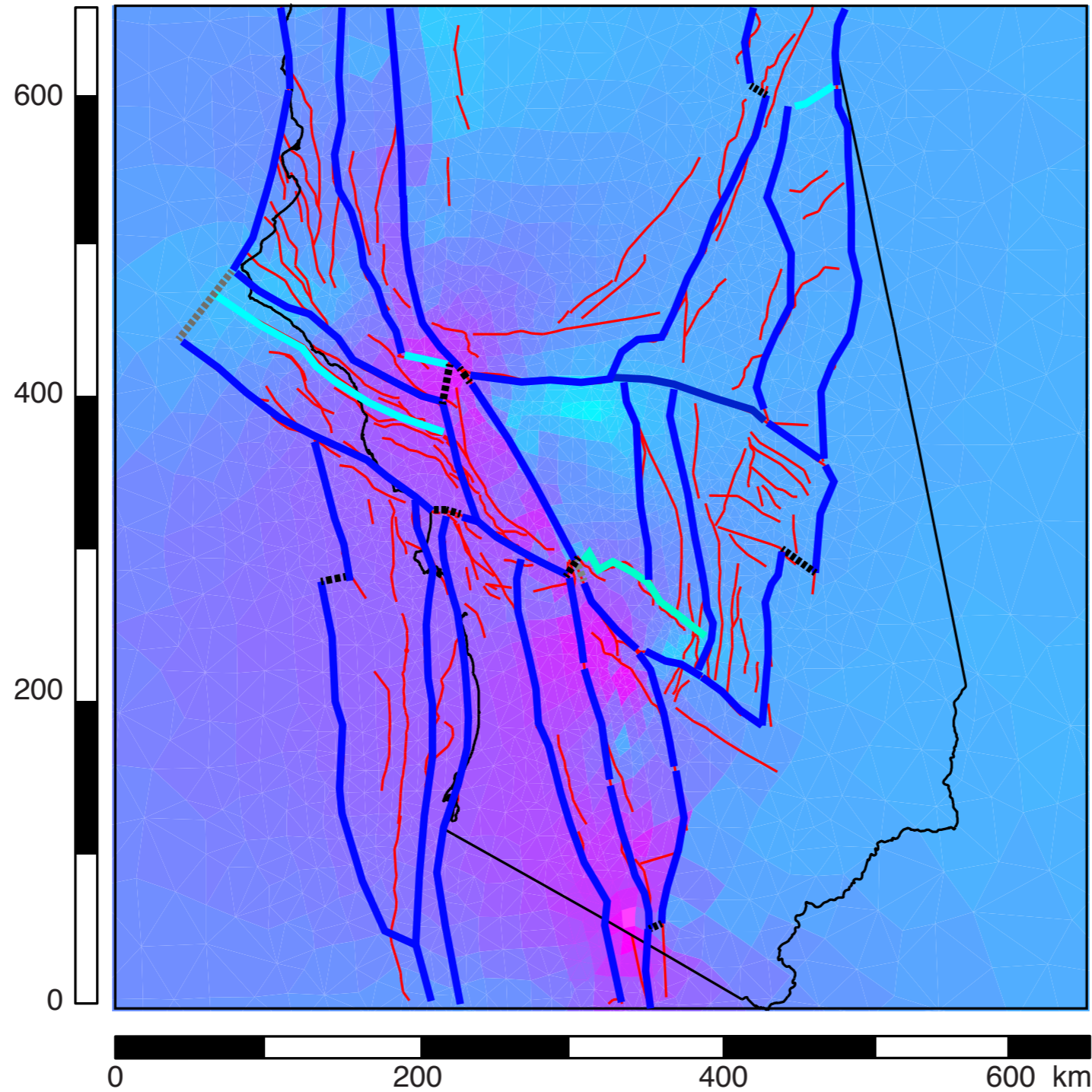


How Sensitive are Inferred Stresses and Stressing Rates to Rheology? Clues from Southern California Deformation Models

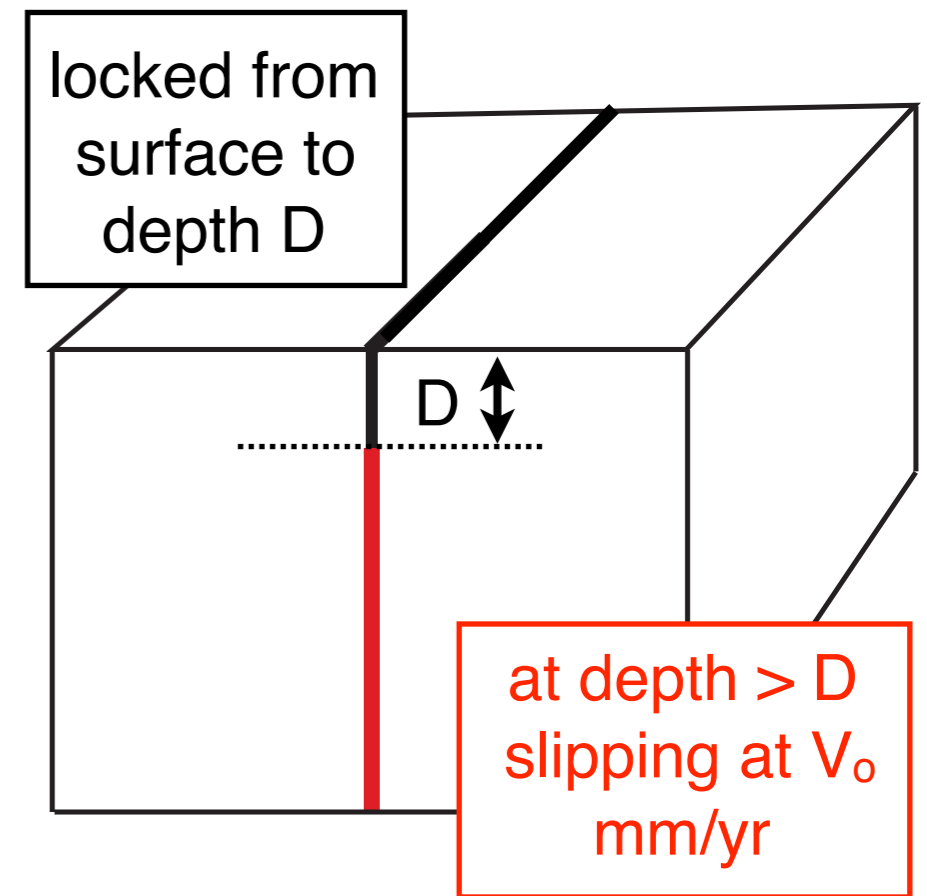
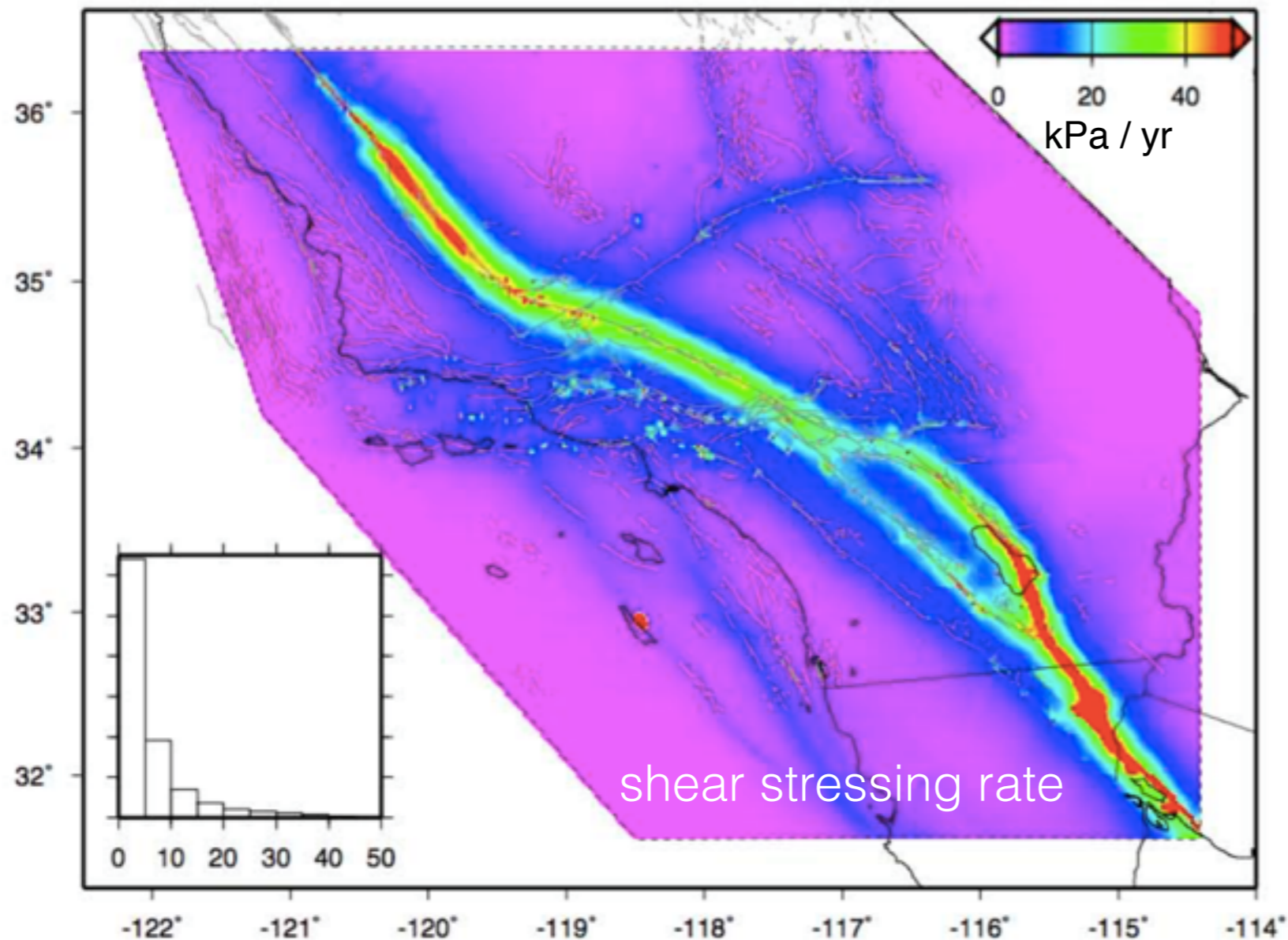
Elizabeth H. Hearn, Capstone Geophysics



SCEC5 theme: Beyond Elasticity

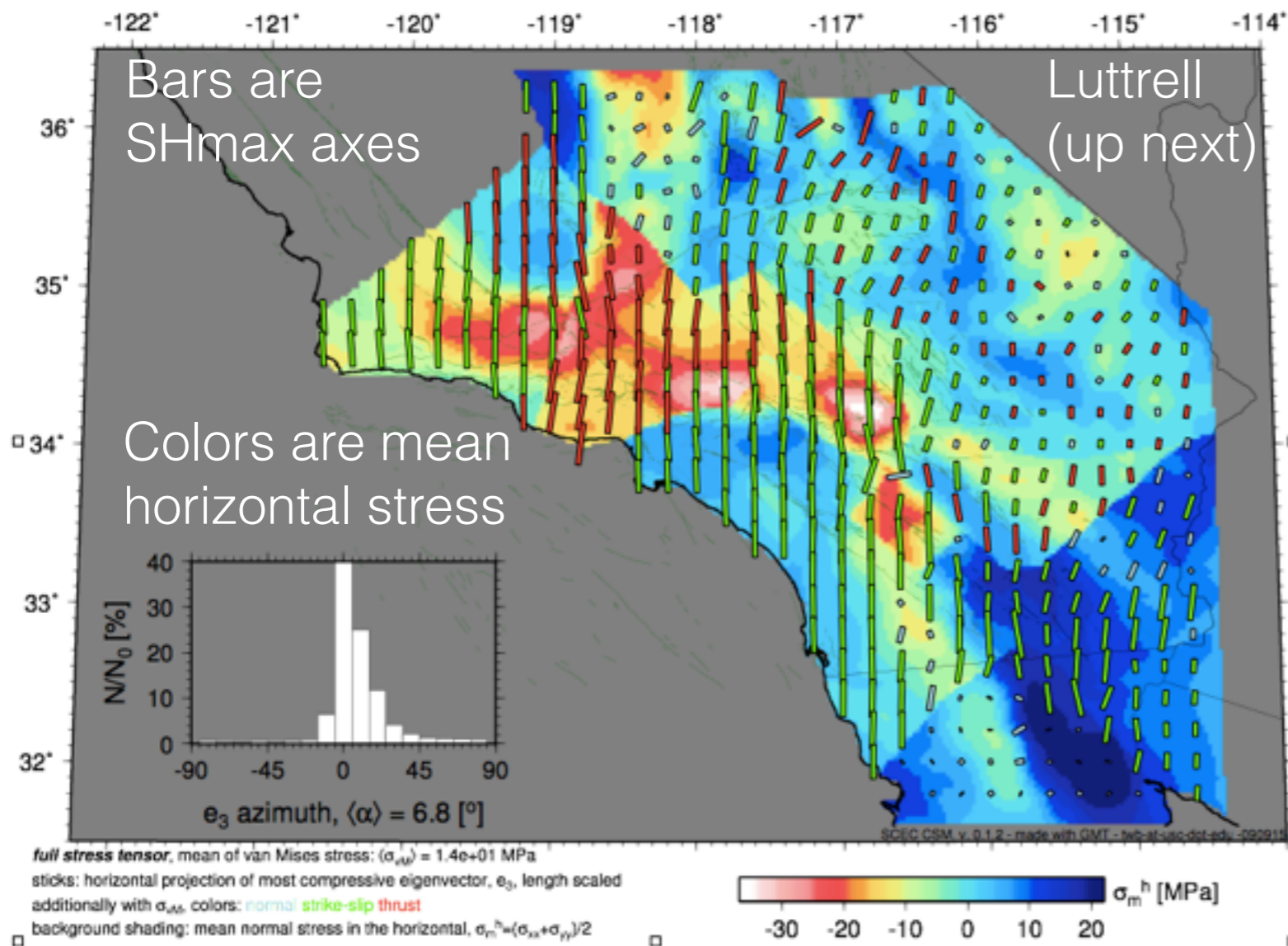
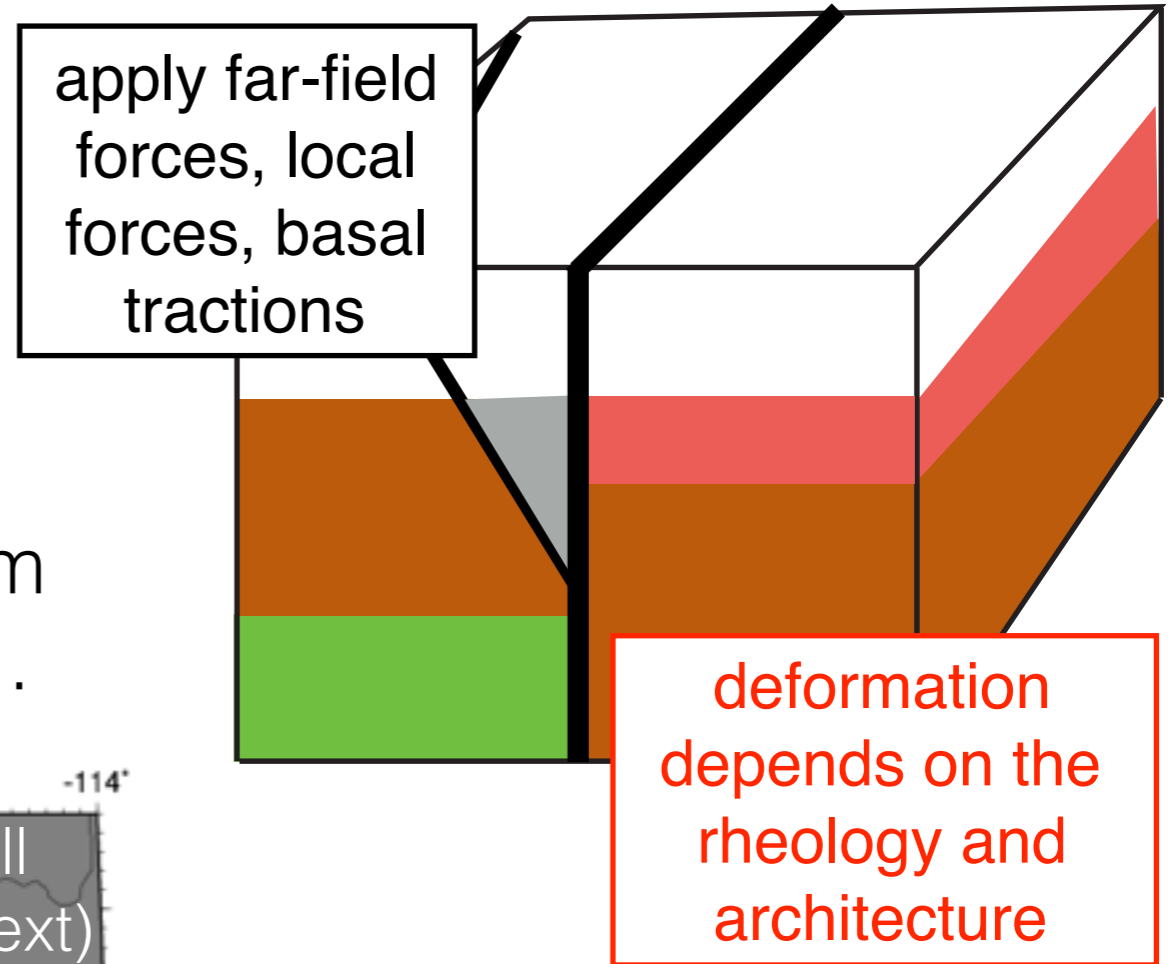
Stressing Rate

- Kinematic models of fault systems
- Slipping faults embedded in elastic volume
- Solve for slipping rates and locking depths



Stress

- Dynamic models of fault systems
- Balance forces driving and resisting deformation
- Tune rheology to match GPS, long-term uplift, geologic slip rates, stress data...



Consistent with larger-scale geodynamics and geologic history

Consistent with *observed reaction to perturbations.*

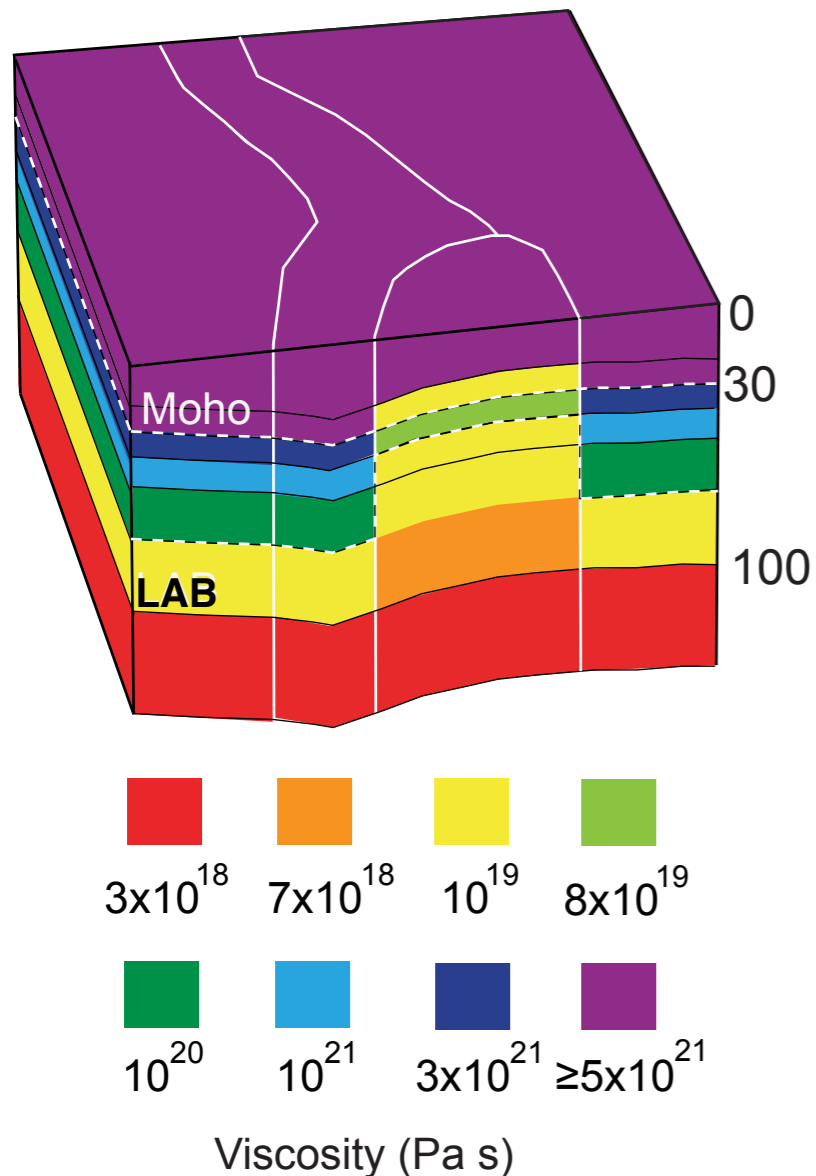
Stress change

- Earthquake cycle (perturbations relative to an average)
- Coseismic, postseismic
- Loading/unloading

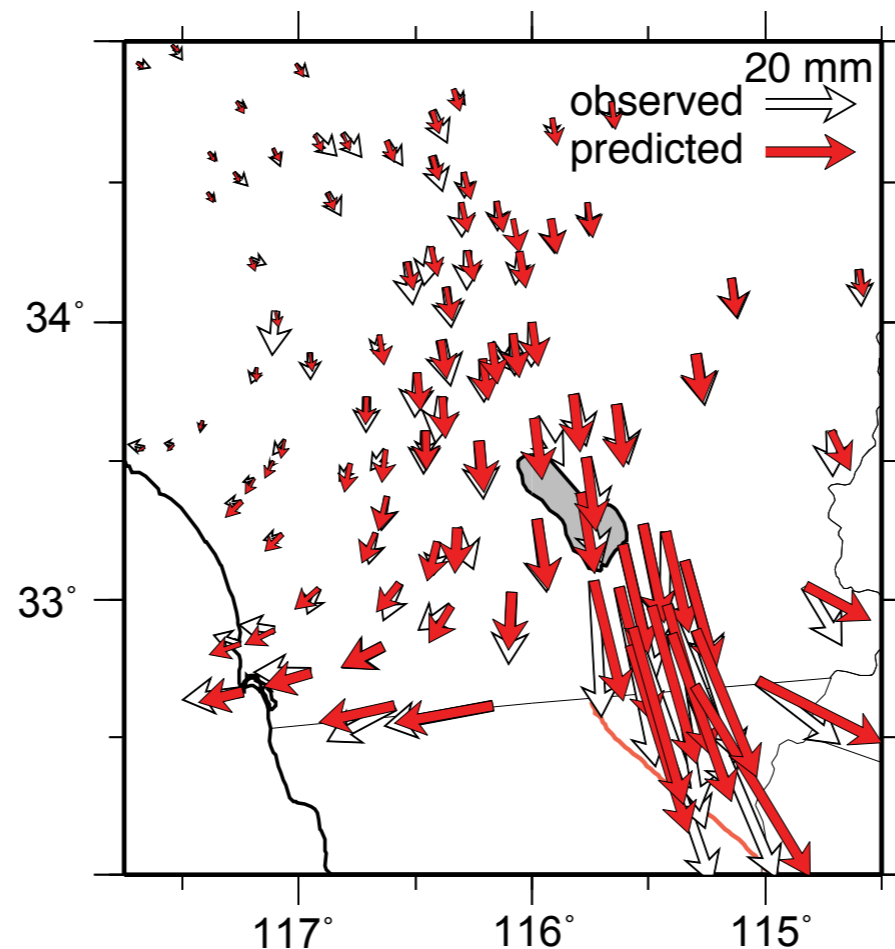
Coseismic and postseismic deformation from the 2010 Mw 7.2 El Mayor-Cucapah Earthquake in Baja California: Lithospheric structure and deformation in the Salton Trough

Huang, Dickinson, Freed, Fielding, Bürgmann, and Gonzalez-Ortega

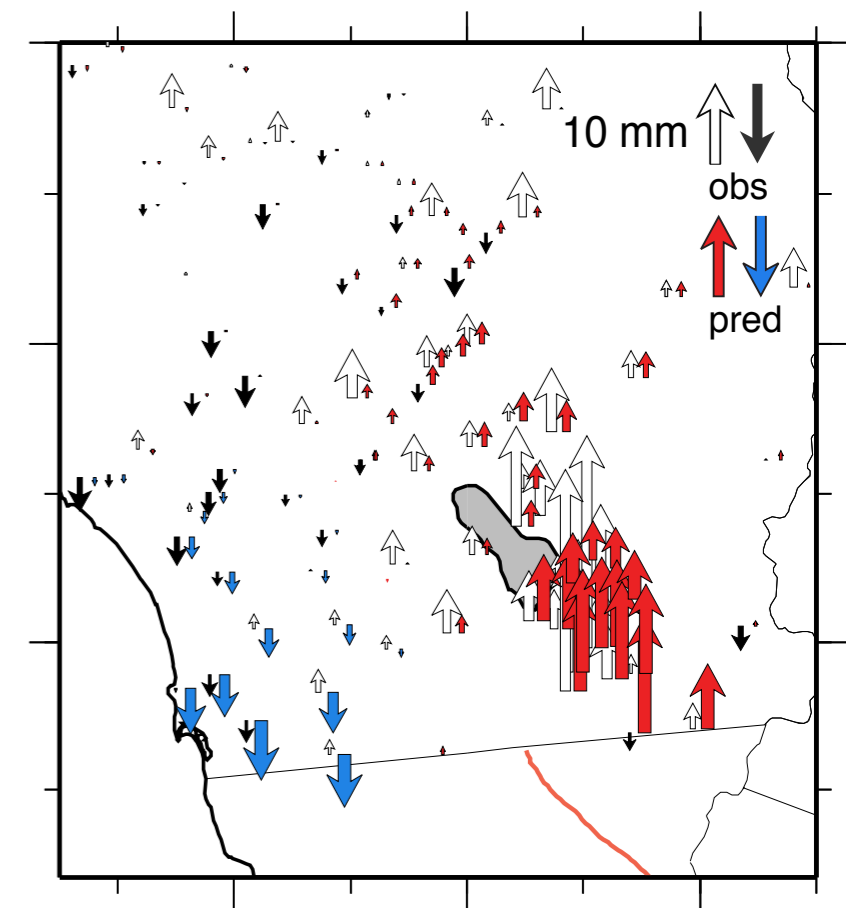
Heterogeneous 3D viscosity structure



Lateral GPS



Vertical GPS

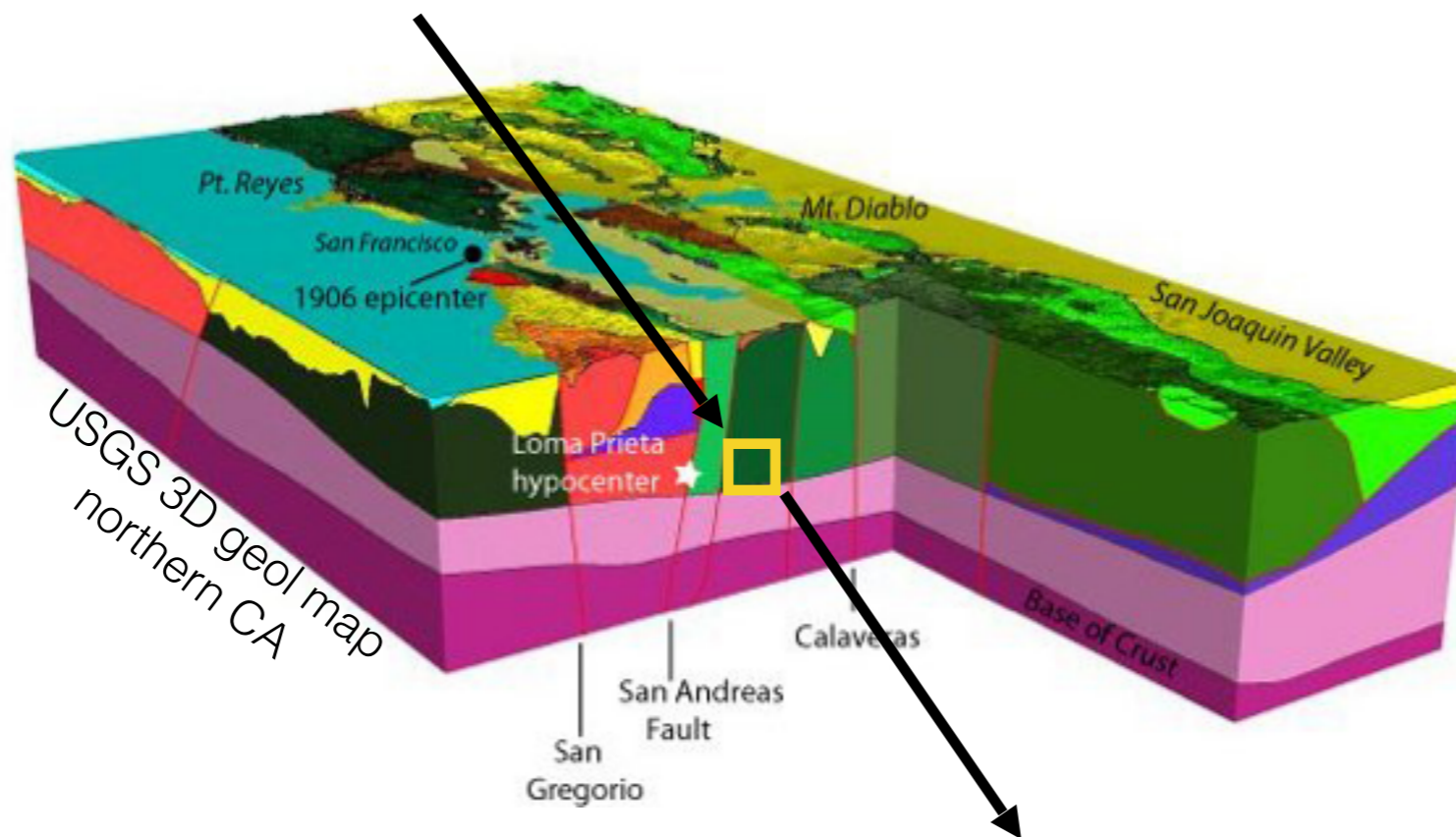


Non-unique estimates of southern CA viscosity from models

<i>POSTSEISMIC</i>	<i>lower lithosphere effective viscosity Pa s, Maxwell if not otherwise indicated</i>	<i>references</i>
<i>El Mayor-Cucapah M 7.2</i>	<i>heterogeneous, Maxwell or power-law. 1-3 e18 Pa s mantle and (local) lower crust</i>	<i>Rollins et al, 2015; Gonzalez-Ortega et al., 2014; Pollitz et al., 2012</i>
<i>Central Nevada Quakes M 6.9 - 7.4</i>	<i>1e18 to 1e19 Pa s</i>	<i>Hammond et al., 2009</i>
<i>Landers and Hector Mine M 7.4 and 7.2</i>	<i>1e18 to 1e20 Pa s; Maxwell, power-law and/or transient flow laws with assumed T;</i>	<i>Pollitz (2015 and earlier refs), Fialko, Freed...</i>
<i>“LONG-TERM” DEFORMATION</i>		
<i>So Cal</i>	<i>From Kirby (1983) and computed T field, then perturbed to fit data</i>	<i>Bird and Kong, 1994</i>
<i>So Cal</i>	<i>4e19 to 1e21</i>	<i>Li et al., 2009</i>
<i>San Andreas</i>	<i>2e18 to 5e19 (below 70 km plate)</i>	<i>Smith and Sandwell, 2004</i>
<i>EARTHQUAKE CYCLES</i>		
<i>SSAF</i>	<i>2e19 - 1e21 layered and transient rheol. or 5e20 uniform Maxwell</i>	<i>Hearn et al. 2013</i>
<i>SoCal-wide</i>	<i>6e18 mantle, 2e20 lower crust</i>	<i>Chuang and Johnson, 2011</i>
<i>SSAF</i>	<i>0.1 - 3e20 mantle asthenosphere</i>	<i>Johnson and Segall, 2004</i>
<i>Carrizo SAF</i>	<i>2 - 5 e19 lower crust/upper mantle</i>	<i>Schmalzle et al., 2006</i>

Rheological constraints from geology, lab experiments and temperatures: SCEC **CRM**

lat, lon, elevation
shear zone or host rock?



temperature (from **CTM**)
admissible rock type(s)
admissible flow law(s)

→ rheology

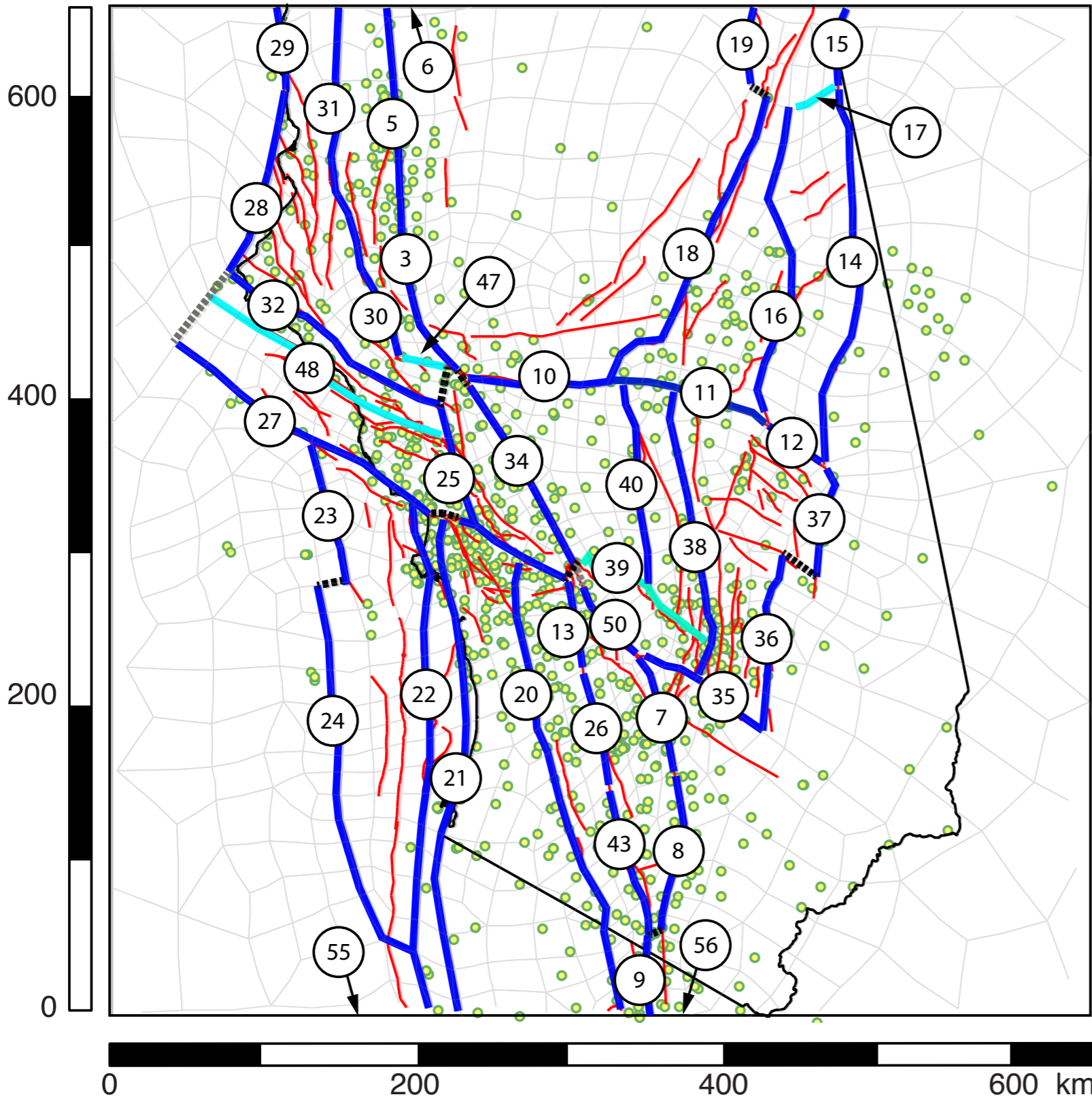
Effects of

1. Viscoelastic lower lithosphere
2. Lithosphere-scale shear zones
3. Plastic upper crust

on modeled fault system behavior and stresses

Start with a kinematic model of southern California

A southern California kinematic model (elastic FEM)

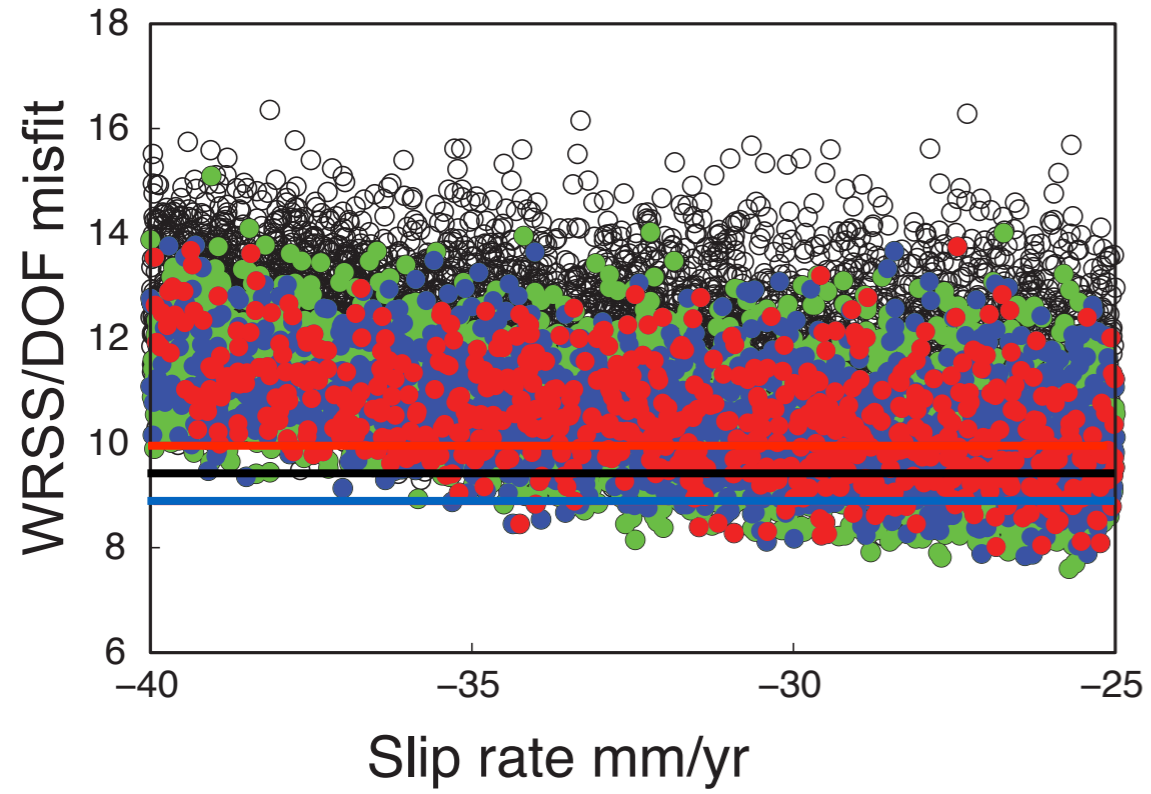


Fault geometry and slip rate ranges from UCERF3

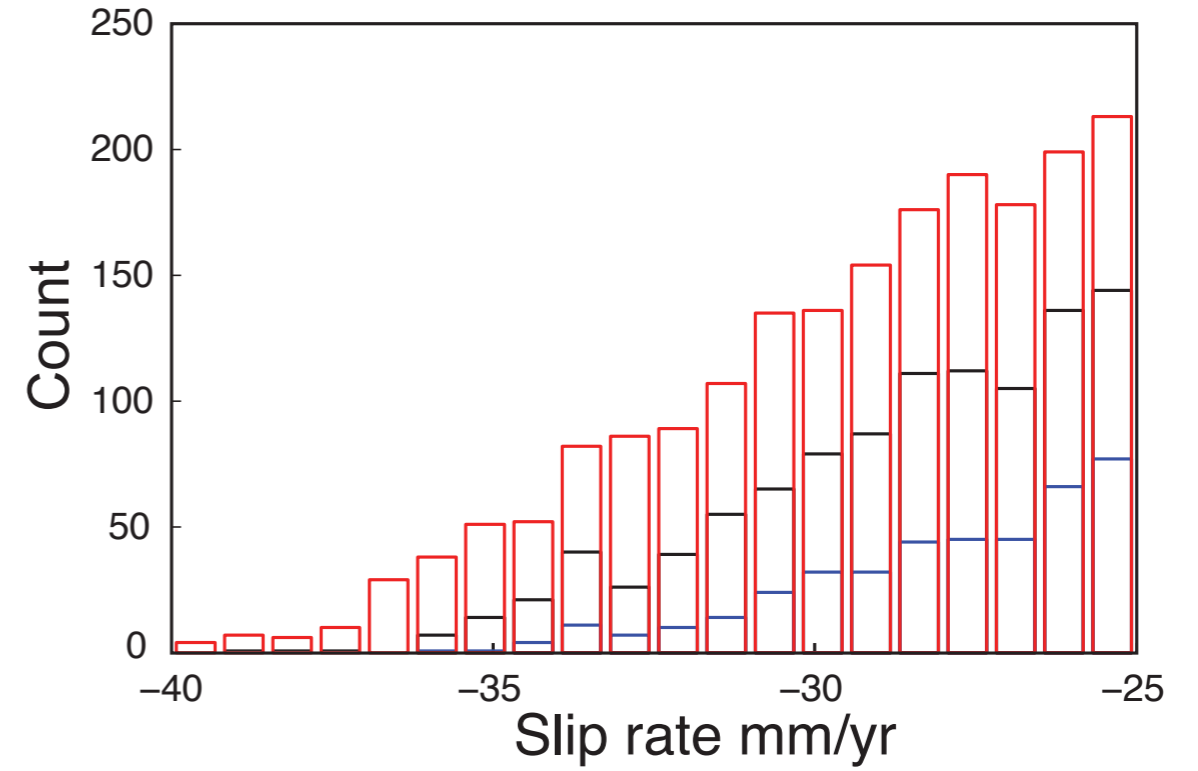
Locking depths:
(1) 12 km
(2) from UCERF3
(3) no locking

Draw slip rates from UCERF3 ranges, run 10,000 cases. Find models that fit SCEC CMM4 GPS velocity field or minimize strain energy

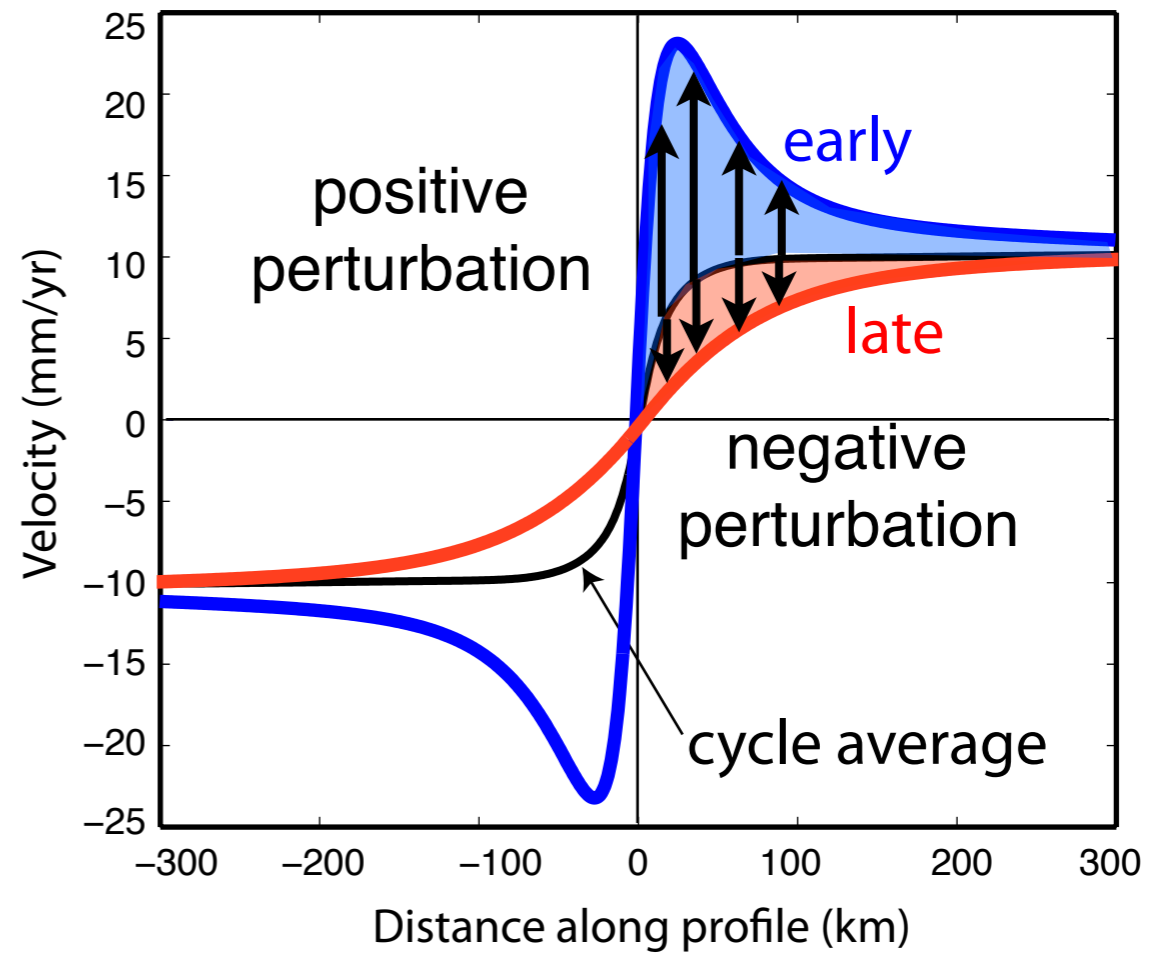
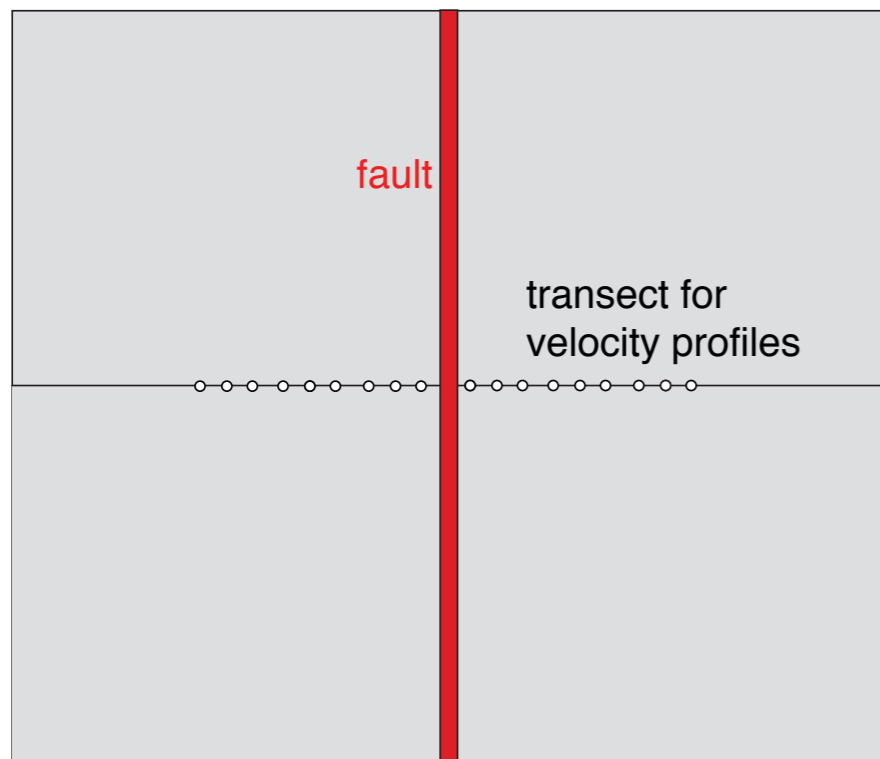
GPS misfit minimization



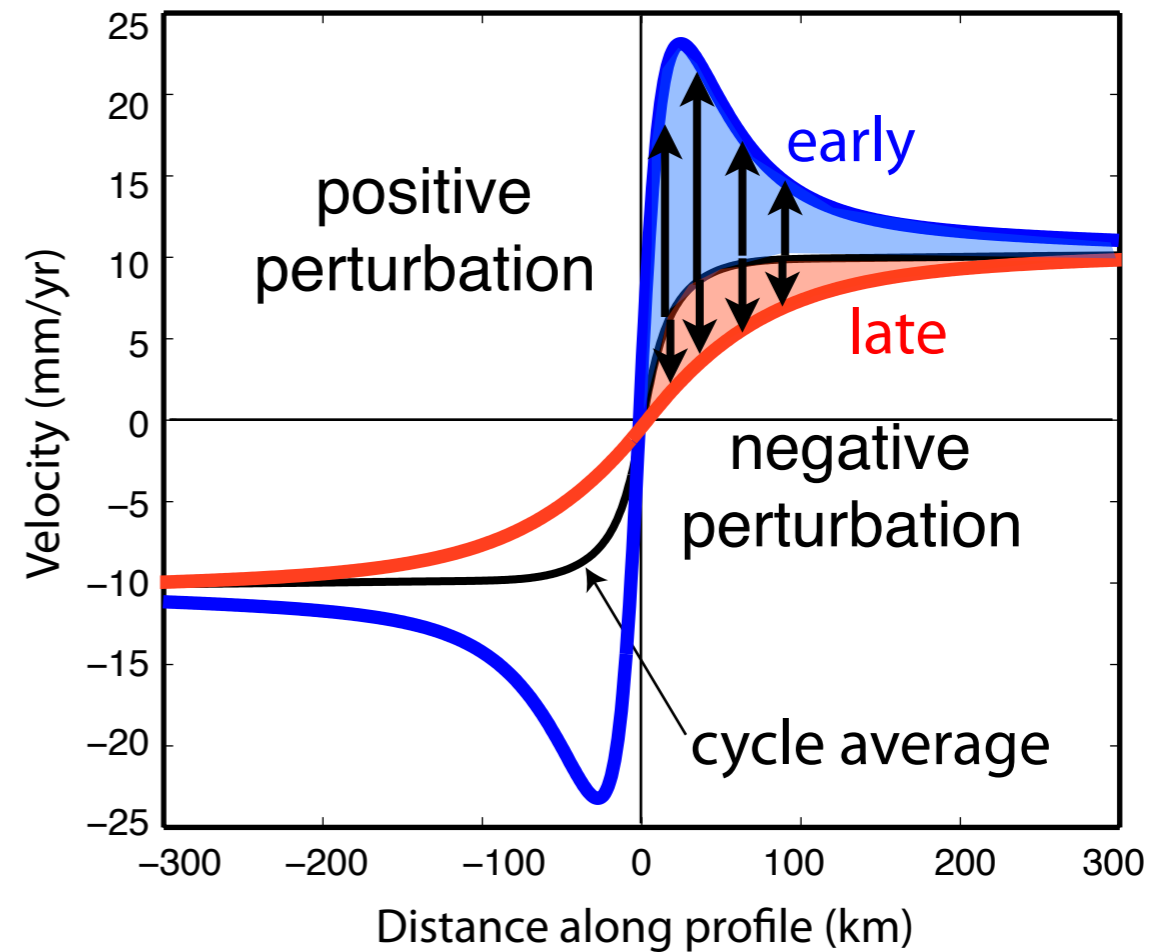
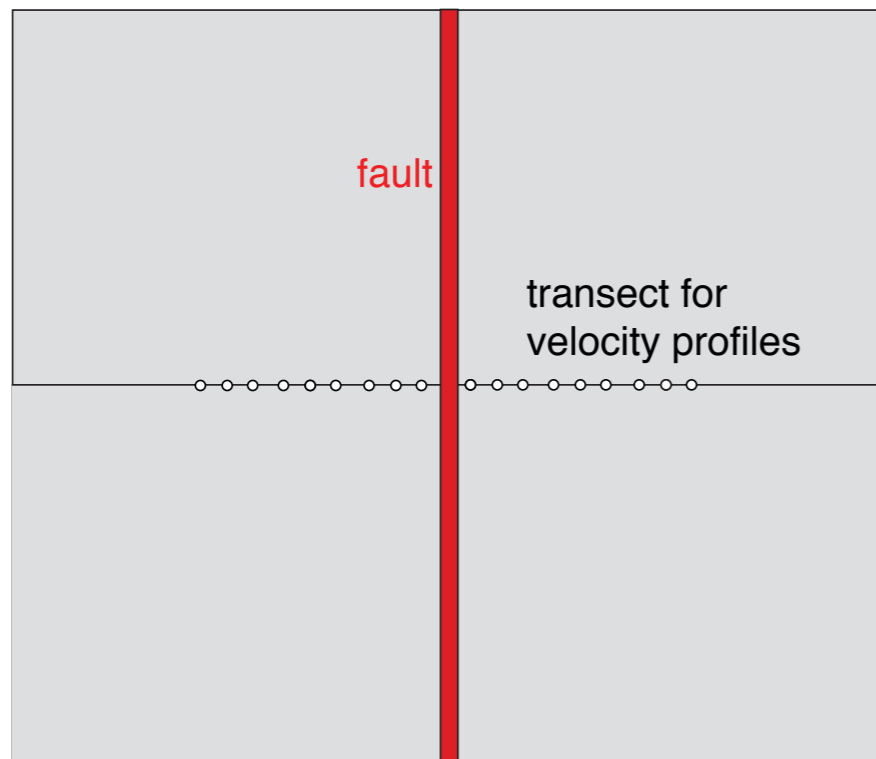
Low WRSS Models



1. Effect of viscoelastic lower lithosphere



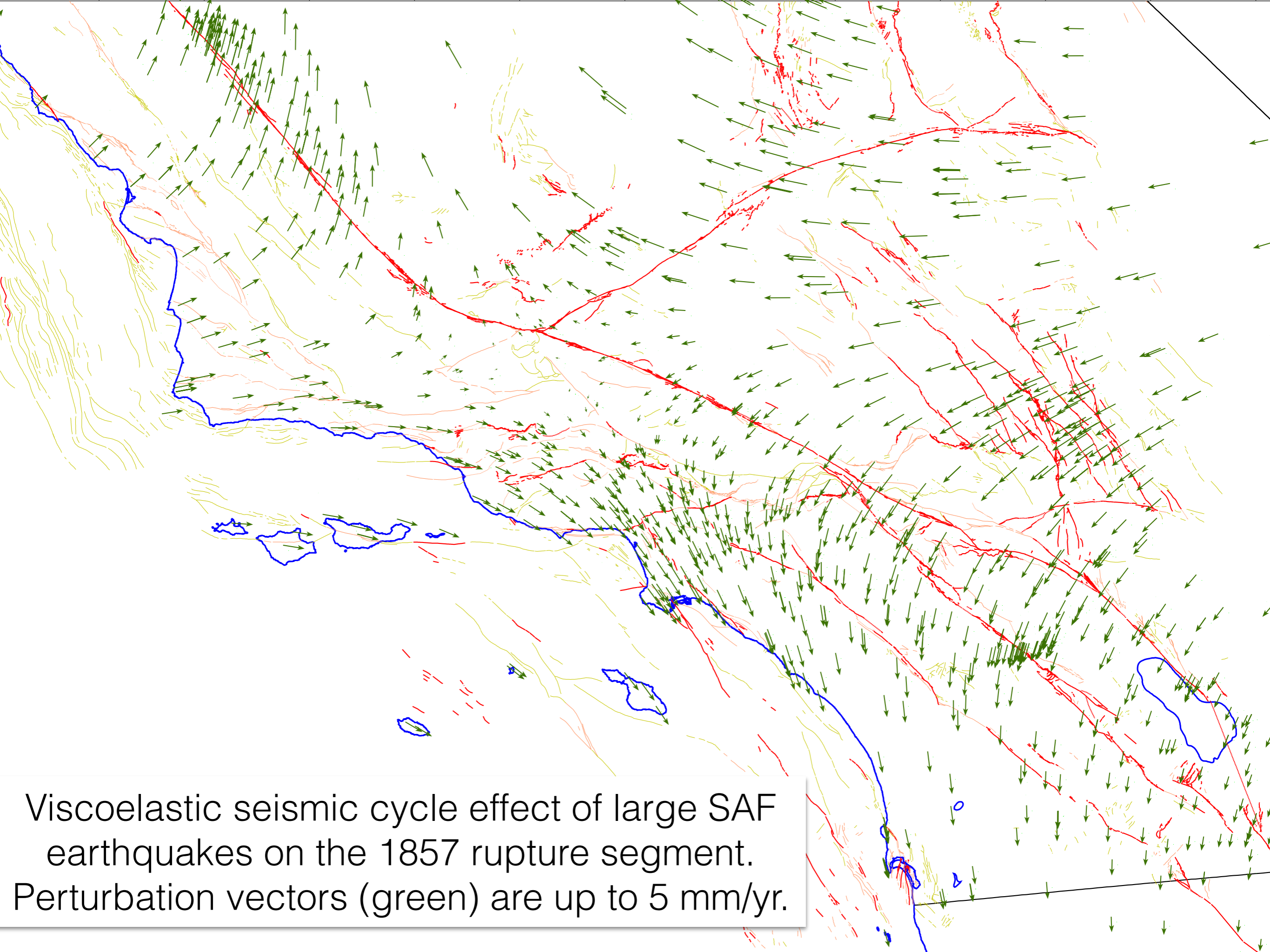
1. Effect of viscoelastic lower lithosphere



Model M1

	elastic upper crust	strike-slip fault
25 km		
30 km	viscoelastic lower crust	$\eta_1 = 3 \times 10^{19} \text{ Pa s}$
	viscoelastic uppermost mantle	$\eta_1 = 1 \times 10^{21} \text{ Pa s}$
50 km		
	viscoelastic mantle	$\eta_1 = 3 \times 10^{18} \text{ Pa s}$
	$\lambda = \mu = 30 \text{ GPa}$	 μ η_1 Maxwell element

- Calculate perturbation
- Correct the velocity field to get interseismic average velocities
- Infer slip rates

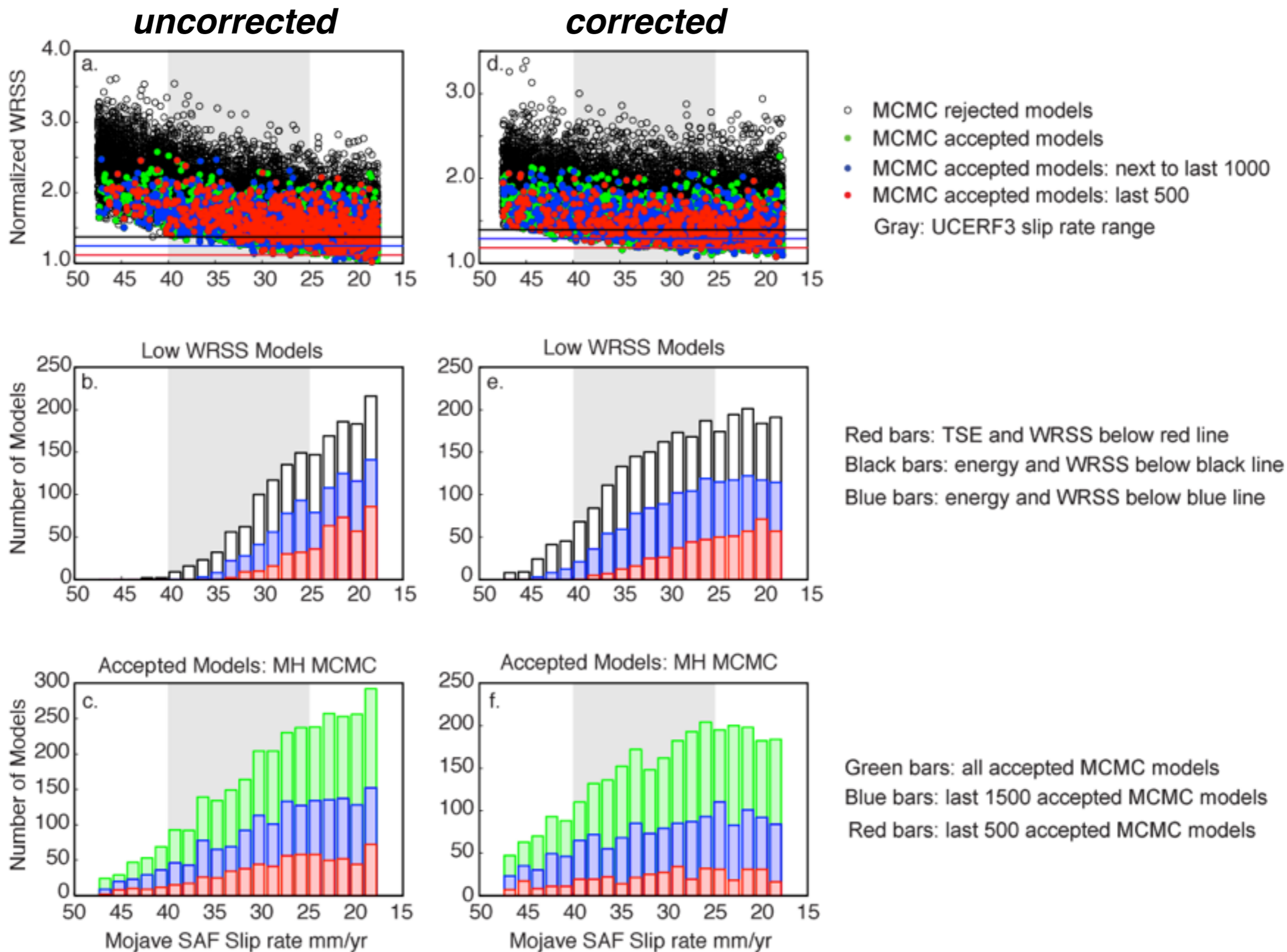


Viscoelastic seismic cycle effect of large SAF earthquakes on the 1857 rupture segment. Perturbation vectors (green) are up to 5 mm/yr.

Correcting for “ghost transient” affects inferred

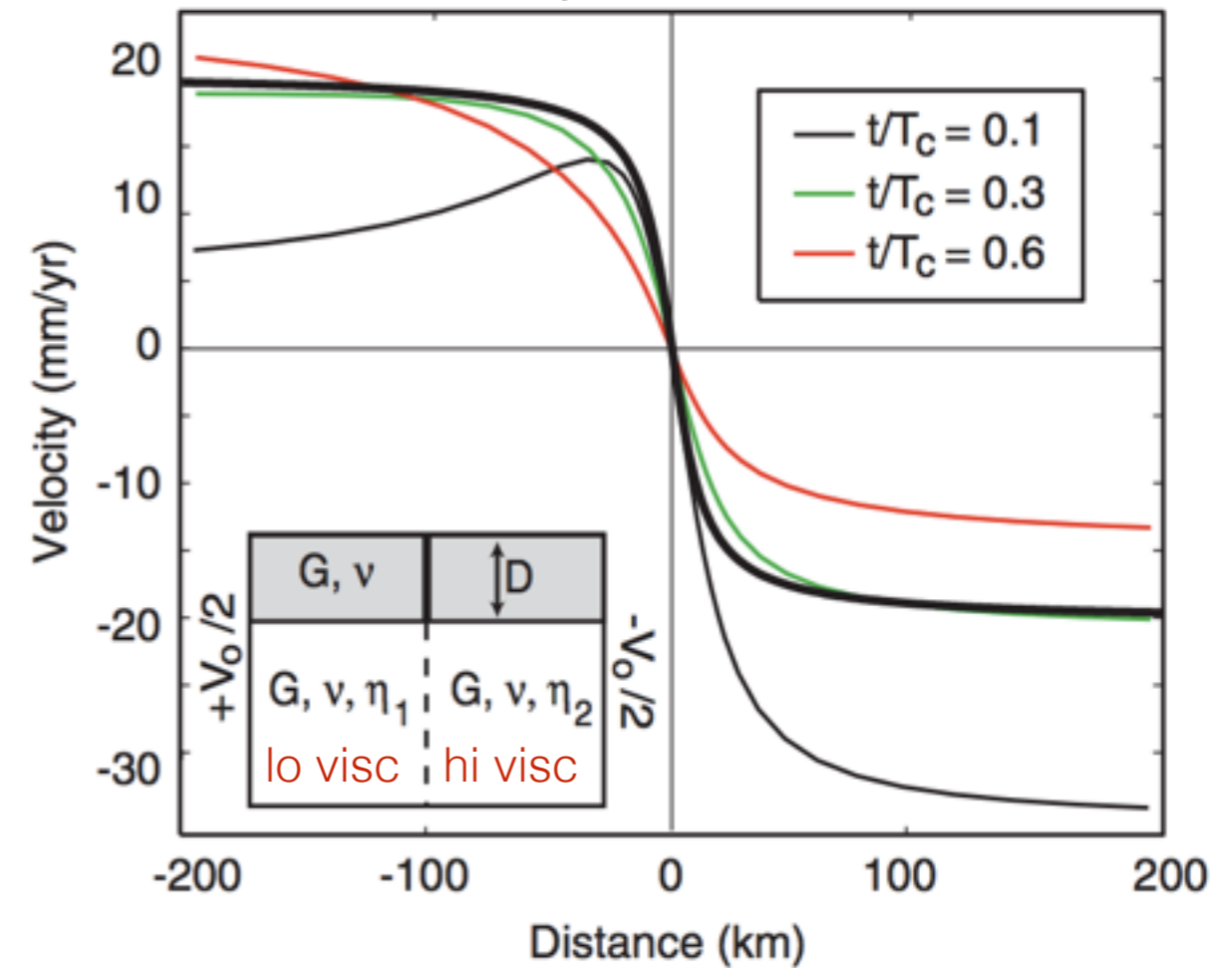
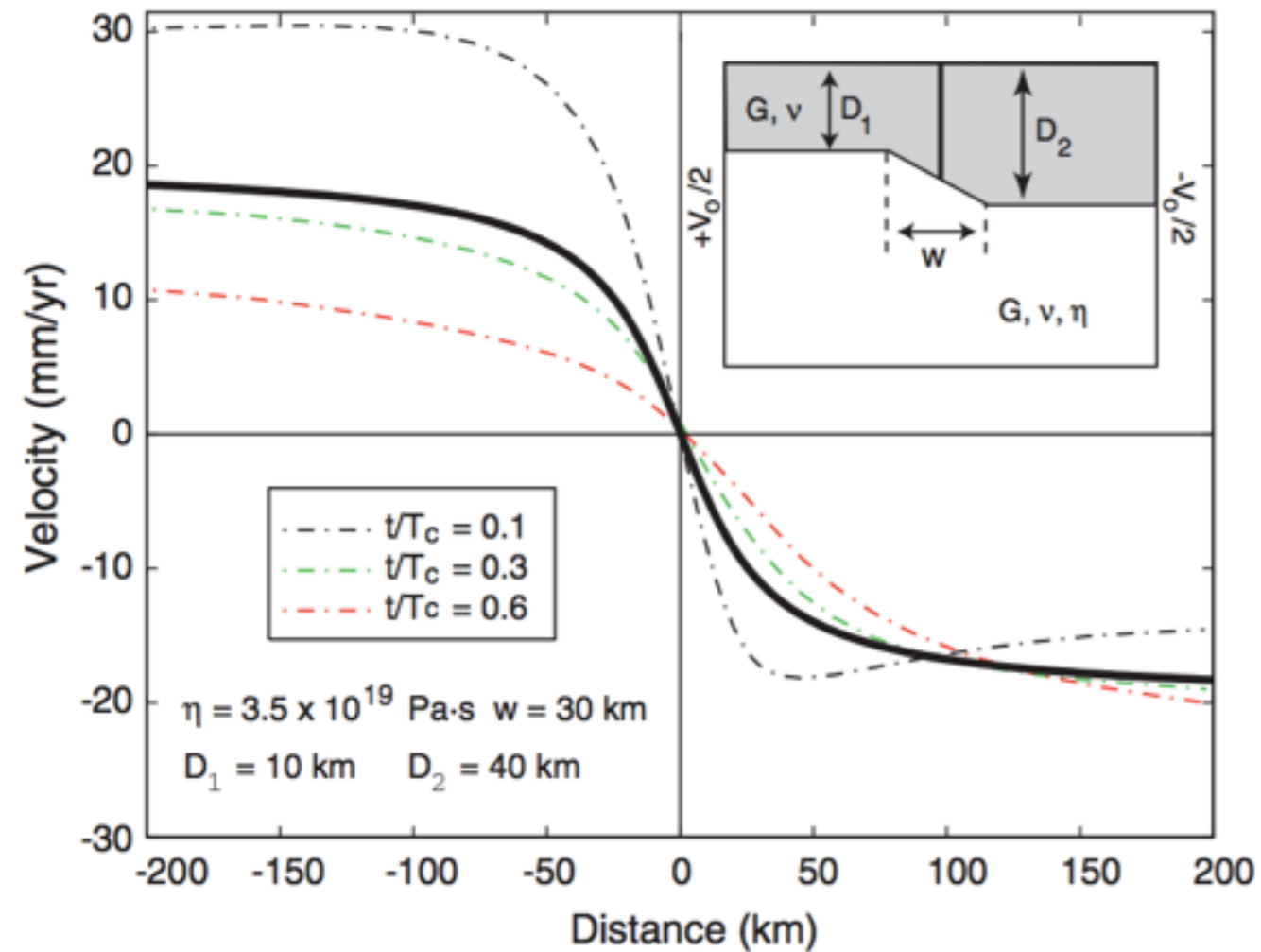
34 SAF Mojave slip rate.

Not a new conclusion: see also Johnson et al., 2007, Chuang and Johnson, 2011, Hearn et al., 2013



Material heterogeneity affects interseismic deformation and surface velocities

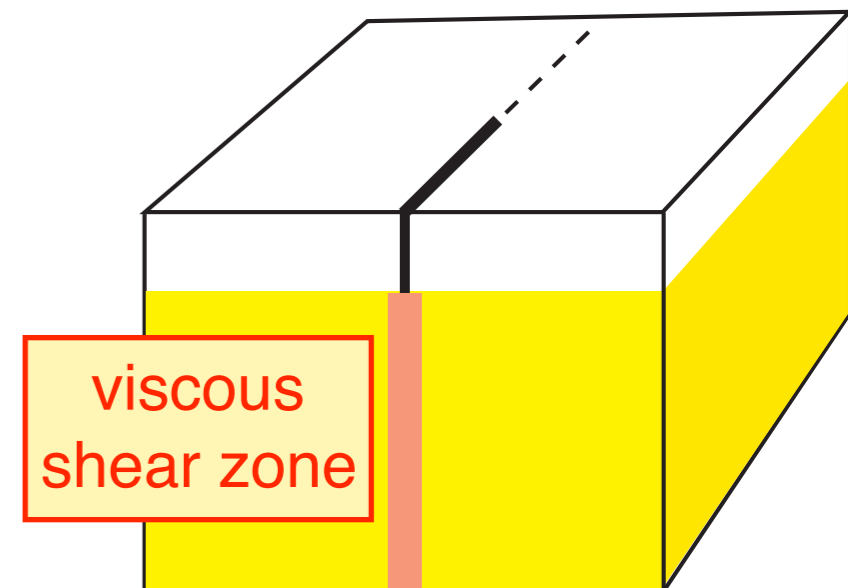
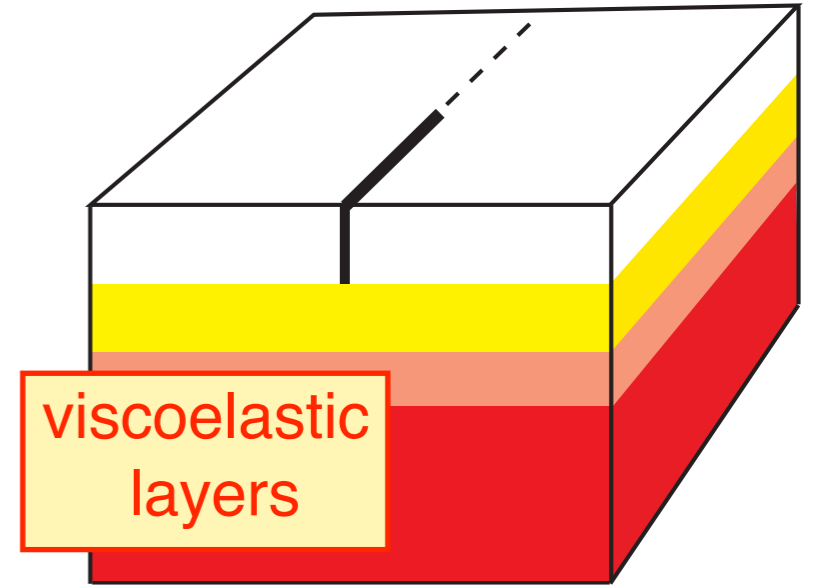
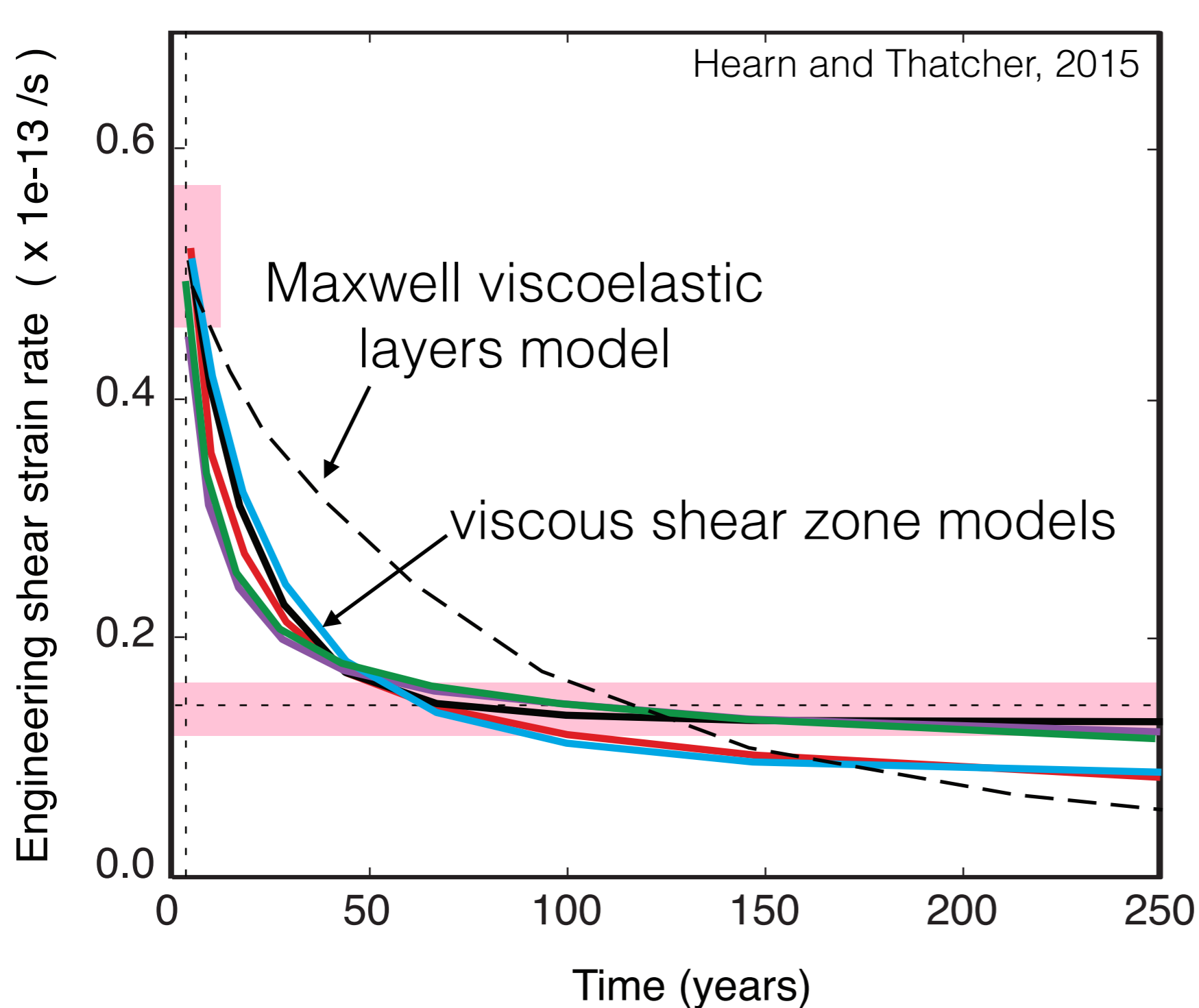
Vaghri and Hearn, 2012



Example: asymmetry

2. Effects of lithosphere-scale shear zones

Viscous shear zones also produce interseismic stress and velocity perturbations

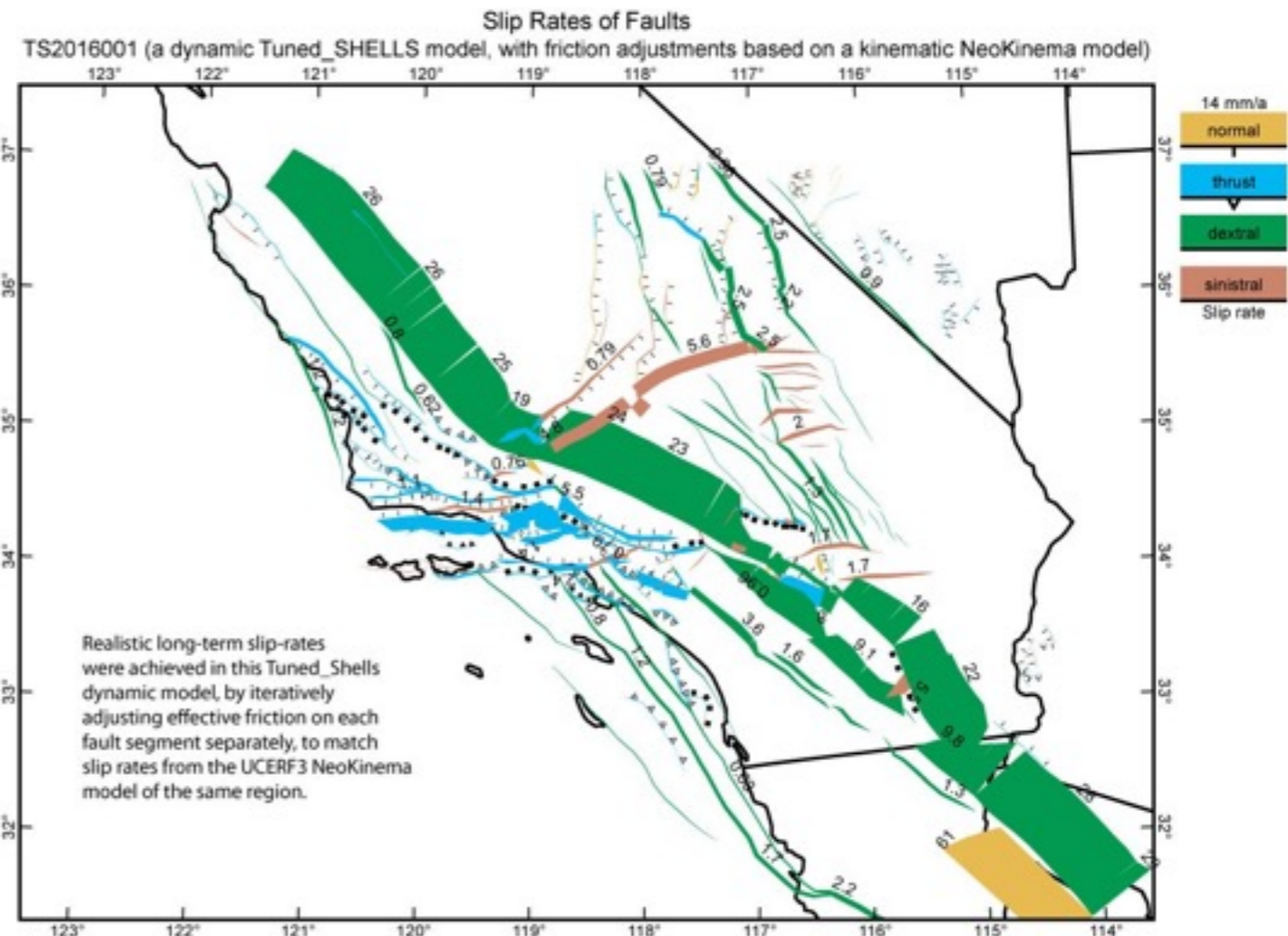
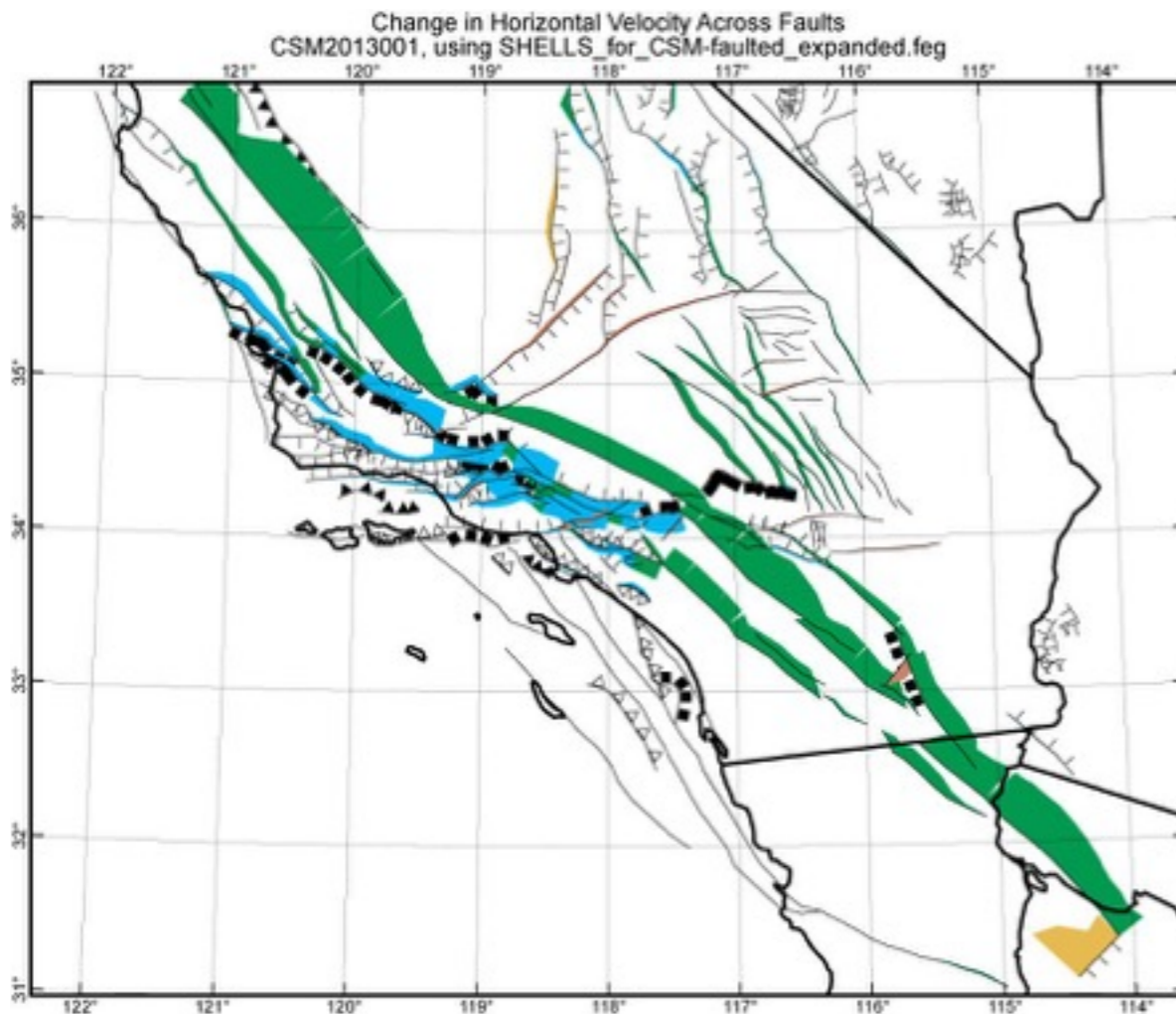


A new way to estimate shear tractions on active faults in southern California

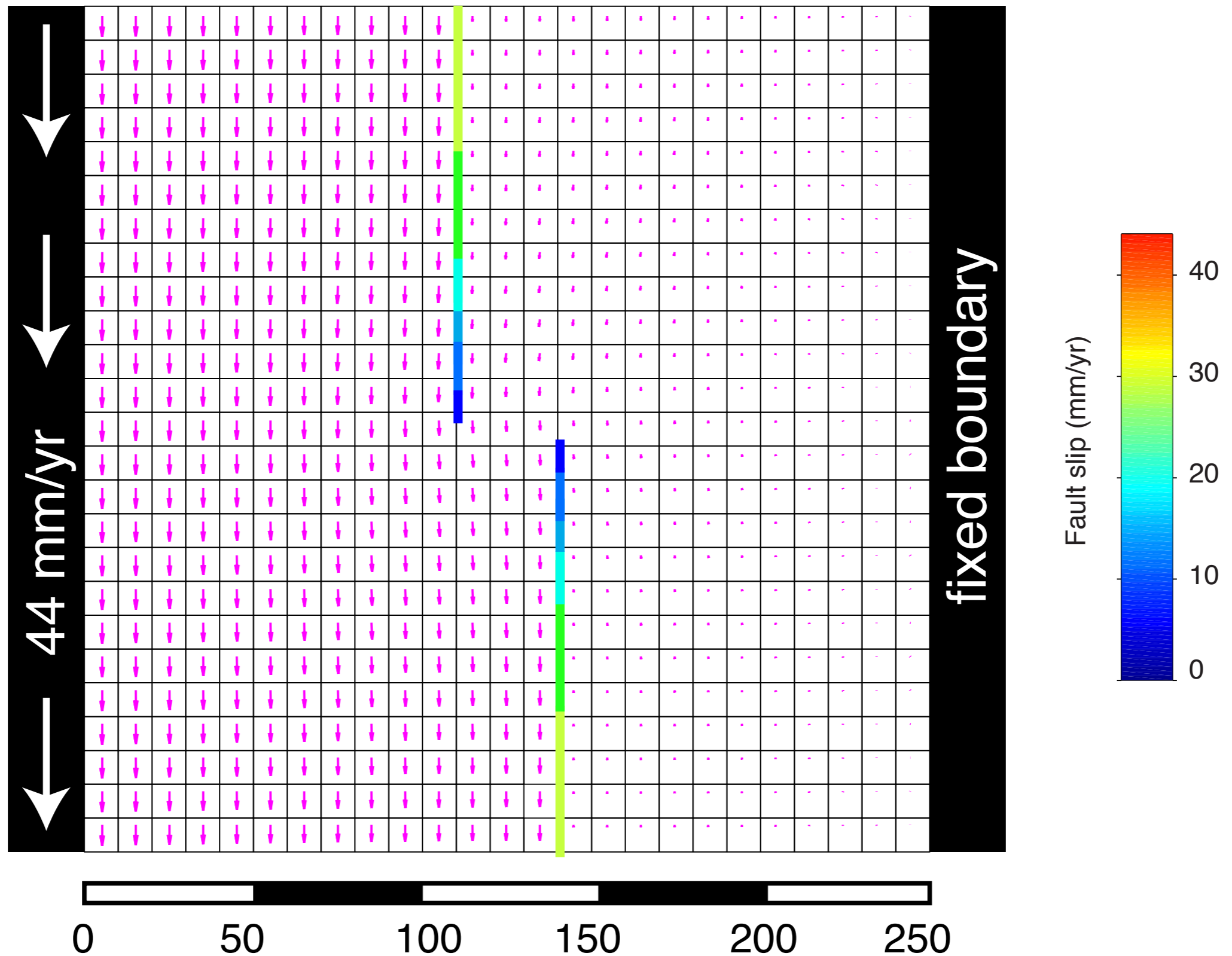
Peter Bird

Model with same effective friction on all faults

Model with effective friction adjusted for each fault segment to match target slip rates

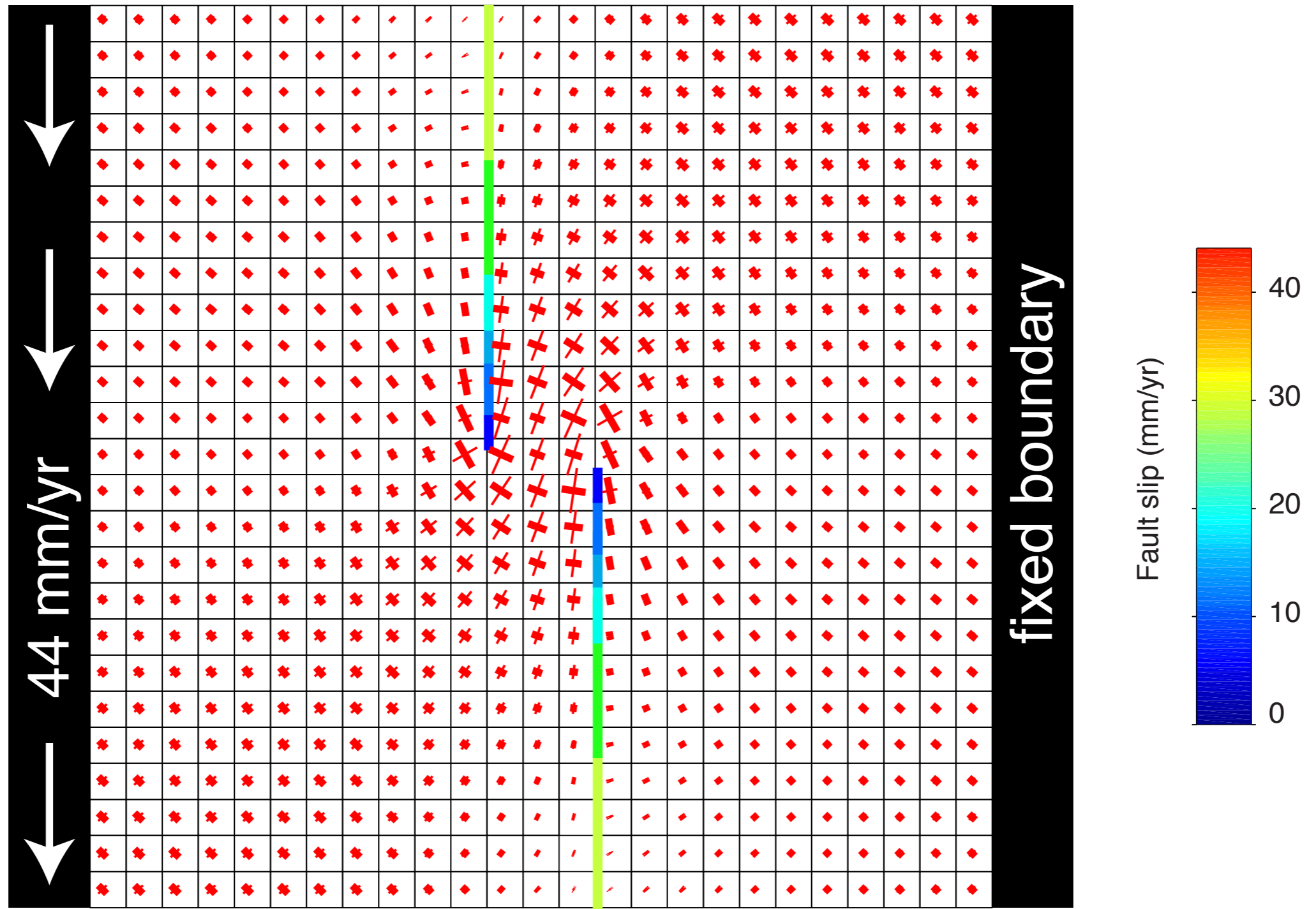


3. Effect of plastic upper crust



Elastic model with traction-free faults and a stepover

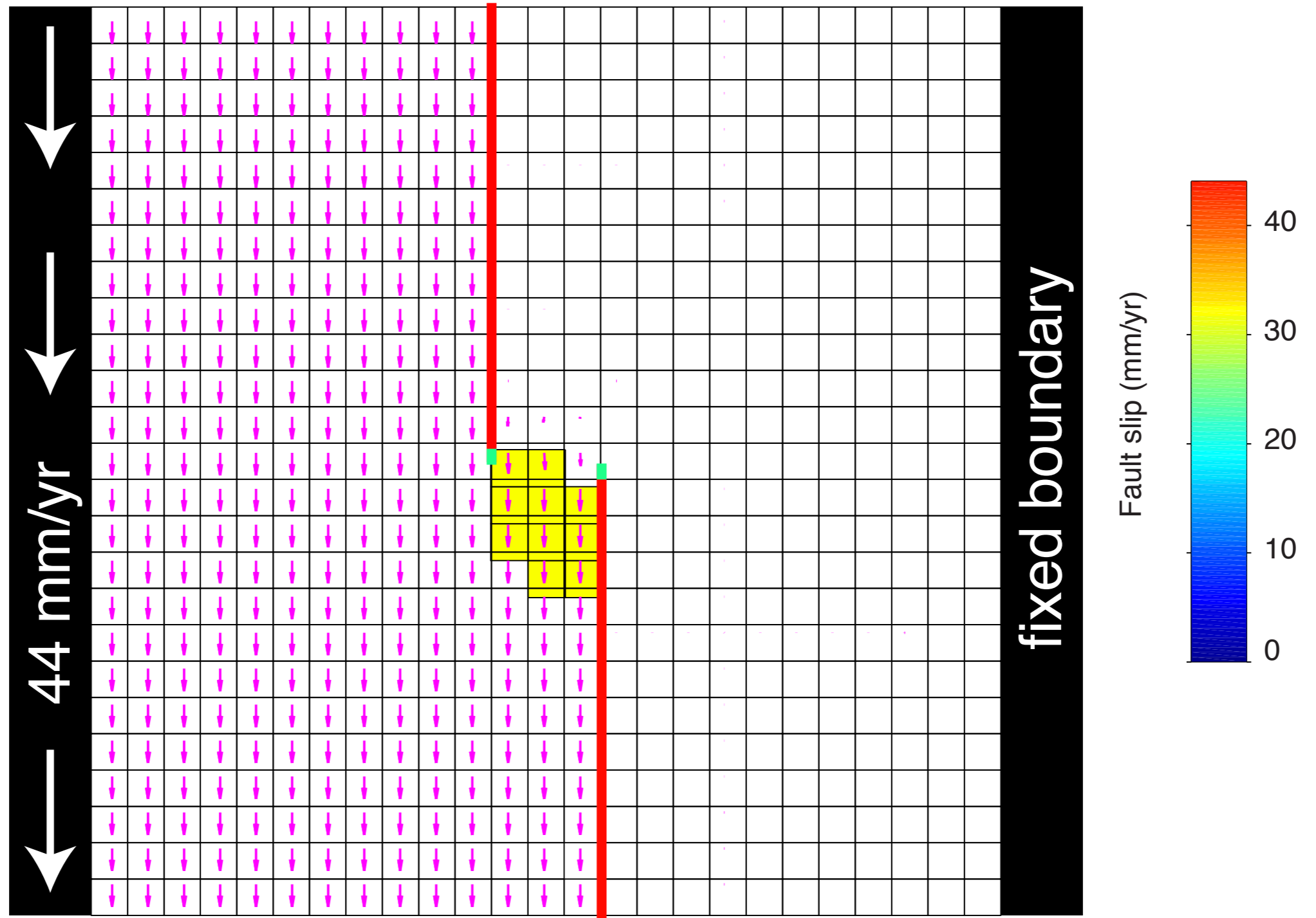
$$\sigma_3 + \sigma_1 = 264 \text{ MPa and counting}$$



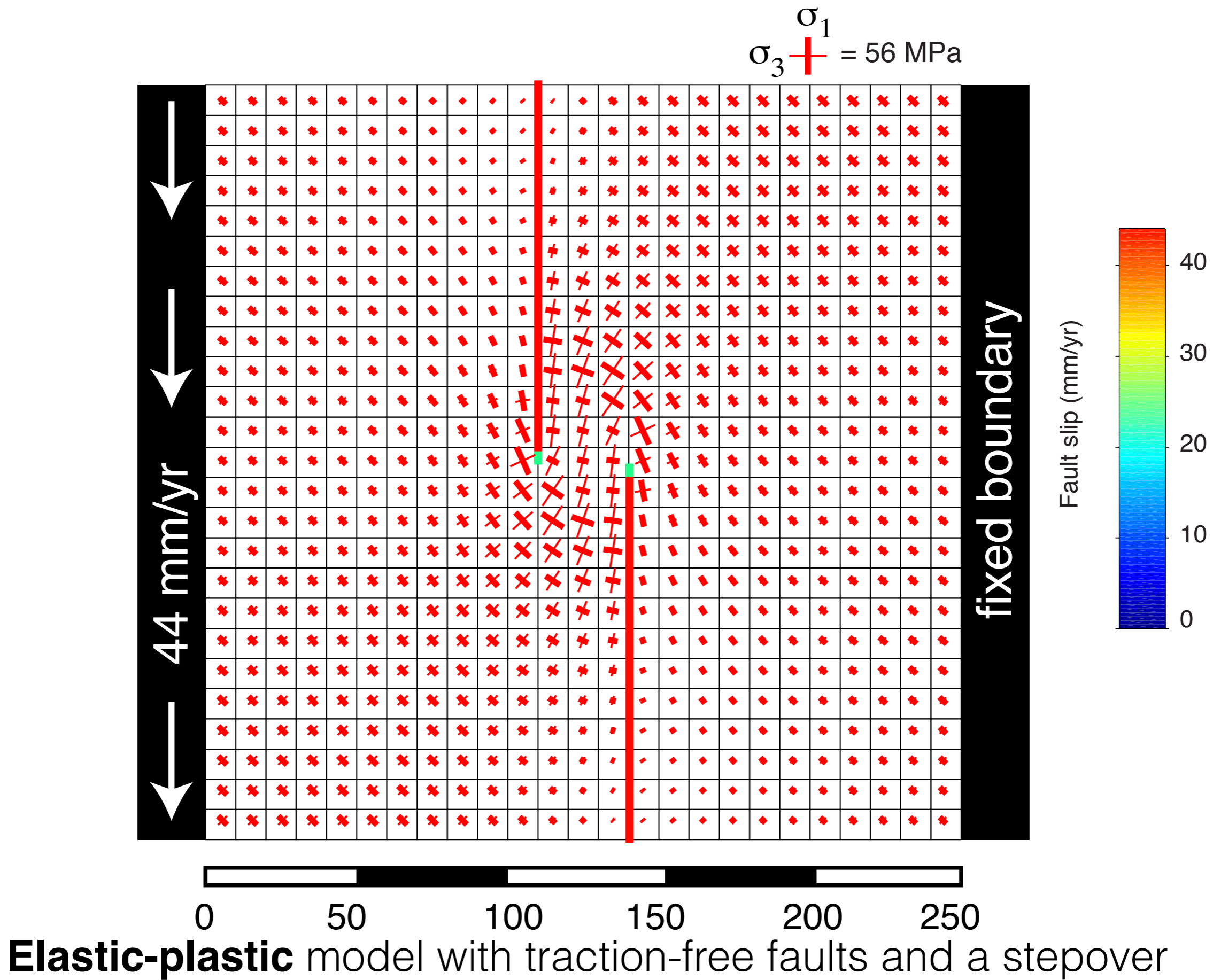
0 50 100 150 200 250

Elastic model with traction-free faults and a stepover

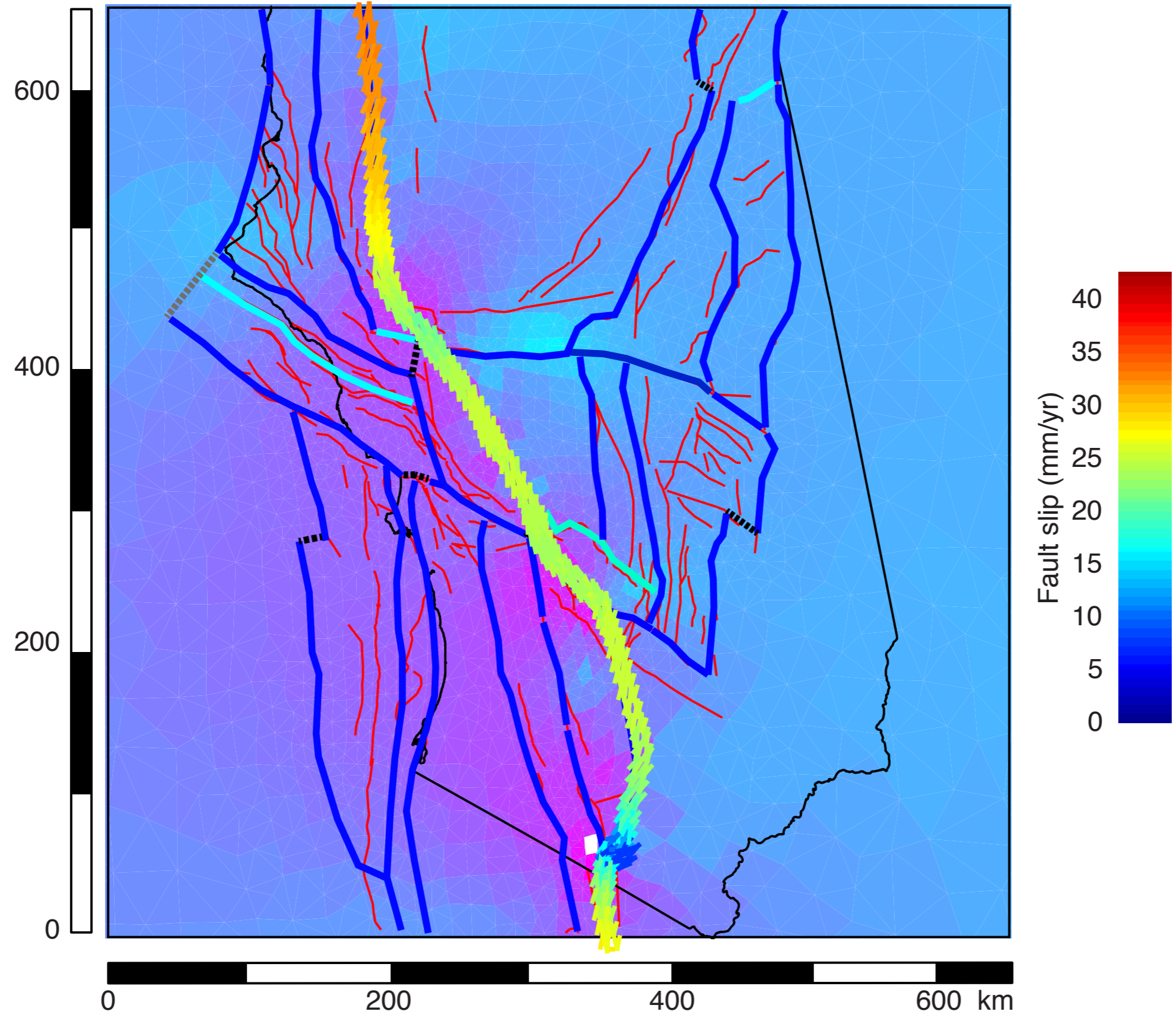
3. Effect of plastic upper crust



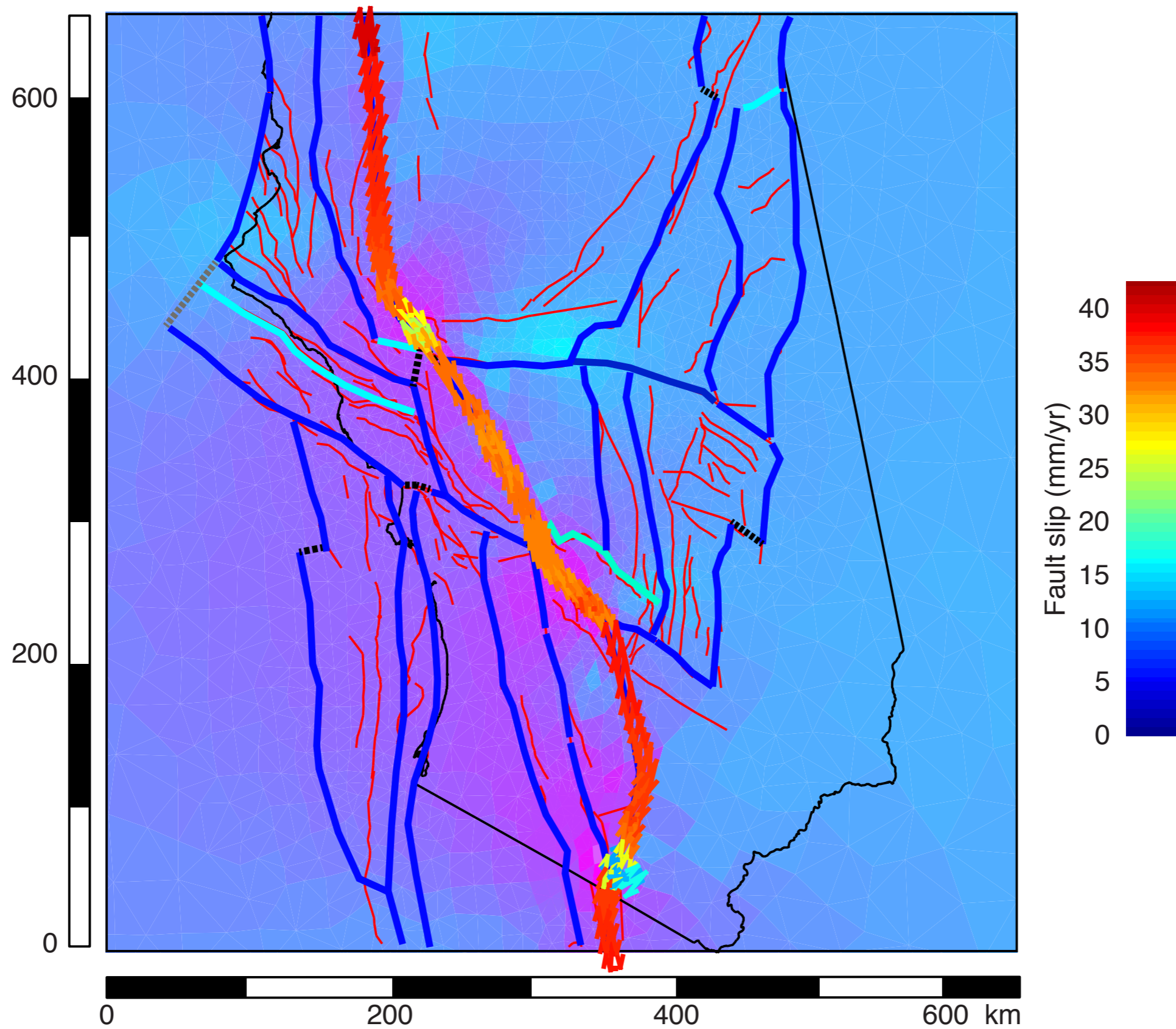
Elastic-plastic model with traction-free faults and a stepover



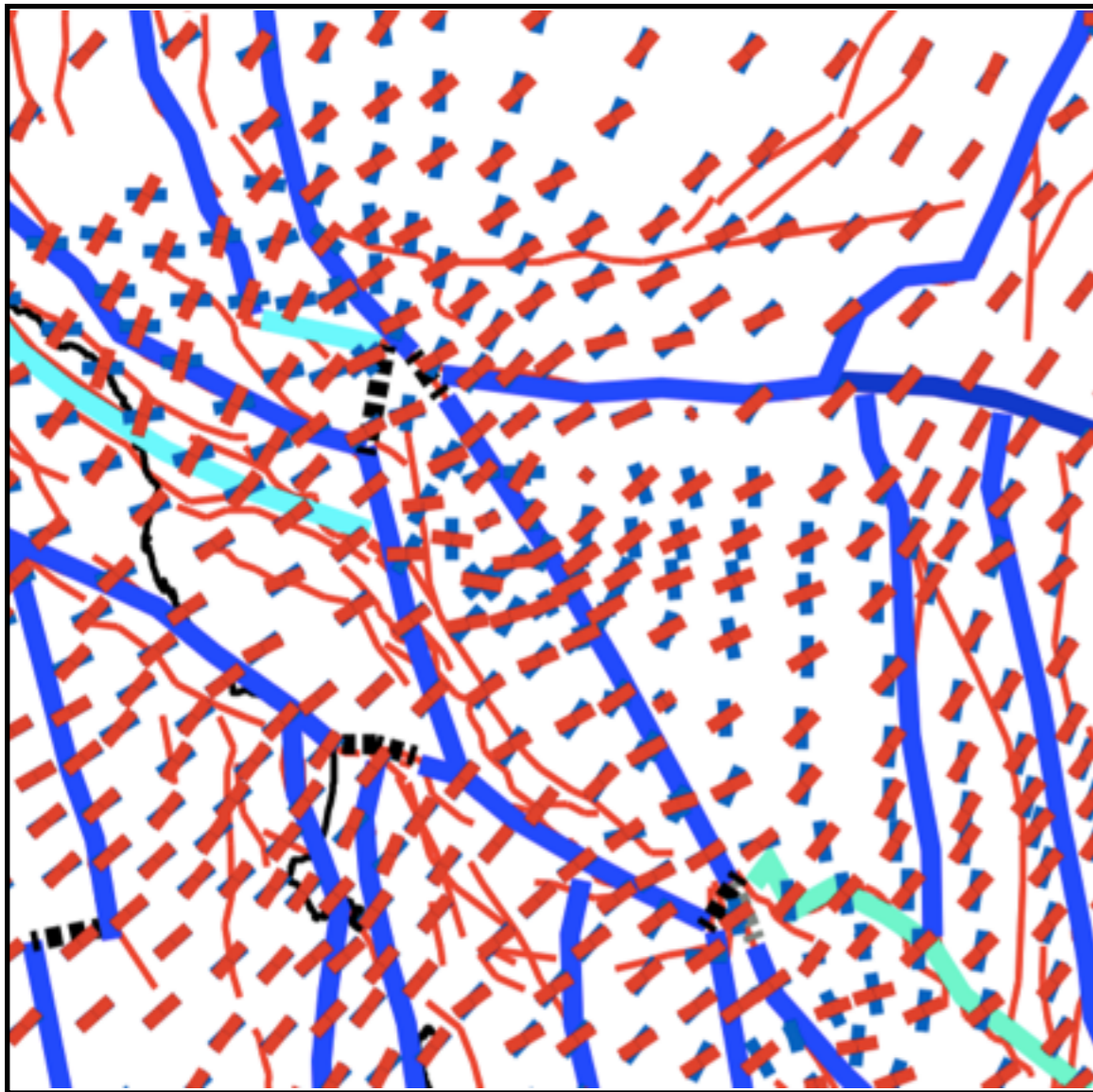
Elastic model with traction-free SAF: Strain energy and SAF slip rates



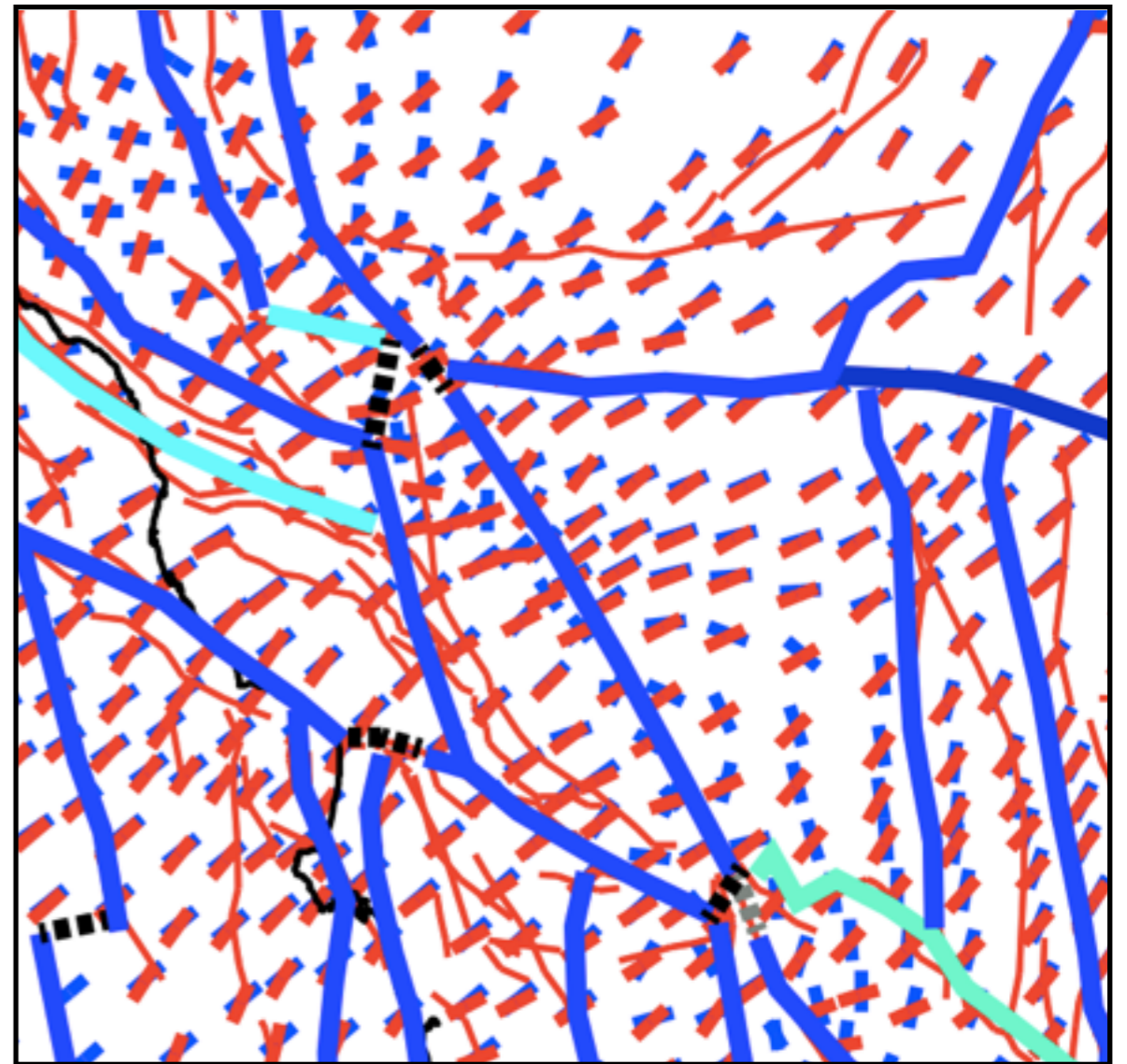
Elastic-plastic model with traction-free SAF: Strain energy and SAF slip rates



Elastic model



Elastic-plastic model



blue: from model

red: from seismicity (CSM contribution
of Yang and Hauksson)

SCEC5 theme: Beyond Elasticity

Rheology influences modeled stresses,
stressing rates, and fault interactions

(geologic vs. geodetic slip rates? supercycles?
earthquake gates?)

1. Viscoelastic lower lithosphere
2. Lithosphere-scale shear zones today's examples
3. Plastic upper crust