## **Different glacier status with atmospheric circulations in Tibetan**

## **Plateau and surroundings**

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#### 1 **Supporting Online Material**

2 Glacial area in the Tibetan Plateau (TP) and surroundings is still an issue to be 3 precisely inventoried. Here we provide a figure of  $\sim$ 100, 000 km<sup>2</sup> in Table S1 based 4 on studies of Yao et al.<sup>1</sup> and Dyurgerov<sup>32</sup>.

5 There are many publications in Chinese journals about glacial fluctuations in 6  $\binom{1}{10}$  China<sup>33-50</sup>. Those studies showed different pictures of glacier fluctuations. Some 7 indicated heterogeneous retreat magnitudes in different areas, with the largest retreat 8 in the margin of the TP and decreasing to the interior  $33-43$ , while others showed a few 9 advancing glaciers in the southeastern TP and central Himalayas<sup>46-47</sup>. However, Jin et  $10$  al.<sup>47</sup> admitted, from their own experiences, that the 'advancing glaciers' might be 11 misled by satellite images with heavy snow and shadow. Two advancing glaciers interpreted from early satellite images in the paper of Liu et al.<sup>46</sup> are also retreating 13 glaciers according to more reliable satellite images. Some studies emphasized the 14 unique response of debris-covered glaciers to climate changes<sup>15,51, 52</sup>. A few studies of in situ observation of glacial length and mass balance do exist<sup>53,54</sup>, but are too sparse 16 to shed light on a holistic picture of glacial fluctuations of the region. Recently some 17 more data are published in the southern TP and Himalayas using new technologies 18 and methods<sup>52</sup>. The situation thus underlines the necessity for a comprehensive study 19 of glacial area based on satellite images and in situ glacial length and glacial mass 20 balance. This is the major purpose of our study in this paper.

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#### <sup>22</sup>**1. Glacier area reduction**

23 Glacier area change is studied using remote sensing images and topographic maps. 24 Among the glacial area analysis of 16 river basins in seven regions, nine are based on 25 literatures and seven are based on our studies (four had not been published). It's 26 important to evaluate its potential uncertainty. The accuracy of glacier delineation 27 based on remote sensing images depends on the resolution of the utilized image, the 28 conditions at the time of the acquisition (especially seasonal snow) and the knowledge 29 of the operator to decide the lake region, glacier in cast shadow, perennial or seasonal snow, debris-covered area and location of ice divides and might be subjective<sup>55</sup>. The 31 mapping uncertainty of our studies is less than 3% for clean-ice glaciers and 4% for 32 debris-covered glaciers in the Boshula Mountain Range, Southeastern Tibet, Koshi 33 River Basin Nepal, Mapam Yumco Basin, Geladandong and Shulenan Mountain. This 34 is similar with previous studies, which reported a mapping uncertainty of  $\pm 2-3\%$  for 35 clean-ice glacier<sup>56,57</sup>, of  $\pm$ 3-4% for debris-covered glacier for ASTER, Landsat TM by 36 comparison with other high-resolution images or estimation<sup>58</sup>. However, it's difficult 37 to evaluate the uncertainty of the result from previous studies (cited in this study), 38 since they were from different operators using different methods. To reduce large 39 uncertainty, we only use the results of the studies with similar resolution images 40 (Landsat or ASTER) and at the end of snow season. The uncertainty is no more than 41 5%.

42 There are totally 7,090 glaciers  $(\sim 13,363.5 \text{ km}^2)$  analyzed for glacial area 43 reduction in this study. We have measured 2,135 glaciers  $(\sim 3154.5 \text{ km}^2)$  based on 44 digitized glacier inventories from topographic maps and remote sensing images in six 45 river basins in the TP and surroundings. The other 4,955 glaciers ( $\sim$ 10209.0 km<sup>2</sup>) were 46 summarized from the results of previous studies<sup>59-69</sup>.

47 The locations of studied glaciers for area reduction are shown in Figure S1 and 48 Table S2. The total glacier area of 7,090 glaciers has decreased from 13,363.5  $\text{km}^2$  to 12,130.7 km<sup>2</sup> in the period between the 1970s and 2000s. The average decreasing rate 50 is  $\sim$ -0.30% a<sup>-1</sup> for the TP and surroundings, but different from region to region, with the largest  $(-0.57\% \text{ a}^{-1})$  in Region I and the smallest  $(-0.07\% \text{ a}^{-1})$  in Region V (Table 52 S3).

53 We have also compared retreat magnitude of different glacier sizes in different 54 regions. Figure S2 summarized the change of glacial area with different sizes in 55 different regions. According to the sizes, glaciers were classified into four types: 56  $\leq$ 1km<sup>2</sup>, 1.0-5.0 km<sup>2</sup>, 5.0-10.0 km<sup>2</sup> and >10km<sup>2</sup>. Although the change rate of smaller 57 glaciers is larger than that of larger glaciers in all the regions, the general pattern of 58 glacial retreat is controlled by regional climate. In a region with the same climate 59 dominance, such as monsoon or westerlies, glaciers undergo similar retreat tendency, 60 no matter what sizes the glaciers are.

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## <sup>62</sup>**2. Glacial length fluctuation**

63 We have analyzed glacial length fluctuation of 82 glaciers from the 1970s to 64 2000s. Among them, the length fluctuation of 13 glaciers is based on our *in situ* 65 observations and that of the other 69 is summarized from previous studies<sup>70-89</sup>. The 66 general pattern of glacial length fluctuation shows more intensive retreat in the 67 Himalayas, with largest retreat of three glaciers in the southeastern TP or the eastern 68 Himalayas and moderate retreat of eight glaciers in the central Himalayas and 20 69 glaciers in the western Himalayas, and decreasing retreat of five glaciers in the 70 interior northeastern TP and five glaciers in the Nyainqentanglha Mountain, and stable 71 or even advancing characteristics in the eastern Pamir regions and west Kunlun 72 Mountains (among the 41 glaciers, 17 retreated and 10 were stable, while 14 73 advanced).

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## <sup>75</sup>**3**.**Glacial mass balance change**

#### 76 **3.1 Current status of glacial mass balance**

77 We have analyzed 15 glaciers to assess the current mass balance status in the TP 78 and surroundings (Table S6-S7). Among them, 11 glaciers are based on our *in situ*  $\mu$  measurements and the other four are summarized from previous studies<sup>90-92</sup>. The 80 results of the 11 glaciers are presented in the following:

#### 81 **The Parlung No.10 Glacier**

82 The Parlung No.10 Glacier (29º17´N, 96º54´E) is a valley glacier (Figure S3a) 83 flowing from 4,900 to 5,625 m, with an area of 2.1 km<sup>2</sup> and a length of 3.5 km. A total 84 of 11 measuring stakes have been distributed on this glacier. Figure S3c shows the 85 significant variations of spatial distribution of net mass balance year by year, 86 accompanied with ELAs fluctuations (Figure S3b) (data for 2009/10 is absent due to 87 measuring stakes falling down). 2008/09 is the most negative mass balance year we 88 have measured, shown both by the spatial distribution of net balance (negative almost 89 on the whole glacier, shown in Figure S3c and Table S7) and ELA (Figure S3b). The 90 ELAs in the past four years have shifted from 5,419m to 5,500m, and nearly reached 91 the glacier summit in 2008/09.

#### 92 **The Parlung No.12 Glacier**

93 The Parlung No.12 Glacier (29º18´N, 96º54´E) is a small cirque glacier with an 94 area of about  $0.2 \text{ km}^2$  and a length of nearly 0.6 km (Figure S4a). Its elevation ranges 95 from 5,130 to 5,265 m. A total of 5 measuring stakes have been distributed on this 96 glacier. The spatial distribution of net balance demonstrates negative net balance on 97 the whole glacier every year (Figure S4c) and the ELAs have already risen beyond the 98 glacier summit during the period of measurement. This glacier is now suffering from 99 significant mass deficit.

#### 100 **The Parlung No.94 Glacier**

101 The Parlung No.94 Glacier (29º23´N, 96º59´E) is a valley glacier with an area of 102  $2.5 \text{ km}^2$ , a length of 2.9 km and its elevation range from 5,000 to 5,635 m (Figure 103 S5a). A total of 19 measuring stakes have been spatially distributed on this glacier. 104 With similar glacier area and altitudinal range, both the spatial distribution of net 105 balance and ELAs variations show a similar pattern with the Parlung No.10 Glacier 106 (Figures S5b-c). The most negative mass balance with the highest ELA occurs in 107 2008/09.

#### 108 **The Parlung No.390 Glacier**

109 The Parlung No.390 Glacier (29º21´N, 97º01´E) flows from an elevation of 5,160 to 5,460 m, with an area of about 0.5 km<sup>2</sup> and a length of 1.2 km (Figure S6a). A total 111 of 5 measuring stakes have been distributed along the central axis of this glacier. 112 Located in the same region with the Parlung No.10, 12 and 94 Glacier, the spatial 113 distribution of net balance and ELAs of the Parlung No.390 Glacier also show a 114 similar pattern (Figures S6b-c).

#### 115 **The Gurenhekou Glacier**

116 The Gurenhekou Glacier (30°11′N, 90° 28′E, area 1.4 km<sup>2</sup>, length 2.9 km) lies near 117 the town of Yangbajin on the southern slope of Nyainqentanglha Mountain. The 118 glacier ranges from 6,040 to 5,525 m in elevation (Figure S7a). A total of 12 119 measuring stakes have been spatially distributed on this glacier. Different from the 120 four glaciers in the southeastern TP, the circumstance of spatial distribution of 121 negative net balance on the whole glacier did not occur on the Gurenhekou Glacier. In 122 contrast, the mass balance is positive(Figure S7c) with descending ELA (Figure S7b) 123 in 2007/08.

#### 124 **The Zhongxi Glacier**

125 The Zhongxi Glacier (30°52′N, 91°27′E, area 1.6 km<sup>2</sup>, length 2.6 km) lies on the 126 northeastern Nyainqentanglha Mountain Range, near the Npen Co lake (Figure S8a). 127 The elevations of the glacier summit and terminus are 6,210 and 5,376 m, respectively. 128 A total of 16 measuring stakes have been spatially distributed on this glacier. The 129 three-year measurements of this glacier show a similar pattern of spatial distribution 130 of net balance and ELA with the Gurenhekou Glacier. Slight positive mass balance 131 occurred in 2007/08 (Figures S8b-c).

#### 132 **The Kangwure Glacier**

133 The Kangwure Glacier (28°28′N 85°49′E, area 1.9 km<sup>2</sup>, length 3.1 km) lies on the

134 northern slope of central Himalayas range, near the Xixiabangma Mountain (Figure 135 S9a). The glacier ranges from 6,060 to 5,690 m. A total of 16 measuring stakes have 136 been spatially distributed on this glacier. Due to lack of mass balance measurement in 137 2006, we take the average between 2005/06 and 2006/07 for 2005/06 and 2006/07, 138 respectively. The data in 2007/08 is absent due to measuring stakes falling down. 139 Located in the northern slope of Himalayas, the spatial distribution of net balance and 140 ELAs show intensive mass loss (Figure S9b, 9c) although the average elevation of 141 this glacier is very high. The spatial distribution of net balance is negative on the 142 whole glaciers in 2005/07.

#### 143 **The Naimona'nyi Glacier**

144 The Naimona'nyi Glacier (30º27´N, 81º20´E) is located in the northern slope of 145 west Himalayas and is a valley glacier with an area of 7.8  $km^2$  and a length of 7.7 km 146 (Figure S10a). Its altitudinal range is between 5,465 and 7,520 m. 16 measuring stakes 147 have been spatially distributed on this glacier. Due to lack of measurement in 2006, 148 we take the average between 2005/06 and 2006/07 for 2005/06 and 2006/07, 149 respectively. This glacier suffers from significant mass deficit, with particular 150 circumstance of negative net balance on the whole glacier between 2005/06 and 151 2006/07 (Figure S10c). An earlier study indicated that the glacier deficit status may 152 have lasted for a long time<sup>93</sup>.

#### 153 **The Muztag Ata Glacier**

154 The Muztag Ata Glacier (38º14´N, 75º03´E) is located in West Kunlun Mountain. 155 This glacier has an area of  $0.96 \text{ km}^2$ , with a length of 1.8 km (Figure S11a). Its 156 altitudinal range is between 5,235 and 5,940 m. A total of 13 measuring stakes have 157 been spatially distributed on this glacier. Due to lack of measurement in 2009, we take 158 the average between 2008/09 and 2009/10 for 2008/09 and 2009/10, respectively. The 159 spatial distributions of net balance are positive almost on the whole glacier during the 160 period, except in 2002/03 (Figure S11c).

#### 161 **The Xiaodongkemadi Glacier**

162 The Xiaodongkemadi Glacier (33°04´N, 92°05´E) is located at the headwaters of 163 the Dongkemadi river, a tributary at the upper reaches of the Buqu River near the 164 Tanggula pass. The Glacier is 1.8 km<sup>2</sup> in area, with a length of 2.8 km (Figure S12a). 165 The elevations of the summit and terminus of the glacier are 5,926 and 5,380 m, 166 respectively. A total of 25 measuring stakes were set up on the glacier. The most 167 significant phenomena of the Xiaodongkemadi Glacier is negative net balance in

168 2009/10 (-1,066 mm), which is higher than that in 2007/08 (-80 mm) and 2008/09

169 (-91mm) by one order (Figures S12b-c).

#### 170 **The Qiyi Glacier**

171 The Qiyi Glacier (39°15´N, 97°45´E) is located on the northern slope of Qilian Mts. 172 Its area is 2.87 km<sup>2</sup>, with a length of 3.8 km (Figure S13a). Its altitudinal range is 173 between 4,304 and 5,159 m. A total of 28 measuring stakes were set up on the glacier. 174 Similar with the Xiaodongkemadi Glacier, the mass balance is much more negative 175 (-648 mm) in 2009/10, comparing with that in 2007/08 (-105 mm) and 2008/09 (-74 176 mm).

177 Generally, all the 11 glaciers show negative mass balance in the five years of 178 measurement, except the Muztag Ata Glacier that shows positive mass balance in the 179 past four years (Table S7). The most negative mass balance is observed in the 180 southeastern TP, while positive mass balance is observed in the eastern Pamir regions.

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#### 182 **3.2. Long-time series of mass balance**

183 There are three glaciers with mass balance starting from the early 1990s or earlier. 184 The Xiaodongkemadi Glacier has been continuously measured for mass balance 185 starting from 1988/89. Otherwise, the Qiyi and Kangwure Glaciers have ten-year and 186 two-year, respectively, *in situ* measurements before 2005. We have reconstructed the 187 past mass balance using the relationship between the measured mass balance and 188 meteorological data (Table S8). For the Qiyi Glacier, the *in situ* mass balance 189 measurement first started in 1974/75 and measured discontinuously since. Wang et al  $(2010)^{94}$  established a statistical model between ELAs and meteorological factors 191 (warm air temperature and cold-season precipitation) on the basis of measurement of 192 the ELAs. The reconstructed ELAs agreed well with the measured data. In this study, 193 based on this model, we have reconstructed annual mass balances  $(B_n)$  in the Qiyi 194 Glacier ( $B_n = 11456 - 2.44 \times ELA$ , n=12,  $R^2 = 0.92$ ). Figure S14 shows the comparison 195 between the simulated annual mass balances with *in situ* measurements. The 196 simulated mass balance agrees well with the measurements.

197 For the Kangwure Glacier, we have used mass balance data of 1991/92 and 198 1992/93<sup>95</sup>, and that between 2005/06 and 2009/10. We used the meteorological data at 199 Dingri meteorological station (about 100 km away from the Kangwure Glacier) to 200 analyze the relationship between mass balance and meteorological factors. A 201 regression model was used to reconstruct the mass balance since 1991/92 based on the 202 meteorological factors (annual air temperature (T) and precipitation (P)) and 5-year *in*  203 *situ* mass balance observations (the linear regression is  $B_n = -350 \times T + 2.36 \times P - 22.4$ , 204  $n=5$ ,  $R^2=0.61$ ) (Figure S15). The reconstructed mass balance data are shown in Table 205 S8.

#### 206 **3.3 Climate controls over glacial mass balance**

207 We've studied glacial mass balance of 11 glaciers on the TP. Mass balance 208 measurement of glaciers with different sizes under different climate regimes 209 demonstrates heterogeneous mass loss as shown in Figure S16. In the monsoonal 210 region where precipitation is decreasing, glacial mass balance is the most negative. In 211 contrast, in the westerly region where precipitation is increasing, glacial mass balance 212 is positive. Glacial mass balance is moderate in the transitional region.

## <sup>213</sup>**4. Contrast precipitation trends between Himalayas and**

## <sup>214</sup>**Pamir regions**

215 It was first proposed by Yao et al.<sup>96</sup>, and Liu and Chen<sup>23</sup> that the amplitude of 216 temperature change increases with elevations. Latest study by Frauenfeld et al.<sup>97</sup>, Oin 217 et al.<sup>24</sup>, Kang et al.<sup>54</sup> and You et al.<sup>98</sup> found that, the TP is warming, and the warming 218 rate increases with elevation before becoming quite stable with a slight decline near 219 the highest elevations. However, by examining Figure 4 in Qin et al.<sup>24</sup>, we found that 220 the warming rate is most intensive between 4,800 and 6,200 m a.s.l, which covers the 221 ablation area of almost all glaciers in the TP. This is still not conclusive, and more 222 studies are necessary to narrow down the uncertainties before a definite conclusion is 223 drawn.

224 Precipitation is a very important factor contributing to the glacial mass balance 225 change. This is particularly the case in the TP and surroundings. The most intensive 226 glacial shrinkage in the Himalayas coincides with the decreasing precipitation 227 accompanied by the weakening Indian monsoon and the least intensive glacial 228 shrinkage in the eastern Pamir regions is linking with the increasing precipitation 229 accompanied by the strengthening westerlies (Figure 4 in the Text). The trends of 230 decreasing precipitation in the Himalayas and increasing precipitation in the eastern 231 Pamir regions are further supported by the Global Precipitation Climatology Project 232 (GPCP) dataset<sup>99</sup> for the period 1979–2010. The precipitation series in the four grids

233 numbered as 1-4 in Figure 4a in the eastern Pamir regions show statistically 234 significant increasing trends at the 99% confidence level using the Mann-Kendall test 235 (Figure S17). The precipitation series numbered as 5-8 and 14-16 in Figure 4a in the 236 Himalayas exhibit statistically significant decreasing trends at the 99% confidence 237 level (Figure S18). The series numbered as 9-13 are not as significant as those of 5-8 238 and 14-16, but still show obvious decreasing trends.

239 Figure S19 shows the seasonality of the precipitation at Linzhi, Bomi and Zayu 240 stations in the Indian monsoon-dominated Himalayas and at Taxkorgen station in the 241 westerlies-dominated Pamir regions, indicating strong seasonality characterized by 242 high precipitation in the summer and low precipitation in the winter. Figure S20 243 shows the spatial features of GPCP seasonal precipitation trend in the summer and 244 winter during 1979-2010. The results show a similar spatial pattern of the seasonal 245 precipitation to that of annual precipitation (Figure 4a) in the text, demonstrating 246 decreasing precipitation in the eastern Himalayas and increasing precipitation in the 247 eastern Pamir regions. The decreasing trend is more intensive in the summer in the 248 Indian monsoon-dominated Himalayas (Figure S20a), confirming the weakening 249 Indian monsoon; while the increasing trend is more intensive in the winter in the 250 westerlies-dominated Pamir regions (Figure S20b), confirming the strengthening 251 westerlies.

## 252 **Supplementary Tables**

**Supplementary Table S1**. Distribution of Glaciers in the TP and surroundings calculated from Yao et al<sup>1</sup> 254 and Dyurgerov<sup>32</sup>





### 255 **Supplementary Table S2**.Data and method for analyzing glacial area reduction in each basin

256 Note: \* represent our studies.





Note:\* represent our studies. Regional average is calculated by  $\overline{RA} = \sum_{i=1}^{n}$  $=$   $\lambda$   $v $\times$$ *n* 259 Note:\* represent our studies. Regional average is calculated by  $RA = \sum_{i=1}^{n} P \times (\frac{T}{R})$ ; RA is regional weighted average; P is the percentage of annual area change on each basin; T is number

260 of glaciers on each basin in 1970s, R is number of regional glaciers in 1970s; *n* is the number of basins on each region.

















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Regions	Glaciers	2002/03	2003/04	2004/05	2005/06	2006/07	2007/08	2008/09	2009/10	Average	Regional average
	Parlung No.10		$\blacksquare$		$-675$	$-283$	$-593$	$-1575$	$\blacksquare$	$-781$	$-1105$
	Parlung No.12	$\blacksquare$	$\blacksquare$		$-1449$	$-1112$	$-1410$	$-2476$	$-2046$	$-1698$	
	Parlung No.94	$\overline{\phantom{a}}$	$\overline{\phantom{a}}$	$\overline{\phantom{a}}$	$-913$	$-254$	$-1079$	$-2018$	$-347$	$-922$	
	Parlung No.390	$\overline{\phantom{a}}$	$\blacksquare$		$\overline{\phantom{a}}$	$-170$	$-1250$	$-1673$	$-982$	$-1019$	
$\mathbf{  }$	Gurenhekou	$\overline{\phantom{a}}$	$\overline{\phantom{a}}$	$\overline{\phantom{0}}$	$-319$	$-196$	497	$-839$	$-703$	$-312$	$-418$
	Zhongxi	$\blacksquare$	$\blacksquare$			$\blacksquare$	264	$-1045$	-789	$-523$	
III	Kangwure				$-1023$	$-392$	$-487$	$-1092$	$-300$	-660	$-757$
	$AX010^{90*}$				$-810$				$\sim$	$-810$	
	$Yala^{90*}$				$-800$				$\blacksquare$	$-800$	
IV	Naimona'nyi				$-658$		$-718$	$-472$	$-276$	$-556$	
	Hamtah <sup>91</sup>		$-1857$	$-1856$	$-790$	$\blacksquare$			$\blacksquare$	$-1501$	$-908$
	Chhota Shigri <sup>92</sup>	$-1400$	$-1227$	144	$-1400$	$-980$	$-930$	130	330	$-667$	
V	Muztag Ata	$\blacksquare$	$\blacksquare$	$\blacksquare$	$-237$	956	79	220		248	248
VI	Xiaodongkemadi	$\blacksquare$			$-917$	$-591$	$-80$	$-91$	$-1066$	$-549$	$-549$
VII	Qiyi				$-955$	$-513$	$-105$	-74	$-648$	$-459$	$-459$

267 **Supplementary Table S6** Recent annual mass balances in different regions in the TP.

<sup>\*</sup>The average mass balances AX010 (1999-2008) and Yala Glacier (1996-2009) were taken to represent the recent mass balance since 2002/03. The red color numbers for the Kangwure Glacier were reconstructed following the met were reconstructed following the method in Section 3.2.

- 270 **Supplementary Table S7** Mass balance of Long-time series for the Qiyi, Xiaodongkemadi and
- 271 Kangwure Glaciers in the TP.



272 \* The red color numbers are the reconstructed mass balance.

# **Supplementary Figures**



276 studies in the seven regions.



 $\substack{277 \\ 278}$ **Supplementary Figure S2** Comparison of the change of glacier with different sizes in different regions. The absolute area changes (km<sup>2</sup>) are shown on top of each column. The number in the regions. The absolute area changes  $(km<sup>2</sup>)$  are shown on top of each column. The number in the 280 parentheses is the percentage of annual area change of this region. 281



285 **Supplementary Figure S3** The Parlung No. 10 Glacier. a) Photo of the Parlung No.10 Glacier, 286 b) variation of ELAs and c) spatial mass balance distribution in past four balance years



290 **Supplementary Figure S4** The Parlung No. 12 Glacier. a) Photo of the Parlung No. 12 Glacier, c) spatial mass balance distribution in past five balance years

c) spatial mass balance distribution in past five balance years



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295 **Supplementary Figure S5** The Parlung No. 94 Glacier. a) Photo of the Parlung No.94 Glacier, b)

296 variation of ELAs and c) spatial mass balance distribution in past five balance years





**Supplementary Figure S6** The Parlung No. 390 Glacier. a) Photo of the Parlung No.390 Glacier, b)

299 variation of ELAs and c) spatial mass balance distribution in past four balance years



**Supplementary Figure S7** The Gurenhekou Glacier. a) Photo of the Gurenhekou Glacier, b) 302 variation of ELAs and c) spatial mass balance distribution in past six balance years





**Supplementary Figure S8** The Zhongxi Glacier. a) Photo of the Zhongxi Glacier, b) variation of

305 ELAs and c) spatial mass balance distribution in past three balance years





**Supplementary Figure S9** The Kangwure Glacier. a) Photo of the Kangwure Glacier, b) variation

308 of ELAs and c) spatial mass balance distribution in past three balance years



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**Supplementary Figure S10** The Naimona'nyi Glacier. a) Photo of the Naimona'nyi Glacier, b)

313 variation of ELAs and c) spatial mass balance distribution



**Supplementary Figure S11** The Muztag Ata Glacier. a) Photo of the Muztag Ata Glacier, b)

#### 317 variation of ELAs and c) spatial mass balance distribution



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319 **Supplementary Figure S12** The Xiaodongkemadi Glacier. a) Photo of the Xiaodongkemadi

320 Glacier, b) variation of ELAs and c) spatial mass balance distribution in past four balance years



321



323 and c) spatial mass balance distribution in past four balance years



325 Supplementary Figure S14 Comparison between measured and simulated mass balance for the

327 Qiyi Glacier.



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- 330 **Supplementary Figure S15** Comparison between measured and simulated mass balances for the
- 331 Kangwure Glacier.
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334 **Supplementary Figure S16** Distinctive features of glacial mass balance under different climate 335 dominances. (No. 10, 12, 94, 390-Parlung No. 10, 12, 94, 390; NN-Naimona'nyi; QY-Qiyi; KW-

336 Kangwure; XD-Xiaodongkemadi; ZX-Zhongxi; GH-Gurenhekou; MA-Muztag Ata)







**Supplementary Figure S17** The linear trends of GPCP precipitation (mm day<sup>-1</sup>) for 1979–2010 341 from the grids numbered as 1-4 in Figure S4g. All the positive trends are statistically significant at 342 the 99% confidence level using the Mann-Kendall test.



348 statistically significant at the 99% confidence level using the Mann-Kendall test.



 $\frac{350}{351}$ **Supplementary Figure S19** The monthly mean precipitation at the meteorological stations of 352 Linzhi, Zayu, Bomi and Taxkorgan during 1970-2009. The precipitation at all the four stations 353 shows a strong seasonality characterized by high precipitation in the summer and low precipitation 354 in the winter.



**Supplementary Figure S20** The spatial feature of GPCP precipitation (mm day<sup>-1</sup>) trend in the 358 summer (a) and winter (b) seasons in the TP and surroundings during 1979-2010. Similar to the 359 annual spatial trends in Figure 4a in the text, the summer and winter spatial trends demonstrate 360 decreasing precipitation in the eastern Himalayas and increasing precipitation in the eastern Pamir 361 regions, with more intensive decreasing in the summer in the Indian monsoon-dominated 362 Himalayas and more intensive increasing in the winter in the westerlies-dominated Pamir regions.

363 364

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