

1 **A case study integrating remote sensing and distinct element analysis to quarry slope**  
2 **stability assessment in the Monte Altissimo area, Italy**

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12

13 **Abstract**

14

15 Over last decade geomatic techniques have been increasingly used for the geometrical  
16 characterization of rock slopes. Terrestrial laser scanning and digital terrestrial photogrammetry in  
17 particular are now frequently used in the characterization of joint surfaces and slope geometry.  
18 Although the use of these techniques for the structural characterization of slopes is widely  
19 documented, limited research has been undertaken to improve our understanding of the importance of  
20 the derived data quality in the construction of slope geometry imported into 3D numerical models. One  
21 of the most common problems encountered in the use of these techniques, especially in case of  
22 slopes with complex geometry, is the presence of occlusions. In this context, the aims of this paper  
23 are to describe how the integrated use of terrestrial laser scanning, digital terrestrial photogrammetry  
24 and topographic surveys can mitigate the influence of occlusions and how the slope geometry gained  
25 from these surveys can be important in slope stability analyses. For this purpose a case study in the  
26 Monte Altissimo area (Apuan Alps, Italy) will be presented. Several geomatic techniques were used for

27 studying a slope overhanging the Granolesa quarry. Special emphasis will be given to the importance  
28 of using Total Station and Differential GPS surveys as tools for data fusion. Moreover, in order to  
29 validate this procedure, the accuracy and precision of the output were determined through comparison  
30 of 3D models derived from laser scanning and digital terrestrial photogrammetry.  
31 Furthermore, two different analyses with the three-dimensional distinct element code, 3DEC, were  
32 carried out in order to highlight the advantages and limitations of using data obtained from terrestrial  
33 remote sensing techniques as opposed to models based on topographic maps.

34

35 **Key words:** terrestrial laser scanning; digital terrestrial photogrammetry; intersection method; rock  
36 slope stability; 3D-distinct element models.

37

## 38 **Introduction**

39

40 Terrestrial Laser Scanning (TLS) and Digital Terrestrial Photogrammetry (DTP) are being increasingly  
41 used in the study of rock slopes (Feng and Röshoff, 2004; Abellán et al., 2006; Coggan et al., 2007;  
42 Tonon and Kottenstette, 2007; Jaboyedoff et al., 2008; Ferrero et al., 2009; Lato et al., 2009;  
43 Sturzenegger and Stead, 2009a and b; Salvini et al., 2011 and 2013). Using these techniques, it is  
44 possible to obtain very detailed information on the structural setting and slope geometry, even in the  
45 case of inaccessible steep slopes.

46 From past experience we note that there are two main aspects to consider during surveys with DTP  
47 and TLS:

48

49 - the distance between the camera/scanner and the slope

50 - the line-of-sight

51

52 Regarding DTP, depending on the camera-slope distance, different survey techniques can be chosen  
53 in order to work at an optimum scale. The scale (S), in photogrammetry, can be calculated using the  
54 following equation:

55

$$56 S_{avg} = f/D \quad (\text{Equation 1})$$

57

58 where  $f$  is the focal length and  $D$  the distance. It is evident that, to achieve an optimum scale and  
59 ensure recognition of the geological features in the slope, it is necessary either to use a high focal  
60 length or to decrease the distance between the camera and the slope. Examples of these two  
61 techniques are documented by Sturzenegger and Stead (2009b), who obtained good results using  
62 high focal lengths (up to  $f = 400$  mm). Vallet et al. (2000), Haarbrink and Eisenbeiss, (2008) and  
63 Salvini et al. (2011 and 2013), demonstrated the use of aerial vehicles to decrease the range between  
64 camera and slope. The use of an optimum photos scale in DTP is important because it influences the  
65 quality of the photogrammetric model. In this research the structural setting of the slope was studied  
66 through the use of stereo-pairs and manual stereo-restitution of features (fractures attitude,  
67 persistence and spacing). It is evident that the larger the scale of the photographs the more accurate  
68 will be the interpretation of geological features.

69 In TLS, the scanner-slope distance controls the resolution of the output (point cloud). If the laser  
70 scanner is set up close to the slope it is possible to obtain a very high resolution point cloud. In the  
71 case of a slope with complex geometry however, it may be necessary to acquire several point clouds  
72 from different scanner positions. However, it must be noted that the new generation of laser scanner  
73 tends to reduce the problem of scanner/slope distance making this technique even more attractive. In  
74 fact such scanners can reach long range and still maintain high resolution during data acquisition. As  
75 for the DTP, TLS point clouds were used to identify geological features. This step is discussed in detail  
76 in sections 3 and 4 and was performed by selecting points in the TLS model which represent the

77 surface under study. It is clear that attaining a point cloud with a high resolution makes the  
78 measurement of features easier and more precise (Ferrero et al. 2009).  
79 With regard to the line-of-sight, this parameter is important for both DTP and TLS as it controls the  
80 number of occlusions in the output point cloud data. Indeed, it is possible to use different lines-of-sight  
81 to obtain photographs and point clouds from different perspectives and thereby decrease the number  
82 of occlusions. The problem with this approach is that we obtain several point clouds and several  
83 photographs that need to be registered in the same reference system. Considering that the final  
84 accuracy of a 3D model is related to the quality of the images and the point cloud registration process,  
85 it follows that the use of a robust methodology for this process forms an essential prerequisite. In this  
86 context, one of the objectives of this paper is to show how the combined use of DTP and TLS, with  
87 Differential GPS (DGPS) and Total Station (TS), can be used to overcome the problem of occlusions  
88 and to register the data in a unique reference system. The survey's methodology presented in this  
89 research make it possible to set up several TLD and DTP stations (with different lines-of-sight) and to  
90 refer all the data to the same reference system. Such "non-static remote sensing surveys" allow  
91 overcoming of problems due to occlusions and slope orientation that generally occur using static  
92 surveys (Sturzenegger and Stead, 2009a and b; Lato et al., 2010).  
93 The quality of point clouds and photograph registration is verified through the comparison of joint  
94 attitude measurements obtained from point clouds, stereoscopic models and compass techniques.  
95 The difference in the accuracy and precision of measurements from DTP and TLS are then discussed.  
96 Data obtained from these techniques can be used for different purposes. In this research the use of  
97 these data for the analysis of the rock slope stability using 3D numerical methods is presented. To  
98 understand the advantages and limitations of these techniques in the investigation of rock slope  
99 stability, a further analysis was undertaken using a 3D model derived from a topographic map at a  
100 scale of 1:1,000 and the results of the two different analyses compared and discussed.

101

## 102 **2. Apuan Alps test site**

103

104 In this paper, the case study of Granolesa quarry is presented. The quarry is located at the foot of  
105 Monte Altissimo, a mountain in the Carrara Marble District (Apuan Alps - Italy) (Fig. 1).

106 Marble extraction is one of the most important economic resources in the area. Extraction operations  
107 have a long history dating back to the Roman Empire. In fact, ancient gravestones made of Carrara  
108 Marble, have been recently discovered and attributed to the VI – II B.C. Extraction continued  
109 throughout the Roman Empire and is well documented in several archaeological excavations and  
110 reports. Mining increased further after the advent of blasting (XVIII century) and the improvement of  
111 infrastructure. Over the last 40 years blasting has been replaced by new techniques of excavation  
112 including the use of diamond wire and diamond saw blade which has decreased the degree of  
113 excavation-induced fragmentation. Nowadays, the Carrara district is the most important marble mining  
114 area in Europe with more than 100 active quarries and a production of around 1,000,000 tons/year of  
115 blocks and aggregated gravel (Carmignani et al., 2007).

116 Regarding the geology, the Apuan Alps belong to the Northern Apennines and are a compressional  
117 fold-thrust belt formed during the Oligocene due to the collision between the Corsica-Sardinia  
118 microplate and the Italian peninsula (Alvarez et al., 1974; Kligfield, 1979; Carmignani and Kligfield,  
119 1990).

120 The stratigraphy consists of two non-metamorphic units (Tuscan Nappe and Liguride units) and a  
121 metamorphic complex. The latter is made up of two tectonic-metamorphic units, the "Autoctono" and  
122 the overlaying "Unità di Massa". The geometry of this complex is the result of two main tectonic-  
123 metamorphic events named D1 (late Oligocene-very early Miocene) and D2 (commencing in the early  
124 Miocene). During D1, there was deposition of non-metamorphic units (following the North-East  
125 direction) and compression caused by the under-thrust of the Adriatic plate. In these conditions,  
126 limestone, dated 180 My, was transformed to marble. Related to this event, is the formation of the  
127 metamorphic foliation (S1) that represents the axial plane of sheath folds which characterize the major  
128 part of the Apuan Alps. The folds range from microscopic to kilometre scale and the axes are

129 generally oriented from North–West to South–East. During the subsequent D2 phase, all the  
130 structures developed in the D1 were re-folded again in a complex anti-formal stack geometry with the  
131 presence of parasitic folds. The D2 folding process is the likely cause of the exhumation of the  
132 metamorphic complex generated during the phase D1. Subsequently during the last stages of the D2  
133 phase, brittle rather than ductile tectonics took place represented by open and kink folds and normal  
134 extension faults with a low and high angle. This extensional phase caused stretching, denudation and  
135 uplift phenomena, bringing higher and lower structures to the same level. Therefore, the Apuan Alps  
136 represent a wide tectonic window within the thickened Apennine Nappe (Carmignani and Kligfield,  
137 1990).

138

### 139 **3. Geomatic surveys**

140

141 As mentioned above, the Carrara marble district represents a very important economic resource in  
142 Italy. Around 1,000 people currently work in the extraction industry in this region. For that reason, the  
143 use of innovative technologies for the study of rock slopes, open pits and underground quarries has  
144 become very important to improve safety in work places. In this context, this research shows the use  
145 of different geomatic techniques for the analysis of the Granolesa quarry and the overhanging rock  
146 slope.

147

#### 148 3.1 Digital terrestrial photogrammetry

149

150 Digital Terrestrial Photogrammetry (DTP) is used in different fields, such as topographic mapping,  
151 architecture, engineering, and, more recently, environmental geology (Lemy and Hadjigeorgiou, 2003;  
152 Poropat 2003; Roncella et al., 2005; Redfern et al., 2007).

153

154 This is a remote sensing technique in which geometric properties of objects are determined by  
measurements of photographic images. For further details about the theory of this technique reference

155 should be made to Slama (1980), Linder (2003) and Kraus (2007). In this work this technique, together  
156 with TLS, is used to collect the geometrical parameters necessary to understand the structural setting  
157 and carry out the stability analysis of the rock slope.

158 The authors have previously discussed the use of different methods for the photogrammetric  
159 acquisition depending on the slope morphology and the accessibility of rock faces (Sturzenegger and  
160 Stead, 2009b, Firpo et al., 2011). In this case, considering the elevation and the complex geometry of  
161 the slope (Fig. 1), an aerostatic helium inflated polyurethane balloon was used (Fig 2A).

162 Four electrical winches connected to dyneema cables, pulled the balloon from the ground and enabled  
163 the remote control of height and attitude. The equipment consisted of an aluminium frame, supporting  
164 two Nikon™ D80 digital cameras at its extremities (Fig 2B). The physical CCD frame size of the  
165 cameras was 2.36 x 1.58 cm with 10.2 million effective pixels.

166 Nikkor autofocus 2.8D lenses were used with a fixed focal length equal to  $f = 20$  mm and a coverage  
167 angle of 70 degrees. Based on the sensor format, horizontal and vertical fields of view varied with the  
168 distance from the outcrops while the image size is represented in Table 1. Image acquisition was  
169 controlled by a PC-driven radio system which ensured both synchronous acquisition and correct  
170 setting of the camera shot angle in relation to the North. This function avoided problems relating to the  
171 yaw angle by maintaining right angles between the frame and the slope. A specifically-designed  
172 software package called Dragonfly (Menci Software s.r.l.) was used for this procedure. Therefore, the  
173 balloon ascended to the maximum desired height and during image acquisition remained completely  
174 stable. Furthermore, up-and-down movements of the photogrammetric equipment were controlled by  
175 pulleys so that the stereo-pairs were taken at various altitudes for each strip. Complete photographic  
176 acquisition was achieved covering the area by four vertical strips with about six stereo-pairs each. In  
177 order to avoid occlusions, these strips were acquired with different lines-of-sight (Table 2). Figure 2C  
178 shows the extent of the vertical strips adopted to cover the whole area while Figure 2D illustrates the  
179 direction of the aluminium frame during every acquisition (these directions are orthogonal to the line-  
180 of-sight).

181 The absolute orientation of the photographs was performed using the LPS module of ERDAS™  
182 IMAGINE software. With the aim to obtain very precise output, 71 natural Ground Control Points  
183 (GCP) were measured using a Leica™ TCRP 1203+R1000 reflectorless Total Station (TS). The  
184 topographic survey will be discussed in-depth in paragraph 3.3. The Root Mean Square Error (RMSE)  
185 in image units (pixel) obtained for each strip during the absolute orientation process is shown in Table  
186 2. As discussed previously the photogrammetric image scale depends on focal length and camera-  
187 slope distance. Considering the geometry of the slope and the fact that the acquisition with the  
188 aerostatic balloon proceeded vertically with fixed focal length, the scale changed in relation to the  
189 different zones of the slope. In the lowest area, where the camera-slope distance was closer, the scale  
190 was approximately 1:180 whereas, at the top of the slope it was around 1:300. Based on the data in  
191 Tables 1 and 2, and on the scale of the photogrammetric images (similar in every strip), it was  
192 possible to calculate the RMSE in ground units (cm). The results of this calculation are also shown in  
193 Table 2.

194

### 195 3.2 Terrestrial laser scanning

196

197 Terrestrial Laser Scanning (TLS) is an innovative survey technique for rapidly obtaining the geometry  
198 of slopes at high precision. Several types of laser scanner currently exist with different measurement  
199 principles and technical specification (Beraldin, 2004; Fröhlich and Mettenleiter, 2004). In this paper, a  
200 Leica™ ScanStation2 laser scanner was used for the survey. This instrument uses the time-of-flight  
201 technology to determine distances to an object and can scan to a distance of about 300 m. Nine point  
202 clouds were acquired to avoid occlusions in the output data (Fig. 3) and their resolution was set to 2  
203 cm at a distance of 100 m. Considering the proximity between the laser scanner and the slope during  
204 the surveys a slope resolution of around 1 cm was achieved.

205 One of the most difficult steps in using TLS is the registration of point clouds. This process allows for  
206 the integration of several point clouds into a unique reference system. To overcome this problem, 49



207 High-Definition Surveying (HDS) targets were used during the data acquisition. These targets are  
208 automatically recognized by Leica™ Cyclone software (used for the TLS data post-processing) so that  
209 the registration can be easily computed by assigning them absolute coordinates. In addition, the TS  
210 was used to obtain very accurate measurements of the targets. This methodology and the registration  
211 process are discussed in-depth in the following section and in the discussion.

212

### 213 3.3 Total Station and intersection method

214

215 In this research use of a Total Station played a very important role allowing for the registration of  
216 photographs and point clouds in the same reference system.

217 A Total Station is an electronic/optical instrument used for topographic surveying which over recent  
218 has been increasingly used in mining surveying and civil engineering (Feng et al., 2001; Gan, 2011,  
219 Francioni et al., 2013). The Total Station represents an evolution of theodolite technology which  
220 through measurement of angles and distance and the use of trigonometry allows determination of the  
221 relative coordinates of surveyed points (X, Y and Z). To determine the absolute coordinates of a point  
222 (northing, easting and elevation), the TS must be set up at a known point and a reference direction for  
223 angle measurement (generally North) is required. The use of DGPS is necessary for both the  
224 measurement of TS origin and the calculation of the reference direction.

225 The TS, like other remote sensing techniques, has the problem of occlusion during the survey. In fact,  
226 to acquire the coordinate of a point, a direct line-of-sight must be established between the instrument  
227 and that point. When, a direct line-of-sight cannot be established, the instrument has to be set up in a  
228 new position. As discussed above, if the TS is moved to new locations, the positions of the instrument  
229 and the reference direction have to be calculated for each new station. In this paper, the Intersection  
230 Method (IM) was used to overcome this problem. IM is a technique for determining the position of the  
231 TS and the reference direction starting from at least two previously acquired known points (Fig. 4).

232 Calculation of the unknown point is based on trigonometry and, today, this can usually be  
233 automatically computed in the field if the TS is correctly set up. Using this technique and efficient work  
234 planning the DGPS only needs to be used at the first station (master).  
235 In this research, the master was taken as station 1 (Figs. 3 and 5). From this station, several reference  
236 points were acquired so that a direct line-of-sight, with at least four of these points, was guaranteed for  
237 all the other planned stations (Fig. 3). Using this procedure, it was possible to set up the TS in nine  
238 different locations (Fig. 3 - the same station used for the TLS survey) and acquire the HDS targets and  
239 the natural GCPs within the same reference system. In total, the coordinates of 13 reference points  
240 used for IM, 71 GCPs for image orientation and 49 HDS targets for point cloud registration were  
241 acquired (Fig 5).

242

### 243 3.4 Differential GPS

244

245 In order to obtain the origin of the TS master station and establish the precise reference direction of  
246 the TS with respect to the North, a DGPS survey was performed using a Leica™ Viva GPS receiver  
247 (Fig. 5). The DGPS measurement was corrected by the use of Leica™ Geo Office software using data  
248 simultaneously recorded by the Lucca (LU) and Borgo a Mozzano (LU) permanent stations and the  
249 Root Mean Square Error (RMSE) achieved was equal to 2 cm. Due to the data post-processing, it was  
250 possible include in the same global reference system the coordinates of reference points for IM, HDS  
251 targets and natural GCPs.

252

## 253 **4. Data processing and deliverables**

254

255 The HDS targets and GCPs were referred to a unique system with a maximum IM error of 8 mm. From  
256 the absolute orientation of the photographs shown in Table 2, the stereoscopic models of the slope  
257 were derived and utilized for the stereo-restitution of geological features.

258 Based on the acquisition of the HDS targets, the registration of the nine point clouds in a one complete  
259 3D model was possible with a mean error of about 5 mm.

260 Clearly, the construction of the 3D model provided data redundancy in some parts of the model. This  
261 problem was overcome by the use of a filter in Leica™ Cyclone software which allows for defining a  
262 minimum distance between two points. In so doing, the software removes the points that are closer  
263 than the specified range making possible the creation of a model with a reasonably constant  
264 resolution. In this case, a maximum resolution of 3 cm was chosen to allow easier and more precise  
265 extraction of geological features. The resolution was then decreased up to 30 cm during the creation  
266 of the 3DEC model. This was considered a good compromise to measure the complexities of the  
267 slope profile without building a too complex model geometry for use in the 3D numerical code. As the  
268 HDS targets and CGPs were in the same reference system, it was possible to merge the  
269 photogrammetric block with the TLS 3D model and obtain high resolution orthophotomosaics of the  
270 slope (Fig. 6).

271 As already documented by several authors (Krosley et al., 2006; Coggan et al., 2007; Sturzenegger  
272 and Stead, 2009a and b; Firpo et al., 2011; and Salvini et al., 2013) geological features such as block  
273 size and shape, and joint attitude and persistence can be restituted from DTP and TLS models.

274 In the stereoscopic model, surfaces restituted by the Stereo Analyst module of ERDAS™ IMAGINE  
275 were represented by triangles drawn co-planar with the discontinuities (Figs. 7A and B). Their dip and  
276 dip direction were calculated using spatial analysis techniques with Esri™ ArcInfo Workstation  
277 software. In fact, using this workstation it was possible to perform a GIS analysis which defined the  
278 joint attitudes through the calculation of “Aspect” (dip direction) and the “Slope” (dip) of each triangular  
279 surface.

280 It must be highlighted that manual stereo-restitution of feature, although providing a high level of data  
281 interpretability during the stereo-restitution, is prone to human error.

282 In contrast, on the point cloud, the attitude of the surfaces was obtained by Leica™ Cyclone using a  
283 semi-automatic procedure. In practice, several points (at least three) representing the surface under

284 investigation were selected on the point cloud, and the software then automatically recreated that  
285 surface (Fig. 7C). The algorithm used for this procedure is a best-fit algorithm developed by Leica  
286 Geosystems. The spatial vector components  $u_x$ ,  $u_y$  and  $u_z$  were obtained from this surface and,  
287 through the use of Equations 2 and 3, it was possible to calculate the dip direction and dip  
288 respectively:

289

$$290 \text{ Dip direction} = \arctan (u_x/u_y) + \theta \quad (\text{Equation 2})$$

291

292 where:

293

$$294 \theta = 0^\circ \text{ for } u_x \geq 0 \text{ and } u_y \geq 0$$

$$295 \theta = 360^\circ \text{ for } u_x \leq 0 \text{ and } u_y \geq 0$$

$$296 \theta = 360^\circ \text{ in all the other cases}$$

297

298 and

299

$$300 \text{ Dip} = \arctan [u_z/(u_x^2 + u_y^2)^{-1}] \quad (\text{Equation 3})$$

301

302 In this study, a validation procedure was followed in order to assess the accuracy of the joint attitudes  
303 extrapolated from TLS and DTP models. The measurements from the two proposed geospatial  
304 methodologies were compared with data from traditional engineering geological compass surveys.  
305 Eight surfaces were chosen at the quarry face for this purpose and, to avoid affecting the results with  
306 a single error during the measurement, every surface was drawn ten times and the mean value used  
307 for the comparison.

308 Table 3 shows the results of this validation procedure with the mean attitude values determined from  
309 the compass, DTP and TLS measurements. Moreover, the box plot representation in Figure 8

310 highlights the variation in degrees between compass and remote sensing measurements (differences  
311 represented in absolute values). The high accuracy and precision of the TLS and DTP is evident from  
312 this box plot where mean values are similar to the compass measurements and the standard deviation  
313 is always less than 10 degrees. Nevertheless, the graph also highlights the higher precision of the  
314 measurements carried out on the TLS model. This is mainly due to the very high resolution of the point  
315 clouds (around 1 cm of resolution on the slope) and to the methodology used for the extraction of the  
316 dip and dip direction. In fact, as discussed above, the procedure used for the attitude determination for  
317 the surfaces in the TLS model is semi-automatic and reduces human error during the surface  
318 extrapolation. In contrast, the surfaces on the stereoscopic model were drawn manually so that the  
319 precision of the measurements is related to the experience and skill of the drawer and tend to be less  
320 precise.

321 It is important however to considerer the flexibility of the DTP when comparing it to TLS. In fact, even  
322 in the worst case scenario when the dip direction of fractures is orthogonal to the line-of-sight, it is  
323 possible from the stereoscopic model to interpret of their attitude. In contrast, this can be very difficult  
324 using the model derived from TLS, especially when the resolution of point clouds is lower than the  
325 fracture aperture.

326 Data from the engineering geological survey, DTP and TLS were all subsequently integrated to define  
327 the structural setting of the slope. The attitudes of more than 500 joint surfaces were measured using  
328 the three surveying techniques. Joint persistence and spacing were measured in the field and on the  
329 stereoscopic model by 3D linear representation whereas the aperture, Joint Compressive Strength  
330 (JCS - Deere and Miller, 1966) and Joint Roughness Coefficient (JRC - Barton, 1973) were obtained  
331 only from the engineering geological survey. Table 4 shows the results gained from this study.

332

## 333 **5. Remote sensing data and 3D Distinct Element Code for the stability analysis of the slope**

334

335 We have previously illustrated how it is possible to use remote sensing techniques for the geometrical  
336 and structural characterization of rock slopes. This information was used in the stability analysis of the  
337 rock slope with the three dimensional distinct element code, 3DEC, (Itasca<sup>TM</sup>, 2008). This code uses  
338 an explicit time-stepping scheme to solve the Newton's equation of motion and it treats the rock mass  
339 as a discontinuum material where the joints are the main control in the analysis of the rock mass. The  
340 discontinuities cut the rock mass in blocks which can be assigned rigid or deformable stress-strain  
341 constitutive criteria depending on the rock mass characteristics. Different constitutive criteria can also  
342 be assigned to the discontinuities which are treated as boundary conditions along which large  
343 displacements and rotations are permitted. Cundall (1988), Hart et al. (1988) and Itasca<sup>TM</sup> (2012)  
344 describe the principles of distinct element code and 3DEC in more detail. This software has been used  
345 by several authors recently in both case studies and conceptual models (Yeung and Wong, 2007;  
346 Brideau and Stead, 2010; Kalenchuk, 2010; Brideau and Stead, 2011 and 2012; Firpo et al., 2011).  
347 In order to make the data amenable for incorporation into 3DEC pre-processing procedure is required  
348 which is described below. Moreover, to understand the advantages and limitations of using terrestrial  
349 remote sensing techniques for the investigation of slopes, two different 3DEC analyses will be  
350 presented; the first using a 3D model derived from TLS and, the second, derived from a topographic  
351 map with a scale of 1:1,000. The choice of using the TLS point cloud rather than the eventual DTP  
352 point cloud derived from image matching algorithms and manual stereo-restitution, was based on  
353 several reasons. Although software with image matching algorithms has allowed very high precision in  
354 the construction of 3D models in the last decade, the results of the calculations are strictly related to  
355 the quality of the photographs. Therefore, the complexity of the slope geometry, experience in the use  
356 of aerostatic balloons, the results of the comparison between TLS and DTP and the actual availability  
357 of a very precise TLS model have all led the authors to use the latter for the provision of data for  
358 analysis using the 3DEC code.

359

### 360 5.1 Creation of the slope geometry in 3DEC

361

362 The Rhinoceros™ SR<sup>4</sup> (CAD suite developed by McNeel and Associates, 2011) was used to manage  
363 the TLS point clouds and create both the 3D model (Fig. 9A) and triangulated mesh. As mentioned in  
364 the previous section, the TLS point cloud was resampled before importing into the 3DEC model. A  
365 resolution of 30 cm was retained and considered sufficient by the authors to include all the relevant  
366 scale geometrical variations in the slope (Fig. 9A). The mesh obtained from this model was then  
367 exported and made compatible with 3DEC using the Itasca™ software, Kubrix (Fig. 9B).  
368 Using this software, it was possible to create a 3DEC model closely approximating the true geometry  
369 of the slope. The structural setting of the slope could have been defined in 3DEC using the data  
370 presented in Tables 3 and 4. In this case, however, the purpose of the analysis was to highlight the  
371 advantages of using a model created from geomatic techniques. Moreover, the joint geometry  
372 achieved from the above mentioned surveys, represents the situation on the slope face but it is  
373 suggested that this may rarely be fully representative of the actual situation at depth inside the slope.  
374 Therefore the authors in these preliminary models decided to use a joint spacing three time wider than  
375 the measured value (1.1 m for K1, 2.3 m for K2 and 7.9 m for K3). This enabled practical computer  
376 runtimes without changing the relative spacing between joints and compromising the objectives of the  
377 research.

378 In order to understand the advantages and limitations of using terrestrial remote sensing techniques, a  
379 further 3DEC analysis was carried out using the slope geometry obtained from the topographic map at  
380 a scale of 1:1,000 (Figs 9C and D). The procedure used to create this additional 3DEC model was  
381 similar to that outlined for the TLS data.

382

### 383 5.2 3DEC analysis

384

385 As discussed above, the geometric models created from TLS and the topographic map were used for  
386 the 3DEC analyses. In order not to introduce increased uncertainty in the calculation and to focus the

387 study on the role of measured slope geometry on slope deformation, the same typical joint properties  
388 were assumed for all the joint sets. Three simulations were carried out for both models using different  
389 joints properties with the aim to investigate the models susceptibility to instability (Table 5).

390 Considering that the entire slope comprises marble with very high strength properties and that the  
391 principal focus of this research was to analyse the role of different slope geometries with respect to  
392 joint behaviour, the rock mass blocks were assumed to be rigid (non-deformable with a density equal  
393 to 2,700 Kg/m<sup>3</sup>).

394 A comparison of the response of the two models was undertaken using the Shear Strength Reduction  
395 Factor (SRF – Dawson and Roth, 1999; Griffiths and Lane, 1999) where the shear strengths of all the  
396 component materials were iteratively decreased until the slope failed. Clearly, considering that the  
397 rock mass material was assumed to be rigid, the SRF was applied only to the joint properties. The  
398 results of this comparison are shown in Table 6 and Figure 10 where it is possible to see how, in the  
399 TLS model, the SRF is consistently lower than in the topographic model.

400 As all the input parameters used in the 3DEC analyses were the same for both models, it is evident  
401 that the differences in the SRF were due to the geometry. In fact, although the topographic map at a  
402 scale of 1:1,000 may be a reliable representation of morphology it is unable to represent some slope  
403 peculiarities such as the overhanging geometry present in this case-study. Therefore, the 3D model  
404 achieved may not always be fully representative of the actual localized slope geometry. Figure 11  
405 highlights the differences between the geometry extracted from the topographic map and the TLS.

406 Figures 11A and B show the 3D model from the topographic map and a 2D section through it  
407 respectively whereas Figures 11 C and D show the same for the TLS model.

408 The slope in the first two images (Fig 11 A and B) is fairly regular with a steep slope while, in the  
409 images in Fig 11 C and D, it is characterized by a more complex morphology with an overhanging  
410 geometry.

411 This difference in the local slope morphology is most probably the cause of the SRF differences  
412 obtained in the analyses. In fact, this overhanging rock mass has a volume of about 14 m<sup>3</sup> (calculated



413 in Rhinoceros - McNeel and Associates, 2011) which is not evident from the topographic model.  
414 Clearly, the overhanging rock volume can influence the results of stability analysis and consequently  
415 the SRF due to the necessity to consider the existence of a key block and block theory (Goodman and  
416 Shi, 1985).

417

## 418 **6. Discussion**

419

420 This paper evaluates the importance of the integrated use of topographic and remote sensing  
421 techniques in slope stability analysis. The methodology described attempts to overcome a frequent  
422 problem related to the accuracy and completeness of data representing the rock slope morphology.  
423 Special emphasis is given to the use of IM as a tool for data integration and the role that different  
424 geometric models can have in the results of 3D slope stability analyses.

425 With regard to the photogrammetric survey, the authors show how the use of an aerostatic balloon in  
426 the investigation of the quarry and the overhanging slope is able to overcome problems related to the  
427 slope height and the complex geometry of the rock slope face. Table 2 shows the errors obtained  
428 during the orientation process; these are considered acceptable in view of the complexity of the  
429 survey. Although the use of the aerostatic balloon was important to avoid hidden zones, the use of this  
430 tool for the photogrammetric survey tends to be complicated. In fact, the instrumentation needs to be  
431 set up over a large area at the foot of the slope and the maximum height that can be achieved is about  
432 300 m. Moreover, the wind and the cost of the helium can represent additional limitations in the use of  
433 this tool. The advent of new technologies, like Unmanned Aerial Vehicles (UAV) (Haarbrink and  
434 Eisenbeiss, 2008; Niethammer et al., 2010; Tahar et al., 2012; Danzi et al., 2013; Salvini et al., 2014)  
435 and the development of new photogrammetric survey techniques including the use of long range focal  
436 length (Sturzenegger and Stead, 2009b), can overcome these problems making DTP even more  
437 attractive and easy to use.

438 Regarding the TLS, it was used to obtain a very high resolution 3D model and to define a new  
439 approach for the point cloud registration. Several scans were acquired using different lines-of-sight to  
440 avoid occlusions at the rock slope face and the IM was used to join the point clouds into a unique  
441 model.

442 IM played an important role in this research as it allowed for measuring a large number of GCPs and  
443 HDS targets in the same reference system. In so doing it was possible to acquire photographs (DTP)  
444 and point clouds (TLS) using different lines-of-sight and to combine all the data into the same  
445 reference system.

446 Using this “non-static survey” methodology it was possible to overcome commonly reported problems  
447 associates with the use of DTP and TLS. As discussed by several authors (Sturzenegger and Stead,  
448 2009a and b; Ferrero et al., 2010; Lato et al. 2010), the use of remote sensing techniques can lead to  
449 problems related to slope orientation and differences in the point cloud density. Ferrero et al., 2010  
450 described the use of a new software code to determine of the rock discontinuities by means of the  
451 application of a segmentation algorithm. They highlight the advantage of using this new approach and  
452 how the accuracy of derived data always increases with a higher point density. Sturzenegger and  
453 Stead (2009 a and b) and Lato et al. (2010) highlight the importance of the line-of-sight in the use of  
454 DTP and TLS and how the use of “static TLS surveys” leads to problems related to the slope  
455 orientation. In these situations, Lato et al. (2010) suggested the use of mathematical correction to  
456 overcome these problems or, when possible, the use of multiple TLS stations.

457 The quality of the DTP and TLS output (gained with this “non-static TLS survey)’ was validated by  
458 comparison of measurements undertaken on the same structures on both models. Measurements  
459 showed a maximum variance of 8° and 3° for the DTP and TLS respectively compared with the  
460 attitude obtained using a conventional engineering compass. This result agrees with previous  
461 comparative studies carried out by Coggan et al., (2007). The DTP variance may partially be due to  
462 the error in the absolute orientation of the images (Table 2) but, is considered primarily to be the result  
463 of possible human error during the stereo-restitution of the 3D features. Different photogrammetric

464 techniques (Kemeny and Post, 2003; Roncella et al., 2005) and software (e.g. 3DM Analyst software –  
465 Adam Technology 2012) can overcome this problem creating a very accurate 3D model from the  
466 photogrammetric block, where it is possible to use the same joint extrapolation technique as used in  
467 the TLS model. Sturzenegger and Stead (2009b) showed that with careful orientation the attitude of  
468 discontinuities extracted from the photogrammetric model can have a similar accuracy to that obtained  
469 from TLS.

470 However, in this paper the results achieved from the validation process confirmed the good quality of  
471 the data that was used to define the geometry and the structural setting of the slope, and were used to  
472 analyse the stability with 3DEC. With the principal aim of highlighting the advantages and limitations of  
473 using terrestrial remote sensing data, two different slope models were analysed with the 3DEC code.  
474 The first was obtained from the topographic map and the second from TLS. The same rock mass and  
475 joint properties were assumed in both models and three different analyses were done with decreasing  
476 joint properties (Table 5). The reduction of joint properties caused a decrease in the SRF (Table 6).  
477 Moreover, careful examination of the graph in Figure 10 shows important differences in the behaviour  
478 between the TLS and topographic model. In fact, in the three analyses carried out, the TLS model  
479 consistently has a lower SRF and, in the third analysis, the SRF falls to below 1. As mentioned  
480 previously this behaviour is related to the different geometry of the models and, in particular, to the  
481 overhanging geometry that characterizes the top of the quarry in the TLS model.

482 A clear example is illustrated in the Figure 11 where it is possible to see the difference in the two  
483 analyses in terms of the sliding process. The morphological profile represented in Figure 11B shows  
484 how the hypothetical failure resembles either a planar or wedge type mechanism whereas, in Figure  
485 11D, rock fall may be indicated. It must be emphasized that a similar overhanging morphology is not  
486 evident from the 2D topographic map.

487 Considering the limitations of the proposed methodology, it should be noted that the data  
488 processing/interpretation time when dealing with accurate and detailed data are longer and the model  
489 construction more complex.

490 Such detailed slope geometry is useful in complex morphologies but may be unnecessary in simple  
491 slopes where an high scale topographic map can be suitable.

492 Another aspect to take into account is that, in the 3DEC analyses, a point of origin is needed to define  
493 the joint sets and this can play a very important role in terms of key block analysis and, consequently,  
494 SRF. It is evident that the effect of this variation is strictly dependent on the geometry of the slope and  
495 joint sets. In this case study, since the thickness of the overhanging part of the slope varies from 3 to 5  
496 meters (calculated in Rhinoceros - McNeel and Associates, 2011), it will always be wider than the  
497 spacing of the joint systems that daylight on the slope (K1 and K2). It follows that a variation of the  
498 seed point will change the volume of the failing blocks and the SRF but will not modify the kinematics  
499 and general results. Using a different seed point, the analysis with a TLS model will always show a  
500 different failure mechanism and a lower SRF than that using the topographic model.

501 The main joint sets used in the analysis were obtained from a statistical study (with an increased  
502 bedding spacing adopted to allow practical simulation times). This approach although reasonable for a  
503 probabilistic slope stability study does not represent the real situation in the slope with respect to the  
504 observed presence of discrete fractures and the influence of brittle fracturing due to blasting or  
505 excavation. Remote sensing techniques can be useful for the definition of brittle fracturing but only the  
506 rock slope face can be characterized by the use of these techniques. For that reason, a statistical  
507 analysis will be always required to generate discontinuities behind the observed rock face. The use of  
508 a stochastic discrete fracture network (DFN) model can be important to optimize the use of  
509 discontinuity data from conventional field surveys, remote sensing techniques and boreholes. In fact,  
510 in DFN modelling, it is possible define for each discontinuity set the orientation, persistence and  
511 spatial location. Several authors have shown the potential of a DFN approach for creating realistic  
512 geometric models that reflect the heterogeneity of fractured rock mass (Elmo, 2006; Pine et al., 2006;  
513 Elmo and Stead, 2010). Nevertheless, it must be considered that trace length and discontinuity shape  
514 can only be approximated and the discontinuities shape presumed; as a result, this can lead to  
515 significant uncertainty in the DFN model and related block size/shape distributions. Elmo et al. (2007),

516 Rogers et al. (2007), Rogers et al. (2009) and Stead et al. (2009) have demonstrated how the block  
517 size distributions is strictly influenced by trace length, truncation bias and discontinuity intensity.  
518 Sturzenegger et al. (2011) have recently illustrated the generation of DFN models using discontinuity  
519 parameters measured from terrestrial remote sensing. They highlighted the advantages and limitations  
520 of using this approach and suggested guidelines for the correct use of remote sensing techniques as  
521 alternative tools for the generation of advanced numerical models for slope stability analysis methods  
522 such as FracMan rock wedge analysis, Elfen and Slope Model (e.g. Cundall and Damjanac, 2009;  
523 Elmo and Stead, 2010). An integrated approach using geomechanical models that combines observed  
524 and measured discontinuities with statistically generated fractured networks may provide the most  
525 efficient remote sensing approach for the near future. Clearly, there remains a need for improved  
526 understanding of the uncertainty associated with the remote sensing measurement of rock slopes and  
527 discontinuities. The use of these techniques, combined with geostatistics is critical potential for the  
528 improved understanding of rock slope failure mechanisms.

529 In this context several important questions must be answered in the future including to what practical  
530 scale we can measure joints in a rock slope, how important is the truncation of measured joints to  
531 block size estimation and the kinematic assessment of rock slope failure mechanisms and finally what  
532 level of data will it be feasible and more importantly practical to include in future geomechanical  
533 models.

534

## 535 **7. Conclusion**

536

537 In the present research, DTP and TLS are shown to provide powerful modelling and analytical tools in  
538 the study of the geometry of a rock slope. DTP was carried out through the use of an aerostatic  
539 balloon. This technique overcomes the problem of occlusions arising from the elevation (up to 300 m)  
540 and the complex geometry of the slope. TLS was performed using a short range laser scanner. A  
541 “non-static TLS survey” was carried out to avoid hidden zones and problems related to the slope

542 orientation and point clouds density. This type of approach was possible thanks to the use of IM and  
543 provides an alternative methodology with which to overcome the biases that commonly affect remote  
544 sensing data.

545 Table 7 summarize the advantages and limitations of these remote sensing techniques based on the  
546 experience gained by the authors in this and previous researches (Sturzenegger and Stead; 2009a  
547 and b; Firpo et al, 2011; Francioni, 2013).

548 Slope geometry simulations using the 3D DEM 3DEC code were carried out with decreasing joint  
549 properties (Table 5) and with two different slope geometry models (one derived from the topographic  
550 map and the other from TLS). These simulations clearly demonstrate that the results of the stability  
551 analysis in terms of the shear strength reduction factor, are influenced significantly by the measured  
552 geometry (Table 6). This phenomenon shows that subtle variations in the slope geometry can change  
553 the simulated failure mechanism and consequently the results obtained using key block theory. Based  
554 on our results, Table 8 summarizes the improvements that remote sensing techniques can produce in  
555 the analysis of slopes with 3D distinct element codes. This is a modified classification based on Stead  
556 et al., (2006) and considers the critical input parameters in distinct element methods and, based on  
557 Francioni (2013) and this study, the potential role of remote sensing techniques in the analyses of  
558 slopes.

559 Finally, it is emphasised that this research represents an initial step in understanding the importance  
560 of slopes geometry detail in the analyses conducted using advanced three-dimensional  
561 geomechanical codes and the role that remote sensing techniques offer in acquiring improved  
562 geometrical slope details. As discussed in the previous section, further studies are needed to  
563 investigate the influence of parameters assumed in the mechanical model. Moreover the results  
564 gained from this research are related to the geometry of the slope under study and should be  
565 investigated further for varied slope geometry case studies and failure mechanisms. A detailed slope  
566 geometry is may be useful in complex morphologies but could be unnecessary in simple slopes where  
567 a high scale topographic map can be suitable and can decrease the processing/interpretation time of

568 the analyses. Each individual slope must be considered with respect to the potential importance of  
569 complex slope geometrical variations and assessed in relation to the geological structure and  
570 preliminary kinematic considerations.

571

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573

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578

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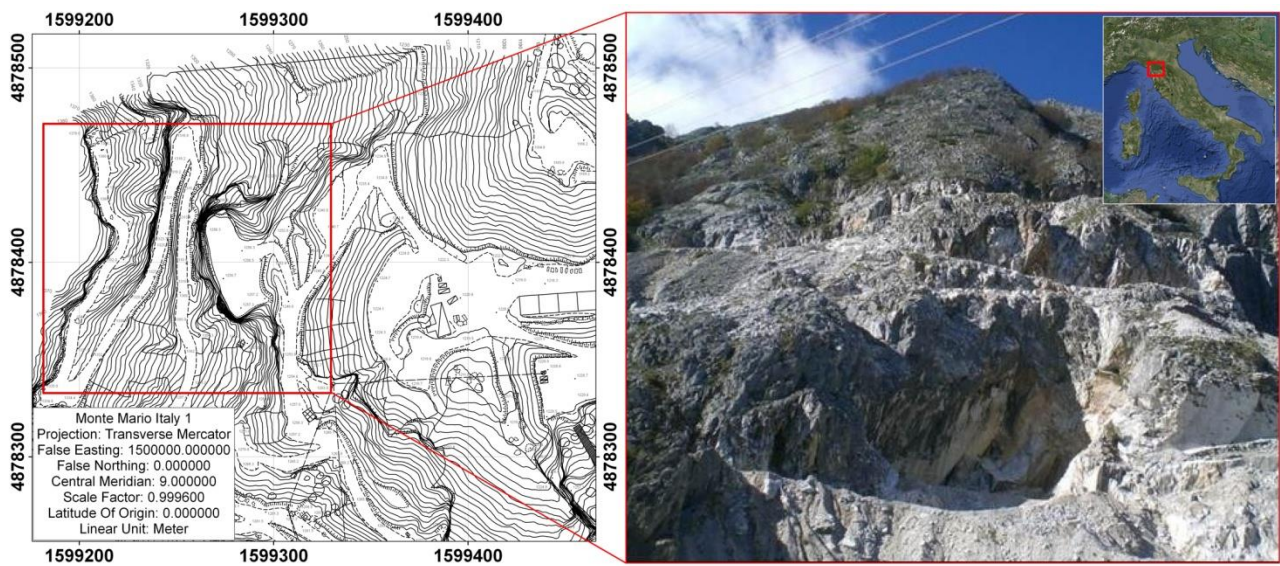
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753 Figures:

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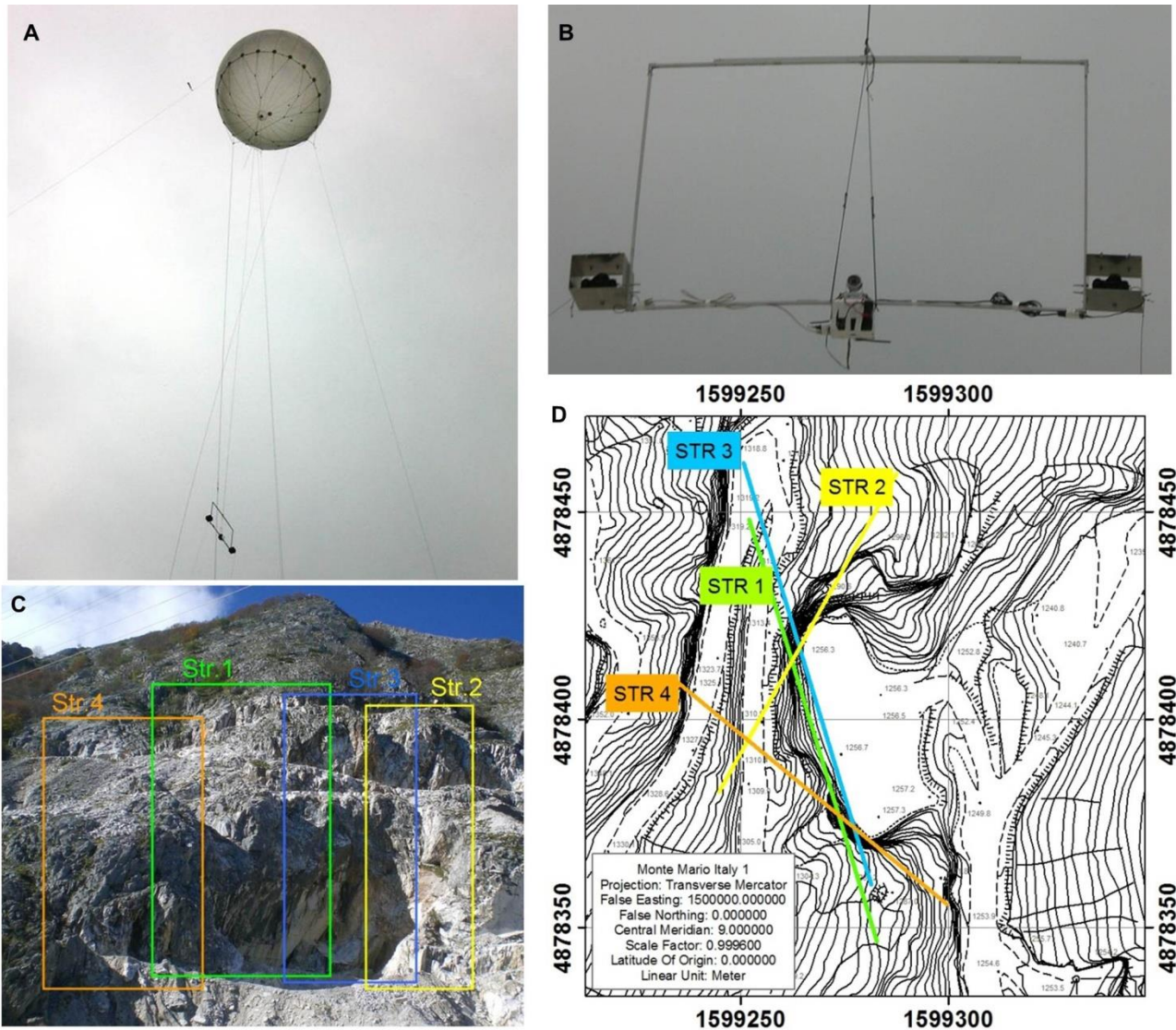
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756 Fig. 1. The Granolesa quarry. The photograph on the right shows the view of rock slope from the East  
757 to the North.

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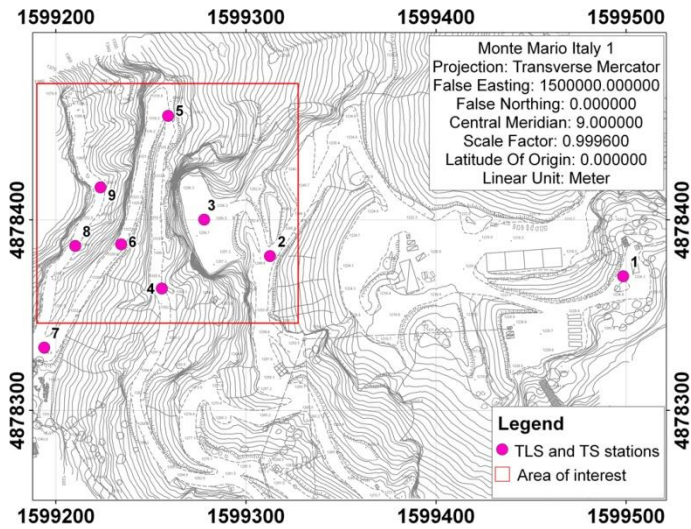
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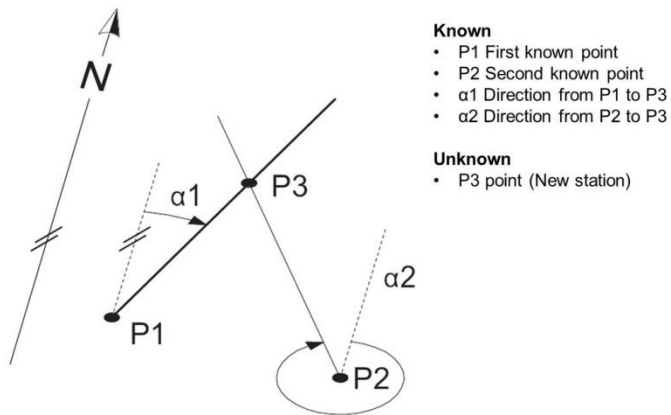
762 Fig. 2. Photogrammetric survey. A) Aerostatic balloon. B) Photogrammetric aluminium frame with  
 763 cameras. C) Vertical photogrammetric strips. D) Direction of the aluminium frame during strip  
 764 acquisition.



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766 Fig. 3. TLS and TS survey stations.

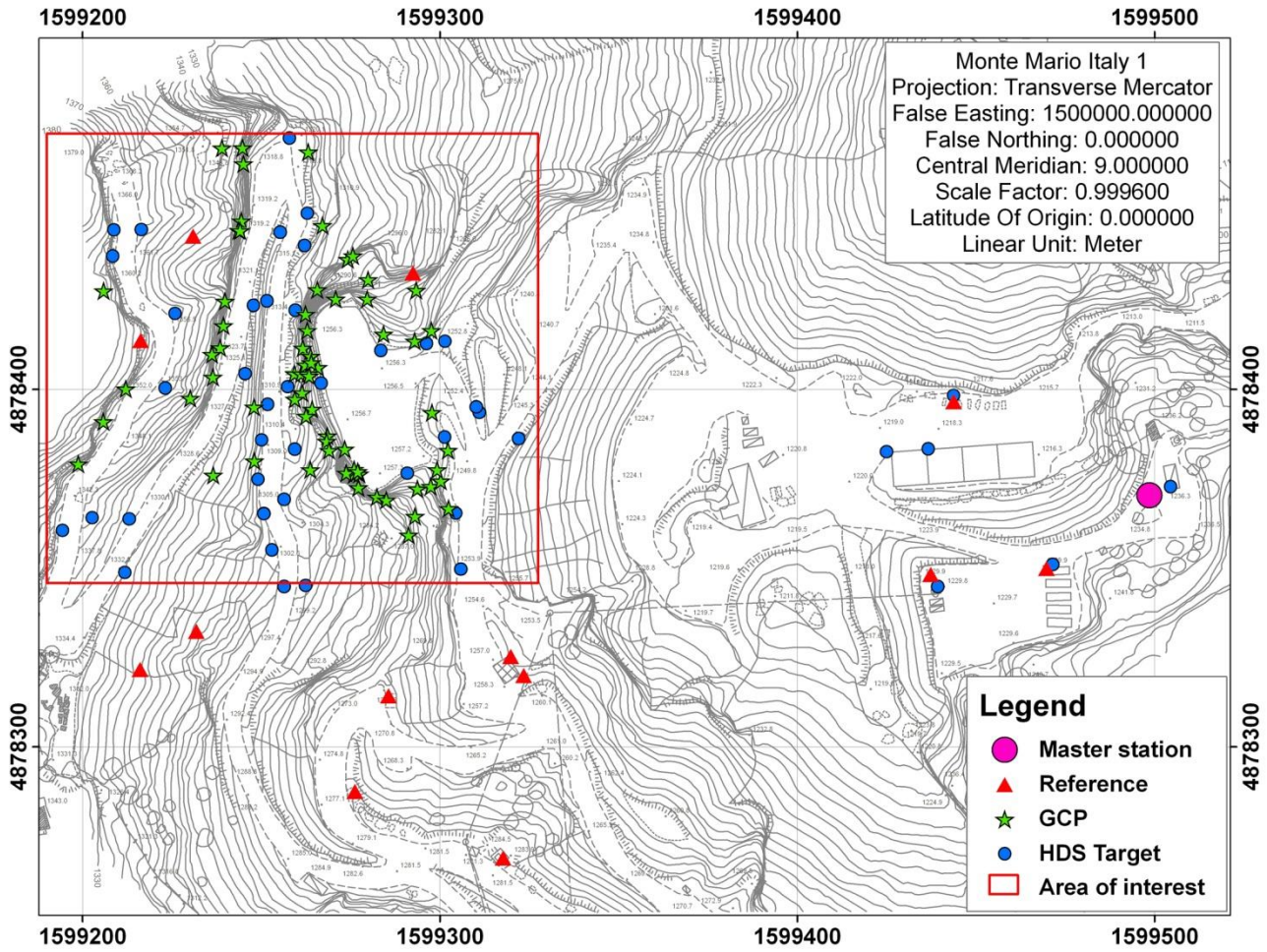
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769 Fig. 4. Intersection Method used to calculate the coordinate of an unknown point (P3) by the  
 770 acquisition of two known points (P1 and P2 - Image from Leica TPS1200+ Applications Field Manual).  
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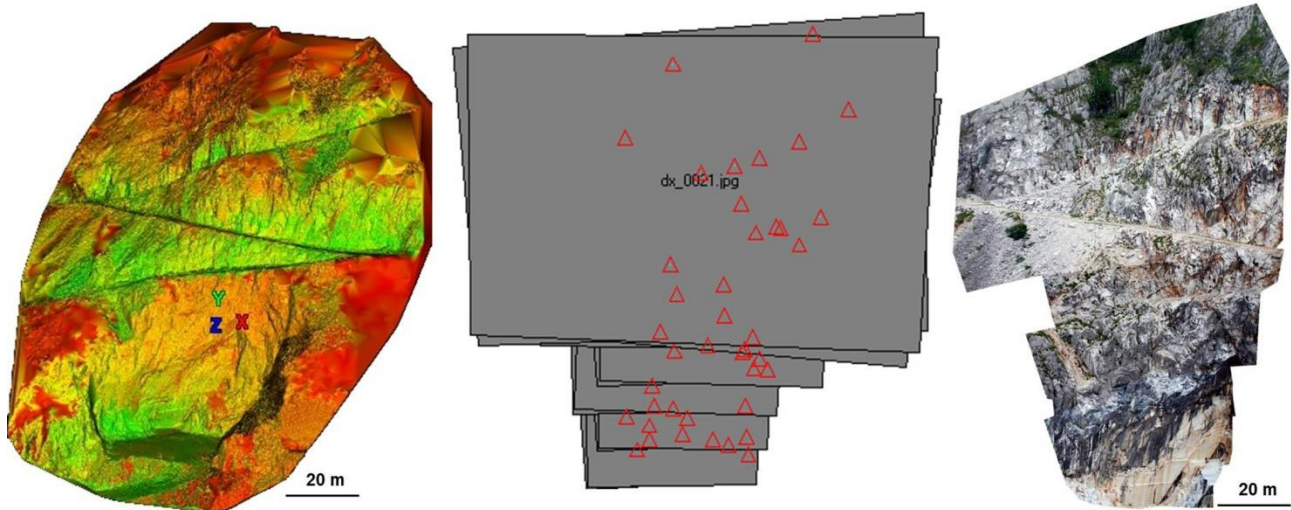




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773 Fig. 5. Map showing points measured with the Total Station and Intersection Method.

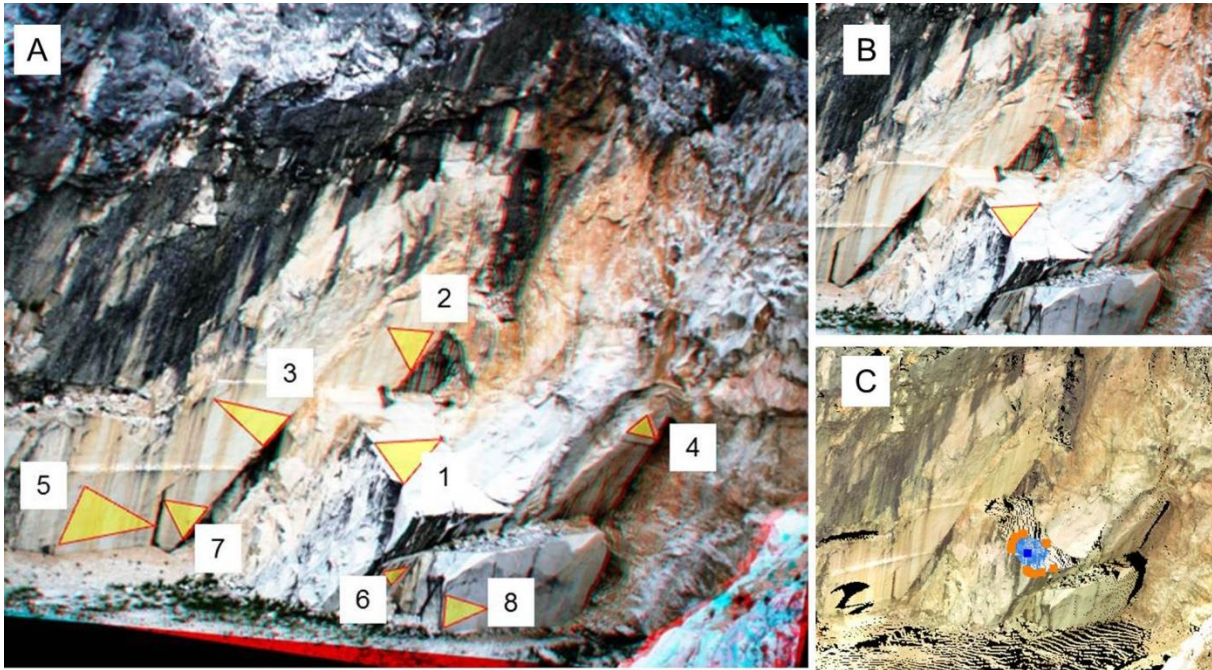
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776 Fig. 6. TLS 3D model (left), photogrammetric block (centre; triangles represent GCPs) and sketch  
777 showing high resolution orthophotomosaic (right).

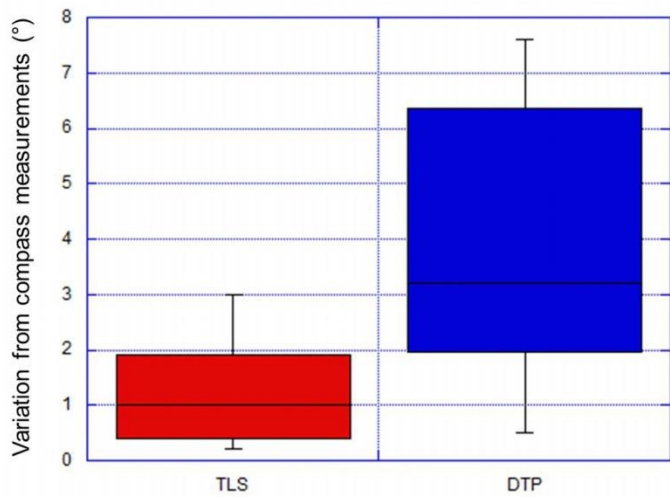
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780 Fig. 7. Measurement of the surface attitudes. A) Stereoscopic model and the surfaces used for the  
781 validation process. B) Surface 1 drawn manually on the stereoscopic model. C) Surface 1 extrapolated  
782 semi-automatically on the TLS model.

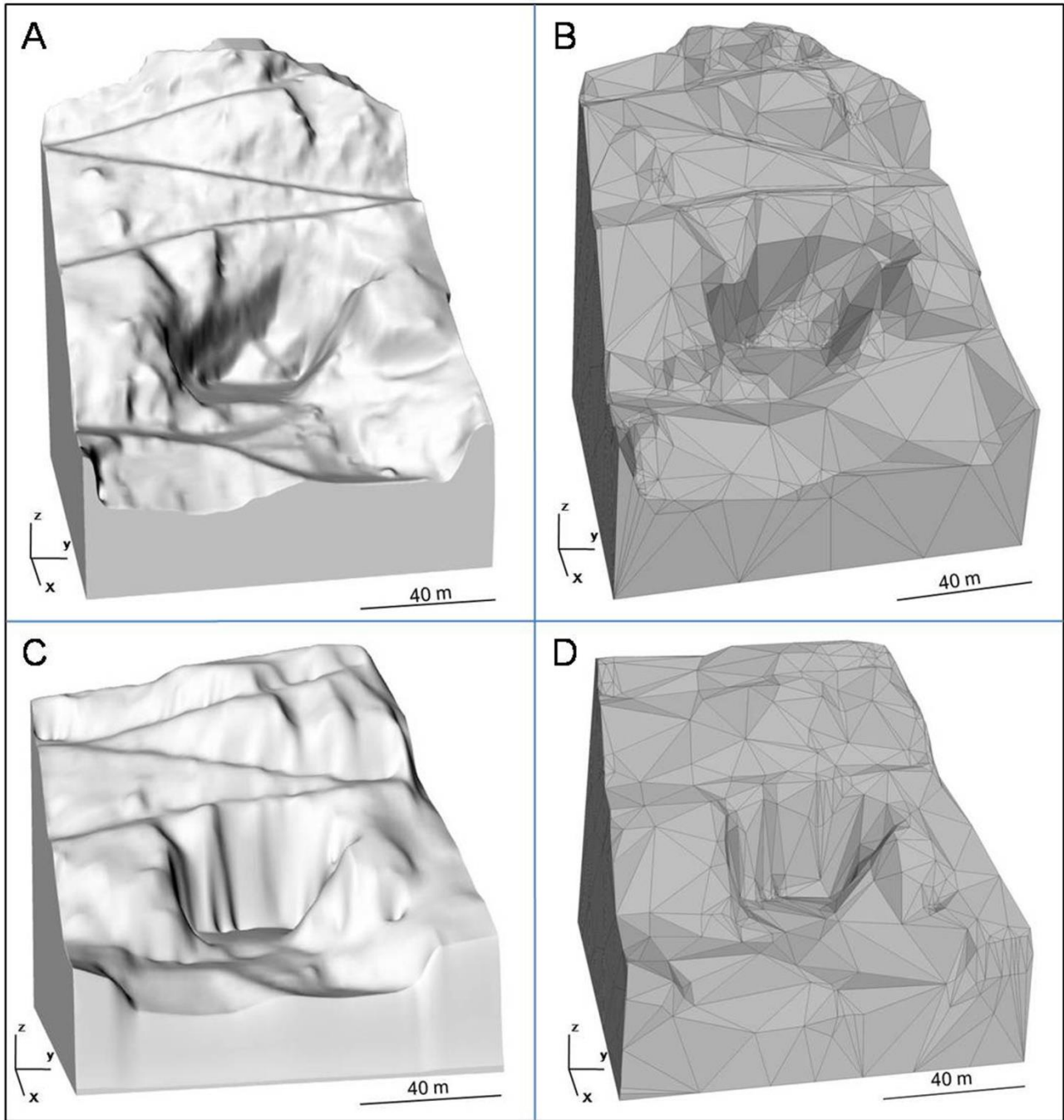
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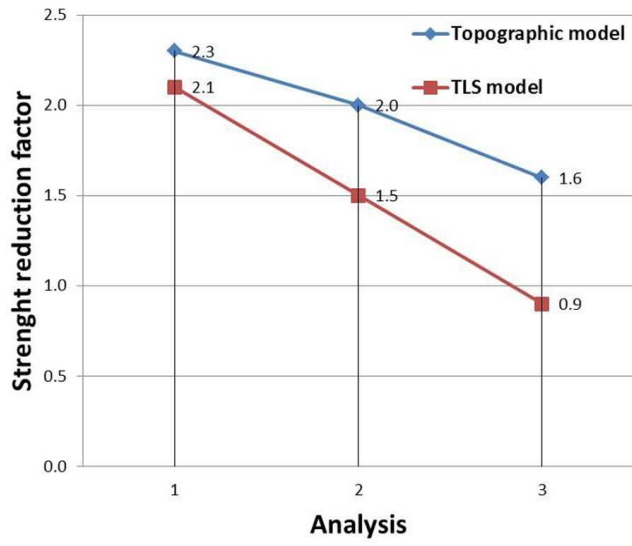
785 Fig. 8. Box plot representation. Differences between compass and remote sensing measurements  
786 (represented in absolute values - degrees).

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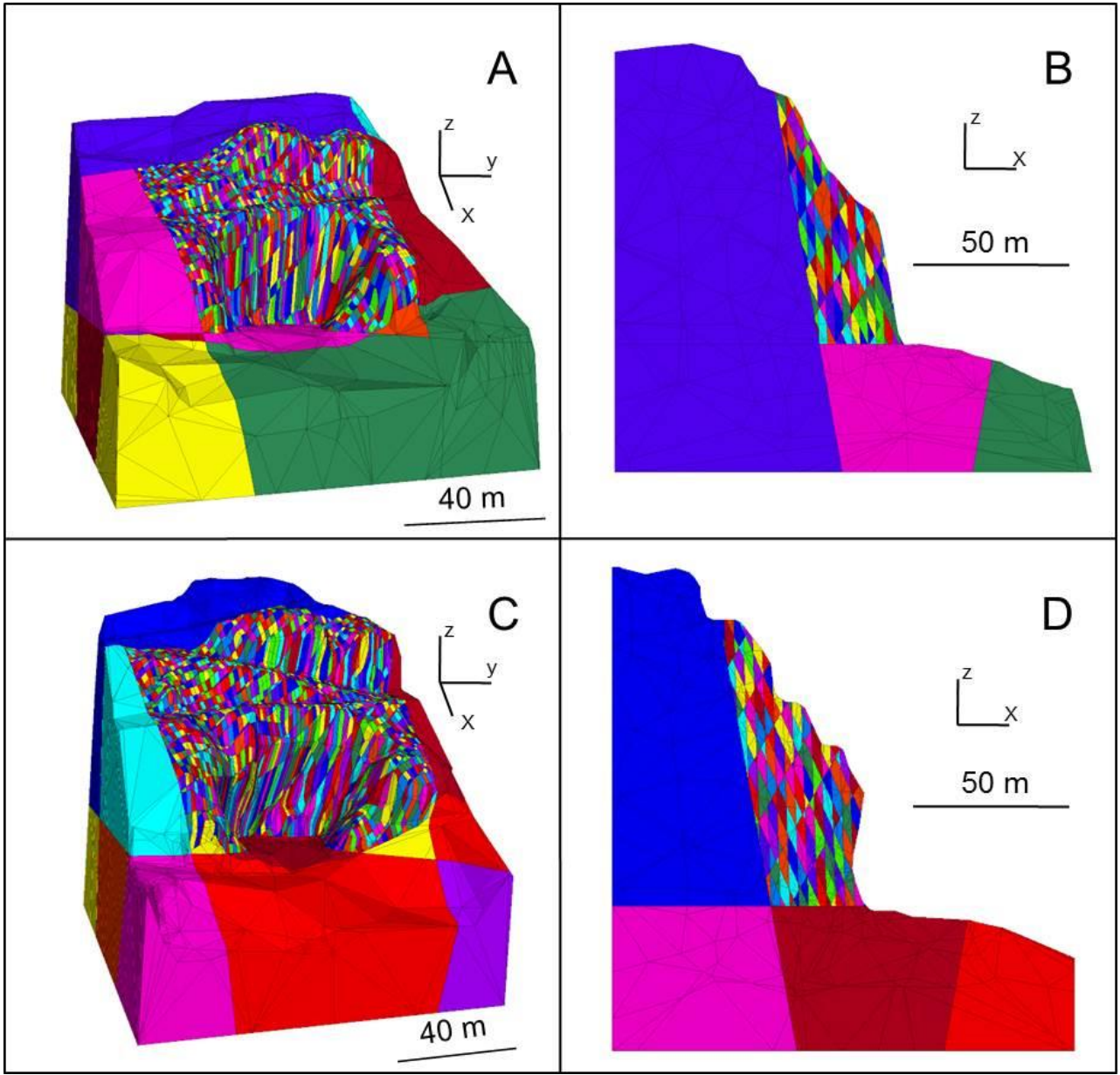
789 Fig. 9. 3D models of the quarry. TLS model created in Rhinoceros (A) and imported into 3DEC (B).  
790 Model from topographic map created in Rhinoceros (C) and imported in 3DEC (D).  
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793 Fig. 10. Change in SRF for the three analyses.

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796 Fig. 11. 3D rock slope models.3DEC model from topographic map (A) and 2D section (B). 3DEC  
 797 model from TLS (C) and 2D section (D).

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806 Tables

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	Image size (Pixel)	Image size (cm)
Width	3872	32.78
Height	2592	21.95

808 Table 1. Photographic image size

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Strip	Lines-of-sight (degrees to North)	RMSE (image unit, pixel)	RMSE (ground unit, cm) areas with highest scale (1:180)	RMSE (ground unit, cm) areas with lowest scale (1:300)
1	260	2.1	3.2	4.9
2	300	3	4.6	7.0
3	260	2.3	3.5	5.3
4	230	4	6.1	9.3

810 Table 2. Lines-of-sight and image and ground RMSE

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Surface	Compass		TLS		DTP	
	Dip Dir	Dip	Dip Dir	Dip	Dip Dir	Dip
1	65	40	65.4	38.1	64.4	36.7
2	240	75	242	73.1	245.9	71.9
3	240	70	240.9	67	243.1	62.4
4	245	70	242.2	69.6	251.9	62.8
5	240	70	239	71.1	238.1	72
6	240	65	240.4	65.2	240.5	58.2
7	240	70	240.3	70.6	245.9	68
8	245	75	244	76	240.4	75.5

812 Table 3. Dip and dip direction values of rock surfaces measured in the field with a compass, on the  
813 TLS point cloud and on the stereoscopic model (DTP).

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Joint System	Dip Direction (°)	Dip (°)	Spacing (cm)	Persistence (m)	Aperture (cm)	JRC	JCS (MPa)
K1	157	80	35.9	7.4	0.3	5	37.5
K2	79	88	75.8	3.0	0.2	4	38.3
K3	228	67	262.2	15.7	1.2	13	29.1

815 Table 4. Characteristics of the main joint systems

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Analysis number	Joint Shear Stiffness jks (Pa/m)	Joint Normal Stiffness jkn (Pa/m)	Friction Angle (°)	Cohesion (Pa)	Tensile strength (Pa)	Dilation angle (°)
1	1,00E+09	1,00E+10	30	5,00E+05	5,00E+05	5
2	8,00E+08	8,00E+09	20	4,00E+05	4,00E+05	5
3	6,00E+08	6,00E+09	10	3,00E+05	3,00E+05	5

817 Table 5. Joint properties used in the 3DEC analyses.

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Analysis number	SRF	
	Topographic model	TLS model
1	2.3	2.1
2	2	1.5
3	1.6	0.9

821 Table 6. SRF achieved from the topographic and TLS models using different joint properties.

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Close range	Advantages	Limitations
Remote sensing		
DTP with aerostatic balloon	Overcomes problems related to elevation, steepness and complex geometry of slope (different lines-of-sight can be used for the survey to avoid problems related to slope orientation). With favourable weather conditions photographs can be acquired with high precision without problems related to the stability of the frame. The high resolution of photographs help in the analysis of slopes.	In practice may be difficult to use Can reach a maximum elevation of 300 m. Not possible to use in case of inaccessibility of the area facing the object under study or unfavourable weather conditions. Expensive compared with the tripod method.
TLS	Easy and fast to use. If integrated with GPS and TLS it is possible perform a "non-static" survey which helps to overcome problems related to point density and slope orientation (different stations with different lines-of-sight can be used for the survey and IM to join all the data in a global reference system). In favourable conditions the output data is extremely precise and representative of the real geometry of the slope.	Not possible to use in case of inaccessibility of the area facing the object under study or unfavourable weather conditions. If the slope is very high occlusion may be present (Sturzenegger and Stead, 2009a). The use of a "static survey" lead to face problem related to point density and slope orientation (Ferrero et al., 2010; Lato et al., 2010)

826 Table 7. Close range DTP and TLS techniques: advantages and limitations (after Francioni, 2013).

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Analysis method	Critical input Parameters (Stead et al., 2006)	Improvement gained from DTP and TLS approach	Limitations
3D Discontinuum modelling (e.g., distinct element, DDA, Lattice-spring)	Slope and discontinuity geometry; intact rock constitutive criteria (elastic, elasto-plastic, etc.); discontinuity stiffness and shear strength; groundwater and in situ stress conditions.	Representative slope geometry is ensured even in case of complex geometry. Major discrete fractures can be identified as well as joint sets (only the rock slope face can be characterized but data for discrete fracture networks can be derived and used to understand the role of brittle fracture, rock bridges and persistence in the analyses; i.e. Sturzenegger 2009b; Tukey et al., 2013).	The data processing and interpretation time when dealing with accurate and detailed data are longer and the model construction is more complex. Only the rock slope face can be characterized. The proposed approach does not provide information about the most suitable constitutive criteria and the joint and rock mass characteristics; nevertheless, some authors (i.e. Fekete et al., 2010; Tatone and Grasselli, 2010) studied the importance of using a 3D approach with high-resolution surface measurements in the definition of these characteristics.

841 Table 8. Improvements remote sensing techniques offer in the definition of critical input parameters for  
842 3D Discontinuum Codes (after Stead et al., 2006 and Francioni, 2013).

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