevis<sup>4</sup>

Received 3 May 2011; revised 8 August 2011; accepted 20 August 2011; published 3 November 2011.

[1] The great 27 February 2010  $M_w$  8.8 earthquake off the coast of southern Chile ruptured a ~600 km length of subduction zone. In this paper, we make two independent estimates of shear stress in the crust in the region of the Chile earthquake. First, we



Compare forearc topography with slip direction to constrain driving stress and compare with stress drop

[Luttrell et al., 2011]

# Previous Incarnations of Absolute Stress Constraint via Topography mid-ocean ridges

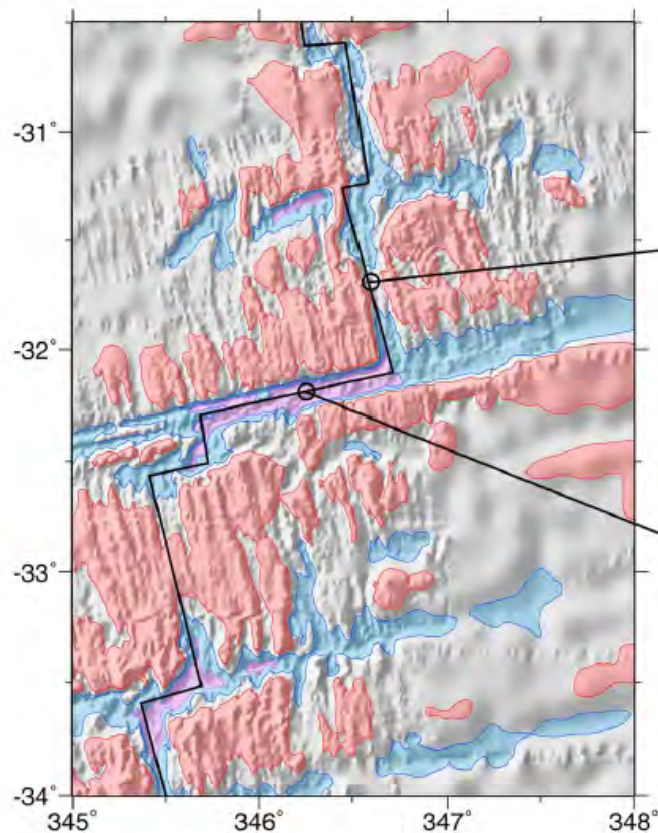
JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 117, B04402, doi:10.1029/2011JB008765, 2012

## Constraints on 3-D stress in the crust from support of mid-ocean ridge topography

Karen Luttrell<sup>1,2</sup> and David Sandwell<sup>1</sup>

Received 9 August 2011; revised 17 December 2011; accepted 19 February 2012; published 10 April 2012.

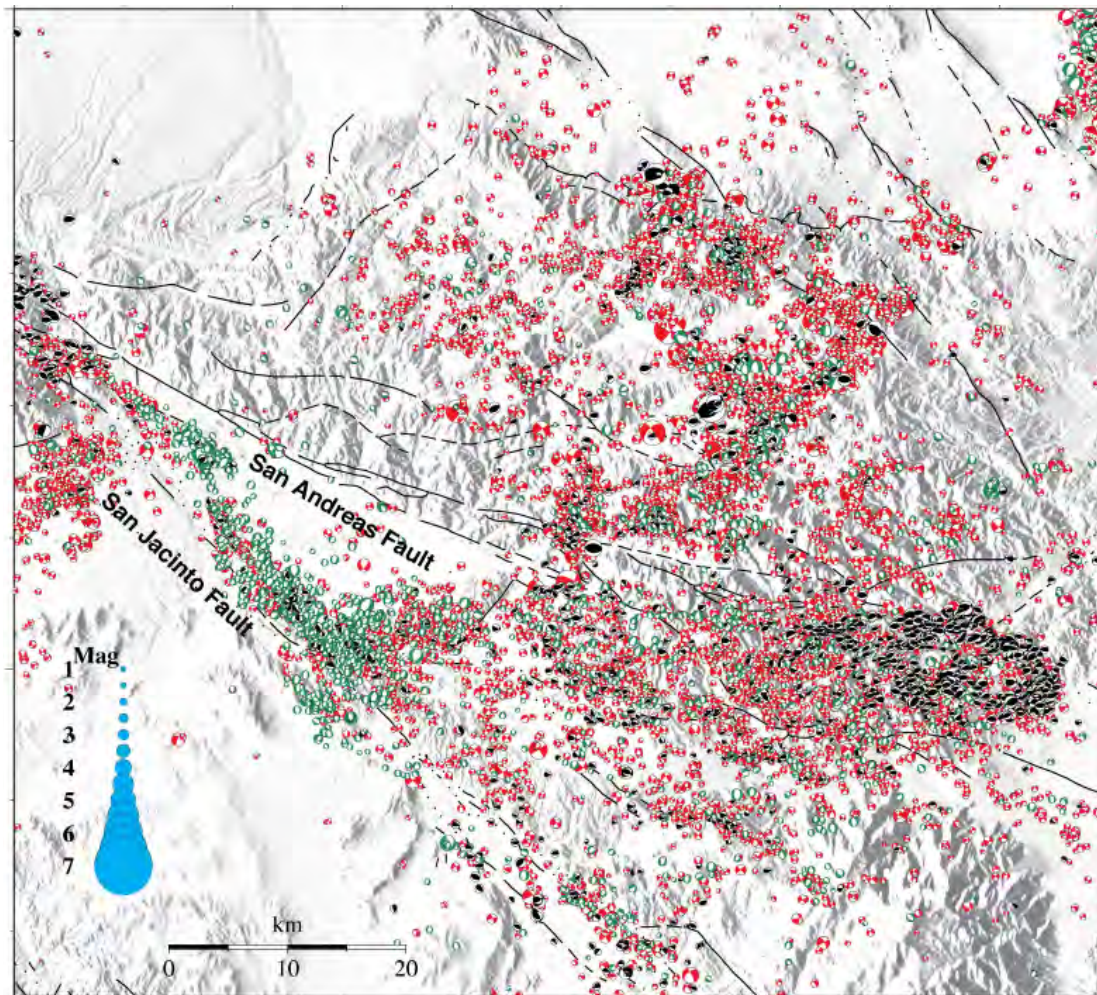
[1] The direction of crustal stresses acting at mid-ocean ridges is well characterized, but the magnitude of these stresses is poorly constrained. We present a method by which the absolute magnitude of these stresses may be constrained using seafloor topography and gravity. The topography is divided into a short-wavelength portion, created by rifting,



[Luttrell and Sandwell, 2012]

Fitting ridge highs/lows and transform lows/highs simultaneously with a single consistent 2-D stress field

# Southern California is a lot more complicated: Varied faulting types



[Yang et al. 2012]

# History of the CSM

## Last year of SCEC 3

September 14, 2011: Palm Springs **Workshop on Strategies for Implementing a Community Stress Model**

## SCEC 4: 2012-2016

October 15-16, 2012: USC **SCEC Community Workshop: Community Stress Model**  
May 29-30, 2013: Menlo Park **SCEC Community Workshop: Community Stress Model**  
October 27, 2014: Pomona **SCEC Community Workshop: Community Stress Model**  
September 13, 2015: Palm Springs **SCEC Community Stress Model (CSM) Workshop**

## SCEC 5: 2017-2021

January 15-16, 2019: Pomona **SCEC Community Stress Model (CSM) Workshop**





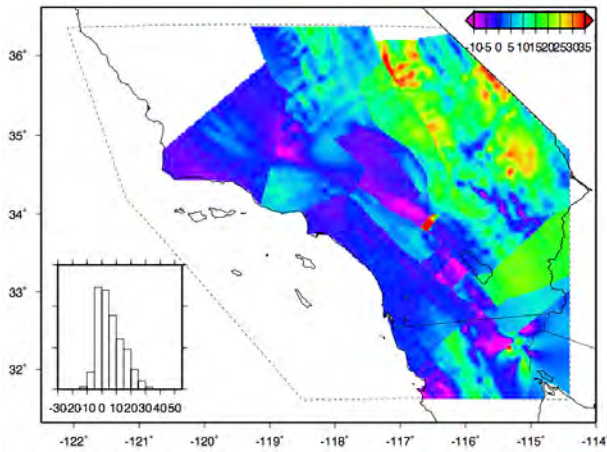
# 2012 Workshop: Initial Contributions

- **16 models submitted**

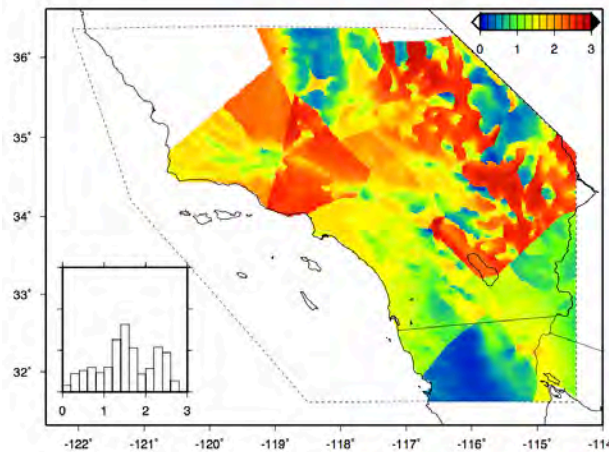
- Stress magnitude (3 models)

- Disagree on magnitude and magnitude of variation

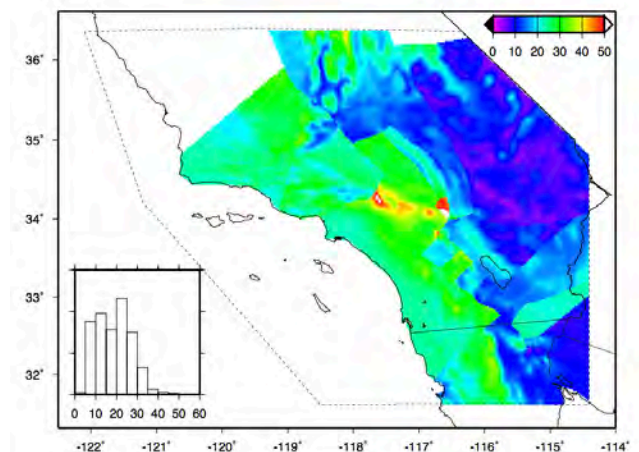
SHmax trend (degrees); depth=5 km



A\_phi; depth=5 km



differential stress (MPa); depth=5 km



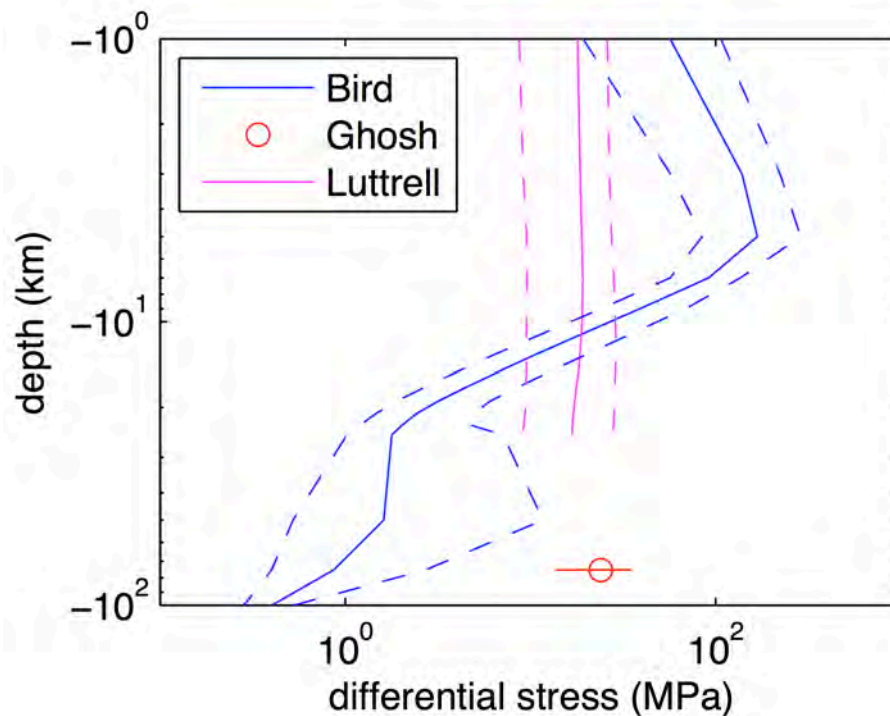
**Luttrell's contribution to CSM 2012: very very very very very preliminary**

# 2012 Workshop: Initial Contributions

- **16 models submitted**

- Stress magnitude (3 models)

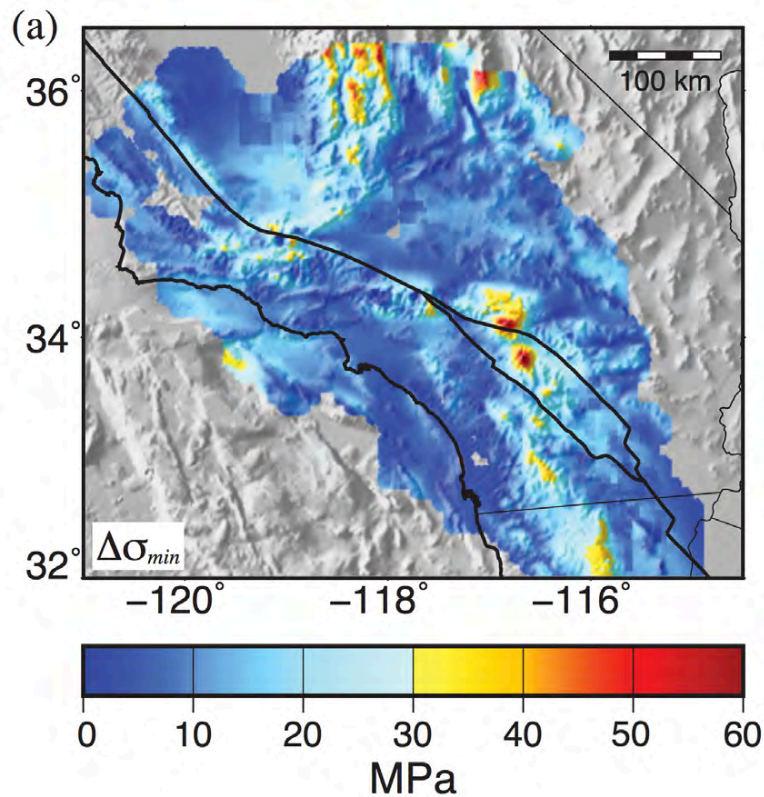
- Disagree on magnitude and magnitude of variation



**“hmmmm, maybe it’s  
worth trying to do this  
for real...”**

**- Karen’s internal monologue,  
2012**

# Eventually, after hitting lots of walls, a constraint on Absolute Stress via Topography



*Geophys. J. Int.* (2017) **211**, 472–482  
Advance Access publication 2017 July 22  
GJI Geodynamics and tectonics

doi: 10.1093/gji/ggx301

## Limits on crustal differential stress in southern California from topography and earthquake focal mechanisms

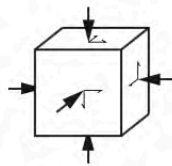
Karen Luttrell<sup>1</sup> and Bridget Smith-Konter<sup>2</sup>

<sup>1</sup>Department of Geology and Geophysics, Louisiana State University, E235 Howe-Russell, Baton Rouge, LA 70803, USA. E-mail: [kluttrell@lsu.edu](mailto:kluttrell@lsu.edu)

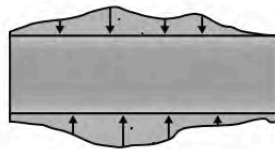
<sup>2</sup>Department of Geology and Geophysics, University of Hawaii, 1680 East-West Road, POST 813, Honolulu, HI 96822, USA

# Simple forward model of stress field

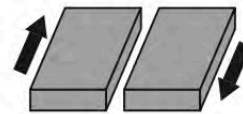
$$\sigma = T + \text{“Other”}$$



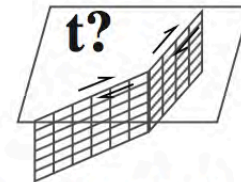
IN SITU STRESS 3D TENSOR  
(FROM  ORIENTATION, Q)



TOPOGRAPHY STRESS  
COMPENSATED IN CRUST



FARFIELD PLATE  
DRIVING STRESS



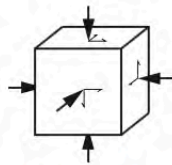
FAULT STRESS  
ACCUMULATION RATE



**Need some additional information or assumptions**

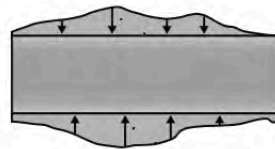
- 1. Assume topography is NOT dominant in Southern California**
- 2. Assume “other” is dominant in Southern California**  
**i.e., topography is ~negligible**

# Simple forward model of stress field

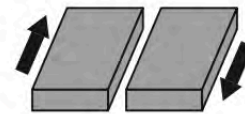
$$\sigma = T + \text{“Other”}$$



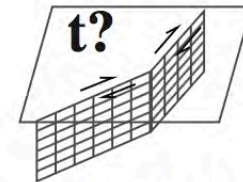
IN SITU STRESS 3D TENSOR  
(FROM   ORIENTATION,  $Q$ )



TOPOGRAPHY STRESS  
COMPENSATED IN CRUST



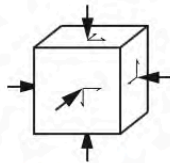
FARFIELD PLATE  
DRIVING STRESS



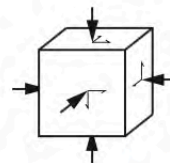
FAULT STRESS  
ACCUMULATION RATE



 $\sigma$ 
 $\approx$ 

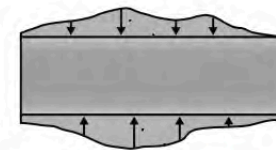
$$\sigma * \Delta\sigma - T$$



IN SITU STRESS 3D TENSOR  
(FROM   ORIENTATION,  $Q$ )



IN SITU STRESS 3D TENSOR  
(FROM   ORIENTATION,  $Q$ )



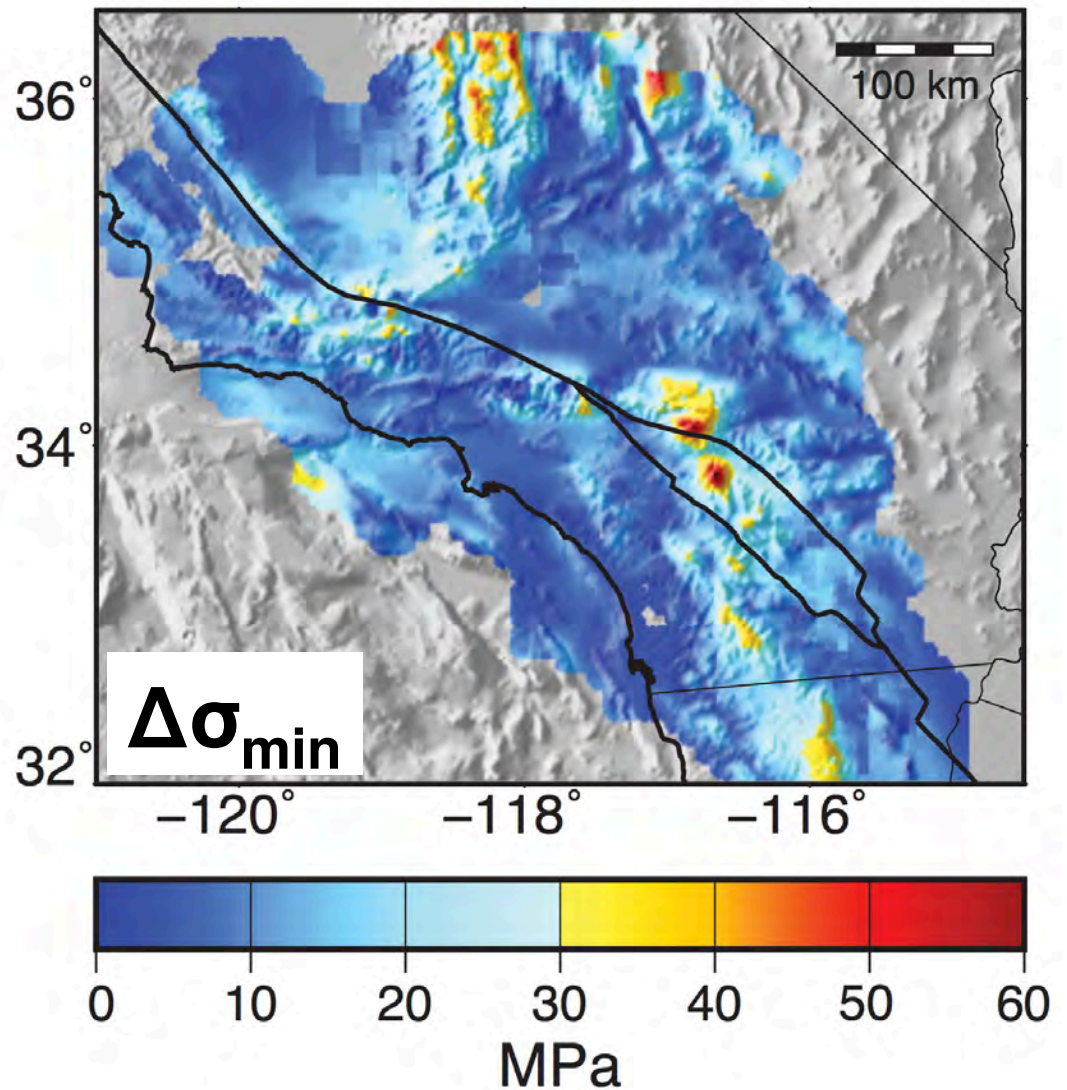
TOPOGRAPHY STRESS  
COMPENSATED IN CRUST

*orientation*

*orientation*

# Minimum in situ magnitude estimate: $\Delta\sigma_{\min}$

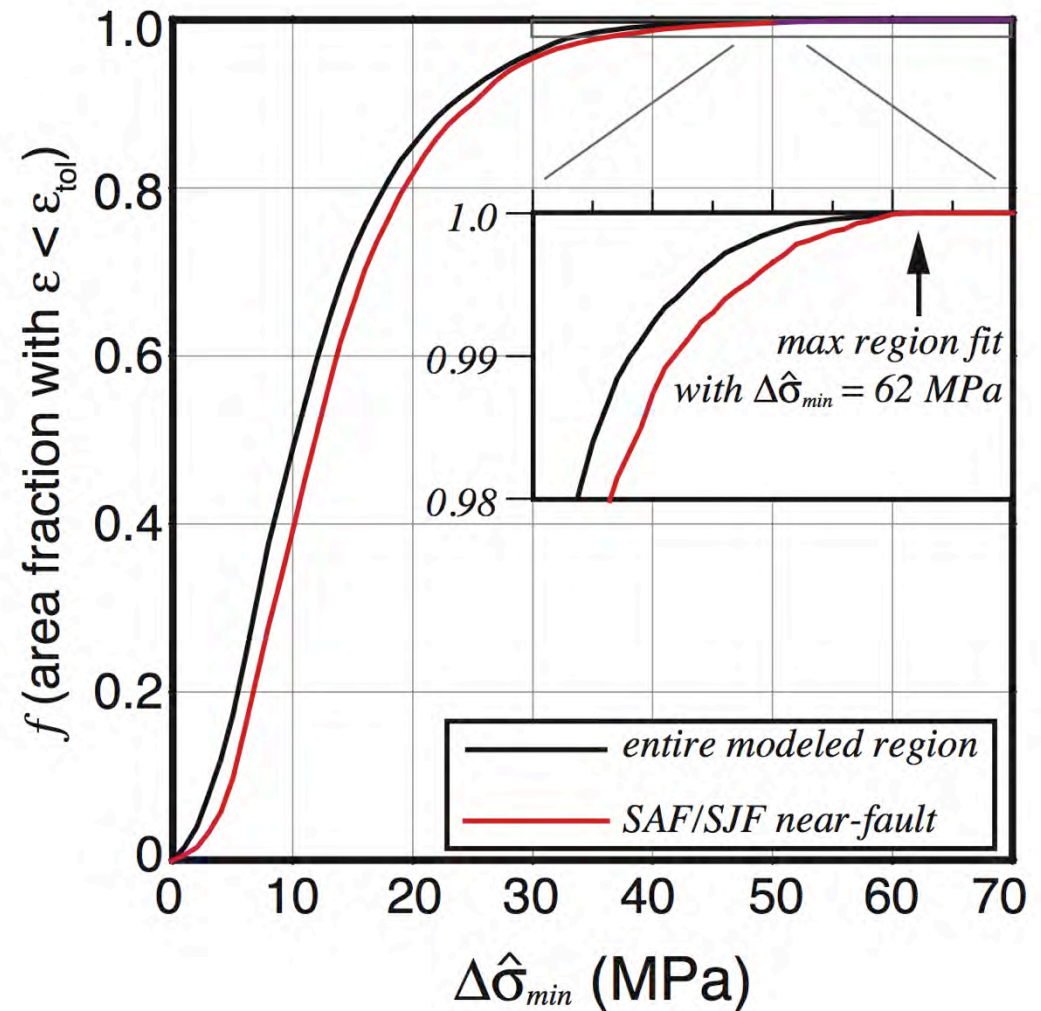
- $\Delta\sigma$  required to maintain *in situ* orientation to within  $\pm 15^\circ$ , despite resistance from topography
- Across SoCal, ranges from  $\sim 10 - 60$  MPa
- This is a lower bound: stress could be arbitrarily higher and fit just as well



[Luttrell and Smith-Konter, 2017]

# Minimum in situ magnitude estimate: $\Delta\sigma_{\min}$

- How does min  $\Delta\sigma$  estimate vary across region?
- CDF of area able to support existing topography for *in situ* differential stress of a certain magnitude
- Similar result if near-fault areas considered separately
- Most rugged topography requires  $\Delta\sigma$  of 62 MPa

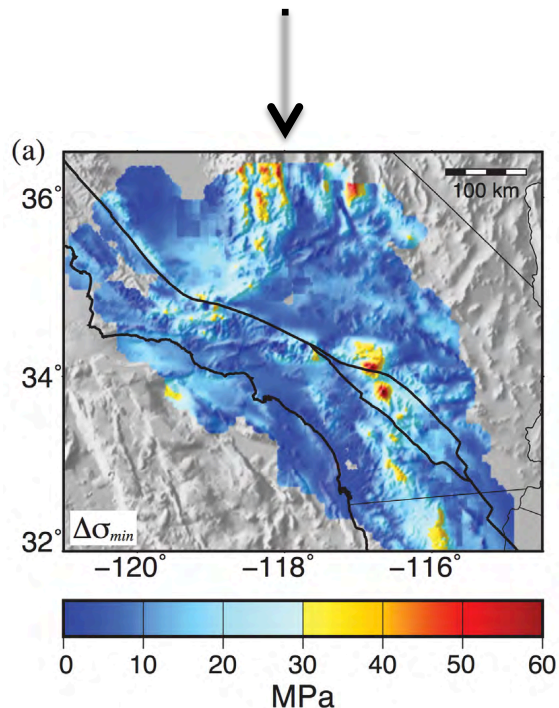


# Which estimate should we use for $\Delta\sigma_{\min}$ ?

Depends on how heterogeneous stress magnitude is...

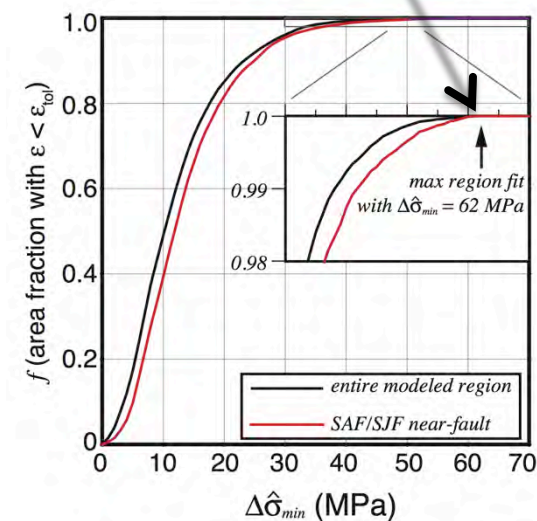
If variations are large relative to mean...

... this is the best estimate of  $\Delta\sigma_{\min}$  at each place



If variations are small relative to mean...

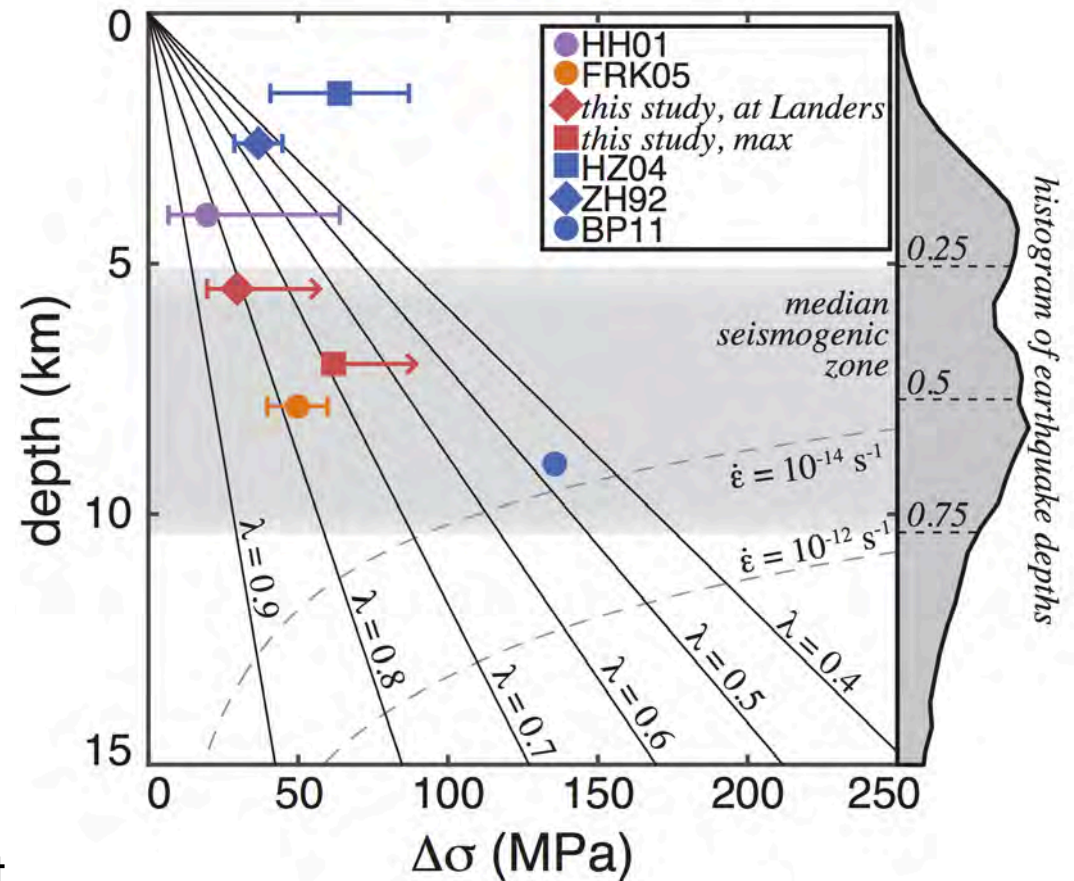
...  $\Delta\sigma_{\min}$  everywhere must be large enough to support max





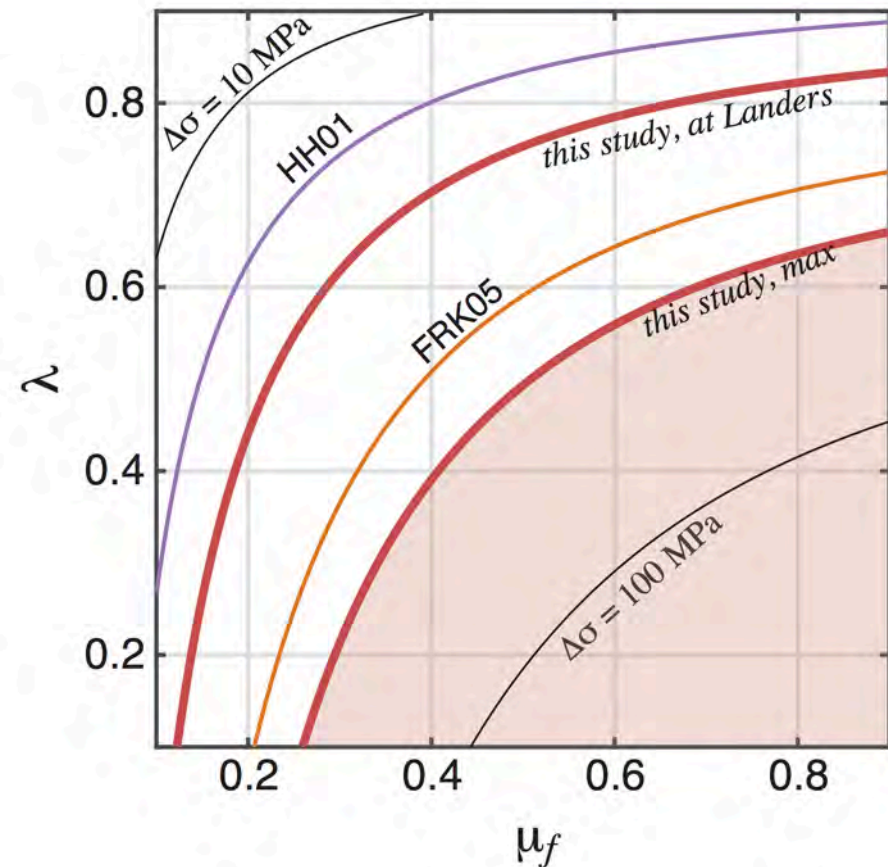
# Do these results make sense?

- Compare with estimates from
  - Shallower drilling
  - Deeper exhumed crustal rocks
  - Landers aftershock rotation
- Max required stress is concordant with shallower and deeper estimates
- Landers region is high, but within error bars
- YSE places a lower limit on fault friction and an upper limit on pore pressure



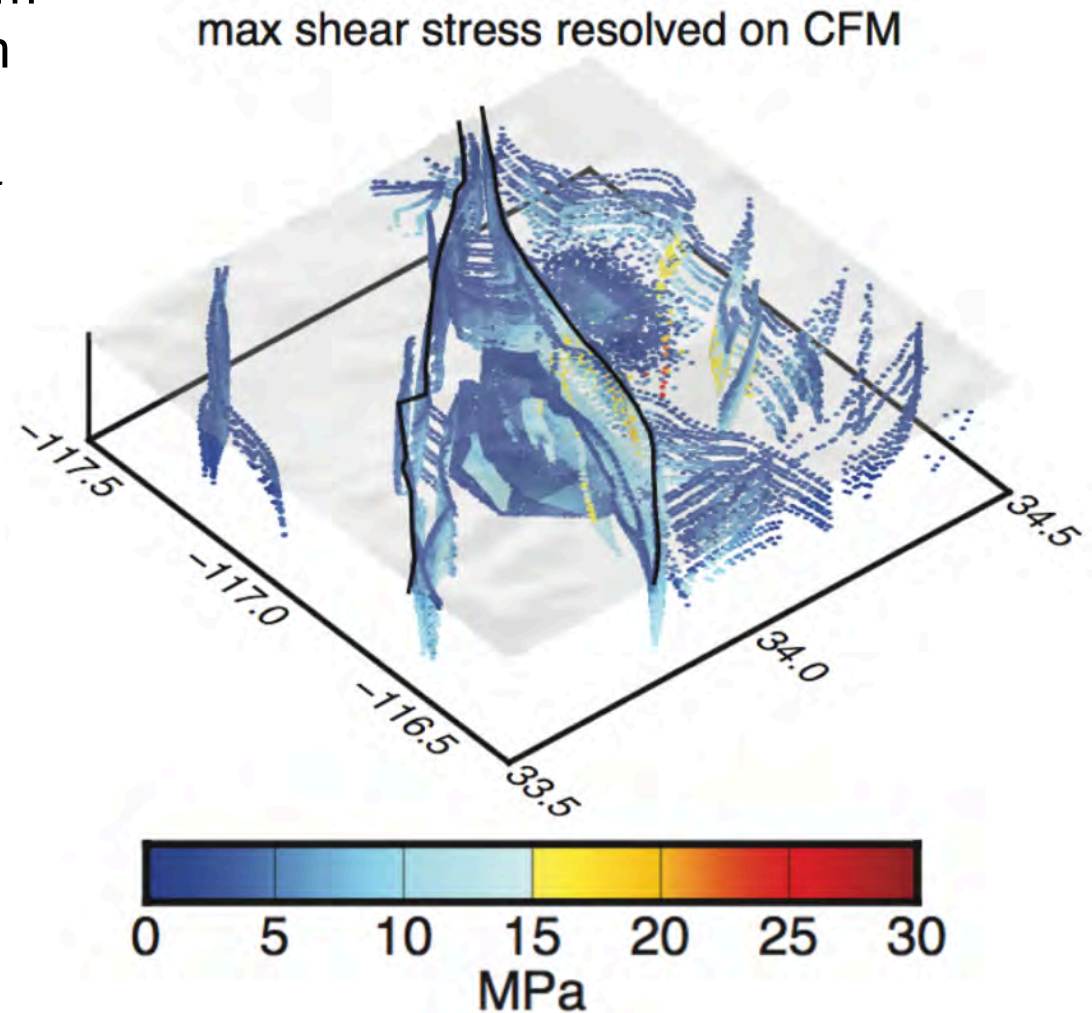
# Do these results make sense?

- YSE places a lower limit on fault friction and an upper limit on pore pressure
- At max required stress,
  - Fault friction can't be very low ( $\mu_f > 0.3$ )
  - Pore pressure can't be very high ( $\lambda < 0.7$ )
- Heterogeneous stress field more permissive



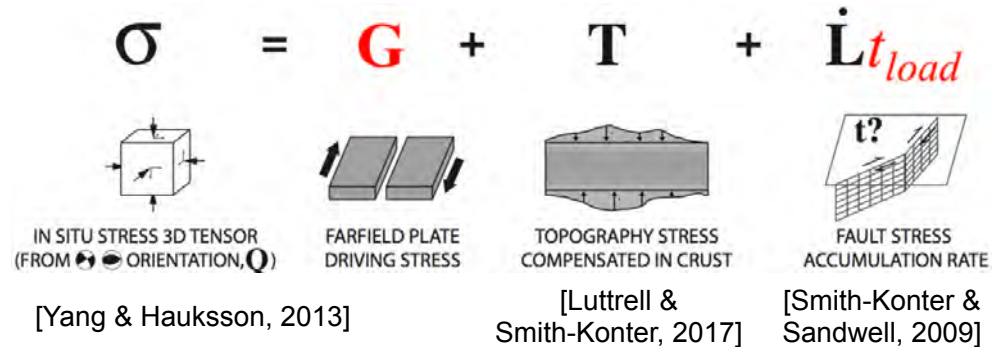
# How much shear stress on the faults?

- Maximum shear stress from depth dependent minimum stress field estimate
  - Based on *Luttrell and Smith-Konter, [2017]*
- Resolve on CFM planes
  - [*Plesch et al., 2007;*  
*Nicholson et al., 2013*]
- Gives the right rake (not surprising, orientation is mostly from focal mechanisms)
- Shear stress generally ~5-20 MPa, varies with depth and fault orientation

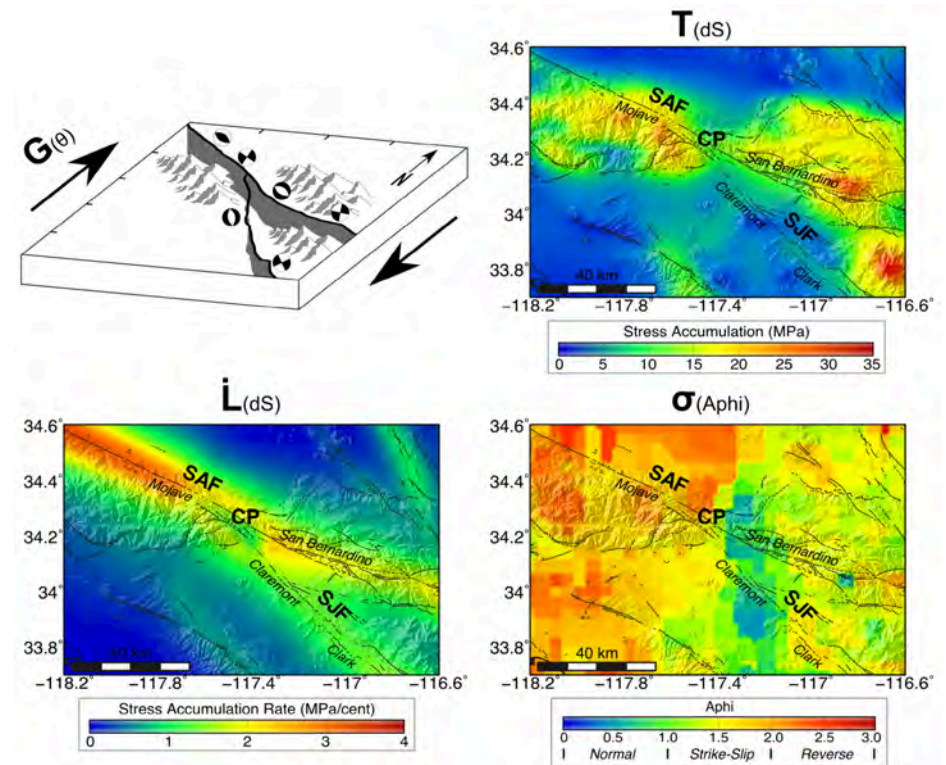


# What's new and current?

- Can we explicitly make sense of the near fault areas using the stressing rate models?
- Cajon Pass is a good place to start
- Two types of free parameters



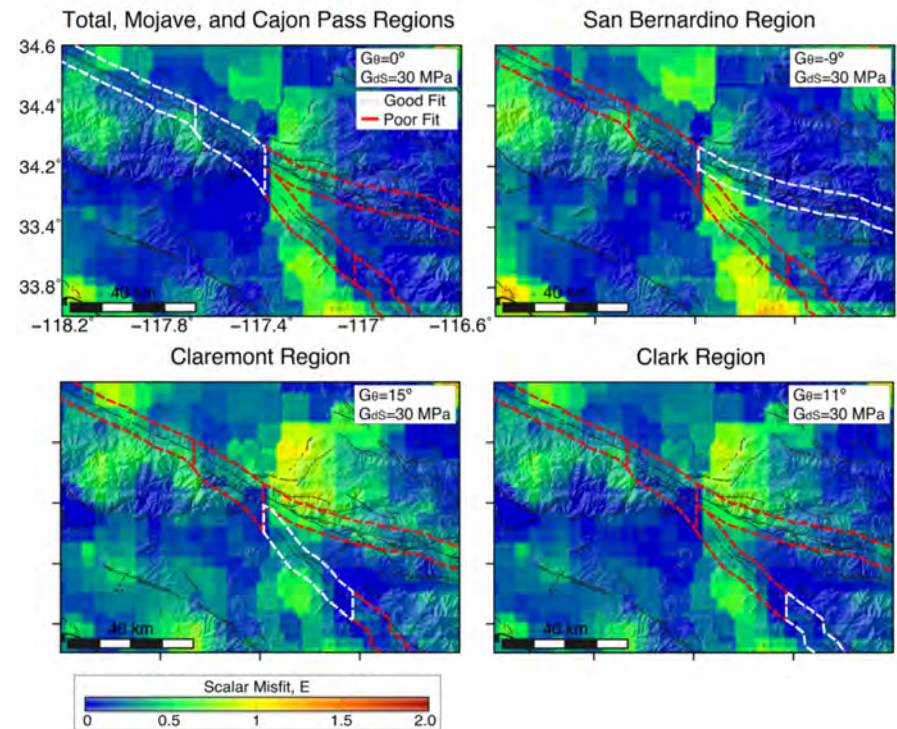
Elliott Helgans, LSU MS student



# What's new and current?

- Can we explicitly make sense of the near fault areas using the stressing rate models?
- Cajon Pass is a good place to start
- Two types of free parameters
- **Bottom line: can make it fit ok if...**
  - Fault loading times are very long (1000s of year)
  - **OR**
  - **G** orientation varies within even this small region
  - ...

## Model misfits along different fault segments



Helgans et al [AGU, 2018]

An aerial photograph of a vast mountain range, likely the Sierra Nevada, showing numerous ridges and valleys. The terrain is rugged and appears to be covered in sparse vegetation or scrubland. The sky is clear and blue. Overlaid on the center of the image is the text "Thanks! Questions?" in a large, white, sans-serif font.

**Thanks! Questions?**