

1 Supplementary Information on

2 **Flip flop of Day-night and Summer-winter Surface Urban Heat Island**  
3 **Intensity in India**

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## 16 **Supplementary Text**

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### 18 **S1. MODIS data and their accuracy**

19 Here we use three data sets from Moderate Resolution Imaging Spectroradiometer (MODIS)  
20 satellite data product: (i) Land Cover (LC) information (ii) Land Surface Temperature (LST)  
21 (iii) vegetation condition (NDVI) and (iv) evapotranspiration (ET). Identification of the urban  
22 clusters is based on the MODIS LC type product (MCD12Q1, 500 m, annual) of the year  
23 2008. Here we use the LC classification types from International Geosphere Biosphere  
24 Programme (IGBP)<sup>S1-S2</sup> with 17 LC classes.

25 We estimate the SUHII with MODIS-Aqua LST dataset (MYD11A2, 1000 m, Version 5) at  
26 eight-day interval. This data product is validated over a widely distributed set of locations  
27 and time periods via several ground-truth and validation efforts<sup>S3-S5</sup> and is frequently used for  
28 surface UHI analysis<sup>5-8</sup>. The MODIS LST data is derived from two thermal infrared band  
29 channels, 31 (10.78–11.28  $\mu\text{m}$ ) and 32 (11.77–12.27  $\mu\text{m}$ ), using the split-window  
30 algorithm<sup>S6</sup>. This algorithm corrects for atmospheric effects and emissivity using a look-up  
31 table, based on global land surface emissivity in the thermal infrared<sup>S7</sup>. The dataset is  
32 comprised of daytime (~13:30) and night time (~01:30) LSTs, quality control (QC),  
33 observation times, view angles, bits of clear sky days and nights, and emissivity estimates.  
34 The QC values provide very important information for filtering of low-quality pixels due to  
35 clouds or other processing failures. For the current study the QC Scientific Data Set (SDS)  
36 for LST are extracted by reading the bits in the 8-bit unsigned integer<sup>S8</sup>. The data pixels  
37 where, the error in the computed LST is less than  $3^0\text{K}$ , are considered for this analysis.

38 Traditionally, the in situ surface air temperature, over the urban and rural regions, is used for  
39 the estimation of UHI intensity. Satellite measured land surface temperature are used in the studies,  
40 where the objective is to understand the space-time variability of the characteristics of UHI, and this is  
41 primarily due to the uniform availability, as well as the global coverage of the satellite data<sup>5</sup>. The

42 satellite measured LST is reported to have a good correlation and lower bias with respect to both the  
43 in situ LST and surface air-temperature<sup>5,8,14</sup>. The MODIS LST data has been applied to characterize  
44 the UHI for the global<sup>5</sup> as well as Asian mega cities<sup>S9-S10,12</sup>. It is observed<sup>S11</sup> that the UHI intensity  
45 computed with MODIS LST for the city of Delhi has similar characteristics to the same derived with  
46 surface air temperature. The study found that the nighttime UHI intensity as well as the UHI hotspots  
47 are well captured from the satellite data and hence can be used as a proxy to the in situ observations  
48 for the estimation of the UHI characteristics. The vegetation indices are obtained from MODIS-  
49 Aqua product (MYD13A, 1000 m) at temporal resolution of 16 days. The QC information  
50 available with the data product is utilised to filter out good quality data to be utilised for the  
51 analysis.

52 We further obtain the evapotranspiration (ET) data from MODIS-Aqua product (MYD13A,  
53 1000 m) at a temporal resolution of 16 days. The available information about the quality  
54 control is used to filter out good quality data to be utilised for this analysis.

55 All three MODIS data products for the present study are obtained from climate data archive  
56 at from the Earth Observing System Data and Information System (EOSDIS) service tool  
57 reverb: <http://reverb.echo.nasa.gov>. The three MODIS datasets namely LST NDVI and ET  
58 datasets were selected during the period between 2003 to 2013 pre-monsoon summer season  
59 (March-May) and winter season (December-February).

60

## 61 **S2. Selection of Urban and Non-urban regions**

62 A cell resizing operation<sup>S12</sup> popularly known as resampling is carried out over the acquired  
63 LC data. The resampling methodology operates a pixel scale calculation to change spatial  
64 resolution of the dataset without losing useful information available with the data. The 500 m  
65 spatial resolution LC images are interpolated to 1000 m resolution dataset to match with the  
66 available LST data. There are many resampling methods available, through a variety of  
67 platforms; here we apply the widely used Nearest Neighbor (NN)<sup>S13</sup> resampling technique.  
68 The NN algorithm maintains the original brightness values of MODIS pixels as well as has a  
69 lower processing time of spatial resampling<sup>S14</sup>.

70 The city clustering algorithm<sup>35</sup> is used to determine the extent of urban area for all the 84  
71 cities. The following steps are followed to estimate UHI intensities over the selected cities:

72 1. An urban map is prepared based on the geographical location of centre of each big city.  
73 The urban map is used to identify first urban node (root node here after) from the resampled  
74 MODIS LC data. The root nodes are queued and processed one by one to identify their  
75 respective urban and non-urban region. A root node is selected from the queue.

76 2. We examine eight neighbouring nodes around the root node. If the land cover type of the  
77 neighbouring nodes has urban land cover, the neighbouring nodes are added into a queue,  
78 assigning attribute of the neighbouring nodes as an urban node, otherwise the attribute of the  
79 neighbouring node is assigned as a non-urban node.

80 3. Step 2 is repeated for each node in the queue till the queue is empty. When the queue is  
81 empty; we quit the search and return the urban map to grow the cluster of the successive root  
82 node.

83 4. After the urban map is returned, suburban area is defined as the buffer zone, which is a  
84 ring around urban area that consists of the nonurban nodes excluding water.

85 5. The urban centres with the count of buffer zone grids covering the 50-150% of urban  
86 nodes are selected for the SUHII analysis.

87 Supplementary Figure S1 shows an example of identified urban cluster by CCA and selection  
88 of corresponding non-urban region along the urban boundary. The figure S1 presents four  
89 different urban clusters namely New Delhi, Hyderabad, Ahmedabad and Varanasi (Figures  
90 S1a, b, c and d respectively) having varied size and they are from different geographical  
91 location in the country. The identified urban cluster is denoted with the black boundary. The  
92 selected non-urban region surrounding the urban centre is presented with the red periphery of  
93 1km radius around the urban boundary. Here, we note that the selected non-urban area does  
94 not include sub-urban belts which have built up areas as dominant land use. The CCA  
95 includes peripheral sub-urban region as a part of the corresponding urban centre as they have  
96 urban built up as the dominant LULC. The surrounding non-urban region largely comprises  
97 of croplands (figure 2i). The mean LST over an urban (buffer) region is calculated  
98 as the spatial average of the LST observed over grid point identified with the  
99 CCA for that particular urban (buffer) region.

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### 101 **S3: Details of monthly UHI analysis**

102 The MODIS data product used here provides four evenly distributed data points  
103 for each month. We obtain 12 data points over each season. This ensures that  
104 each month of the season is well represented with sufficient and even sampling  
105 in the present UHI analysis. However, the representation month of May in the  
106 pre-monsoon summer season day time SUHII estimation remains poor. The SUHII  
107 for 53 urban centres located mainly in the central India could not be  
108 evaluated due to data quality and availability problem. The result obtained  
109 from the monthly UHI analysis is presented in figure S3. We compute the SUHII

110 individually for both the pre-monsoon summer months (March-May), daytime (figure S3, b-  
111 d), nighttime (figure S3, f-h), winter daytime (December-February; figure S3, h-j) and winter  
112 nighttime (figure S3, l-m). We observe similar characteristics of UHI across the months in a  
113 season with similar diurnal variation. The reported negative seasonal UHI over pre-  
114 monsoon summer months (figure S3a) is observed over March and April  
115 separately (figure S3, b-d). However the representation month of May in the pre-  
116 monsoon summer season remains poor due to data quality and availability  
117 problem (figure S3d). The observed positive pre-monsoon summer night time  
118 SUHII (figure S3e) is also prevalent with the same magnitude during the three  
119 individual months of the season. The statistical significance of diurnal difference  
120 is less than 0.05 only for the seasonal SUHII and not observed for individual  
121 months. This probably attributes to the sample size. The seasonal SUHII is positive  
122 during the winter days in the urban locations of the Gangetic basin, and negative in west  
123 central and southern India (Figure S3, i); the same is observed for all the three individual  
124 months of the season (figure S3, j-l). The *night* time positive but non-significant winter  
125 SUHII (figure S3, m) remains positive in all locations except one but is statistically  
126 significant over the individual months of the season for a majority of locations. These  
127 observations of the monthly UHI characteristics during both the seasons reconfirms that the  
128 estimated average seasonal SUHII values remains unchanged when the time windows fall in the  
129 different days.

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#### 131 **S4: Details of BC Emission Inventory dataset**

132 A multi-pollutant emission inventory was developed for India for 1996–2015  
133 including emissions from industry, transport, residential, agriculture and informal industry  
134 sector including fuel consumption, process and fugitive emissions and solvent use<sup>S15-S16</sup>.

135 These five sectors were disaggregated further into 13 source categories and ~75  
136 technologies/activities for estimating 2010 emissions. Black carbon emissions arise largely  
137 from traditional biomass technologies, characterized by inefficient combustion and  
138 significant emissions, are widely used in residential cooking and “informal industries”  
139 including brick production, food and agricultural product processing operations like drying  
140 and cooking operations related to sugarcane juice, milk, food-grain, jute, silk, tea and coffee.  
141 In addition, seasonal agricultural residue burning in field is a discontinuous source of  
142 significant emissions. The dataset estimated magnitudes of BC emissions for 2010 of 1.1  
143 Tgy-1, with average emissions in winter (December-February) of 108 Ggmon-1 and in pre-  
144 monsoon summer (March- May) of 88 Ggmon-1. While BC emissions from industry and  
145 transport are assumed invariant with season, differences in the two seasons arise largely from  
146 greater use of biomass fuels for residential heating activities in winter DJF (22 Ggmon-1), but  
147 greater agricultural residue burning in pre-monsoon MAM (15 Ggmon-1). Spatial distribution  
148 from on-road diesel transport uses GIS based shaped files with road densities of national  
149 super-highway and highway networks, state highways and city level grids, while that from  
150 light industry and gasoline transport follows urban population density, with large point  
151 sources assigned to specific location. Spatial and temporal (monthly) distribution of BC  
152 emissions from the residential sector is based on district-level values of user population of six  
153 different cooking fuels, while that from agricultural residue burning is based on state-level  
154 crop production, and climatological MODIS active fires<sup>S16</sup>. From the dataset, emission  
155 densities of BC in northern India ranging 60-120 Tggrid-1mon-1, in winter months, are  
156 larger than those of 20-100 Tggrid-1mon-1, in pre-monsoon summer months.

#### 157 **S4: Role of BC in SUHII modulation**

158 We found that the opposite seasonal patterns of SUHII between pre-monsoon summer and  
159 winter exist in north and central India (Figure 1) with positive winter daytime SUHII, which

160 is in contrast to negative pre-monsoon summer daytime SUHII. Daytime land-surface  
161 temperature could also be influenced by the atmospheric abundance of radiation absorbing  
162 constituents, such as black carbon (BC) aerosols, which are pollution particles emitted from  
163 incomplete combustion of fuels that strongly absorb radiation over the entire solar  
164 spectrum<sup>30</sup>. Radiation absorption from BC can lead to heating of the atmospheric layers in  
165 places where these particles are abundant at instantaneous rates of up to several degrees  
166 Kelvin per day<sup>S17</sup>. However, it should be noted that BC results into increased air temperature,  
167 but decreased LST and hence the direct impacts of BC do not explain the winter day-time  
168 SUHII behavior. On the other hand, the lowered LST possibly results into decreased ET with  
169 a modification in the latent heat flux that results into a feedback to LST. This is a complex  
170 process and needs model driven studies to understand the same.

171 There is a strong haze from aerosol pollution over north and central India<sup>S18</sup>, particularly in  
172 winter months. Such aerosol, which includes a significant fraction of BC, can lead to surface  
173 radiation balance changes that can affect land surface temperature and consequently, SUHII.  
174 To examine the possible role of BC aerosols, the spatial plot of SUHII was overlaid on BC  
175 emission fluxes, spatially distributed on a 25 km grid that was calculated in a recent  
176 emissions inventory for India<sup>S15-S16</sup> [Details in Supplementary Information S3]. The BC  
177 spatial distribution in winter revealed larger emissions in the Indo-Gangetic plain and central  
178 India, as well as some clusters in the west coast (Ahmedabad-Mumbai belt), east coast and  
179 south India, corresponding to the density of users of biomass fuels. On average, BC  
180 emissions and emission density are larger in northern India in the winter months (DJF) than  
181 in the pre-monsoon summer months (MAM). A BC-mediated mechanism of radiation flux  
182 change would only manifest during daytime, from solar radiation absorption by BC. This  
183 change in SUHII indicates urban land-surface temperatures exceeding those in adjoining  
184 areas in the winter months, co-located with increases in BC emission density, while falling



185 behind adjoining areas in pre-monsoon summer months. Further, at approximately 10 more  
186 sites in northwest and central India, negative SUHII values are reduced (smaller blue circles  
187 as in Figure S6a) in winter compared to SUHII values in pre-monsoon summer, once again  
188 indicating that land-surface temperatures in urban areas with higher BC emissions increase  
189 compared to temperatures in adjoining areas in the winter months. A correlation analysis of  
190 point SUHII with BC emission density ( $\text{Ton grid}^{-1} \text{ mon}^{-1}$ ) however, failed to show a  
191 significant association (Supplementary Figure S6c,d). This is not unexpected because  
192 changes in the radiation balance would be linked to the atmospheric concentration of BC in  
193 the surface layer and its vertical distribution rather than to emissions. In winter months, the  
194 prevailing meteorology in north India leads to low mixed layer heights and poor ventilation  
195 rates<sup>S19</sup>, which concentrate pollutants close to the surface. Thus, emission distributions, such  
196 as those analyzed in this work, reveal a possible influence of BC emissions on SUHII.  
197 Further analysis is needed, utilizing such emissions in regional climate models to calculate  
198 columnar concentrations of BC and subsequent radiative effects with modifications in ET,  
199 which could influence the SUHII characteristics.

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Sr. No.	Urban centre	Conclusion
1	Bangalore <sup>S20</sup>	The Land Surface Temperatures(LST) is computed from Landsat Thematic Mapper (TM), Enhanced Thematic Mapper (ETM) and MODIS LST data for different land use classes of the city. The estimated LST for years 1992, 2000 and 2007 reveals that increased urbanisation has resulted in higher LST due to high level of anthropogenic activities. LST-NDVI relationship is investigated and reported to have a negative correlation.
2	Bangalore <sup>S21</sup>	The UHI is evaluated with the help of Landsat ETM Plus data. LST-land cover relationship indicates that the city core has a significantly lower mean temperature than the outgrowth zones, mainly due to the presence of water bodies and vegetation. The urban core temperature is observed to vary from 1-7°C within different land cover classes.
3	Chennai <sup>S22</sup>	The LST difference between the urban and the surrounding area is assessed with the help of Landsat ETM Plus data for the year 2000. The results indicate that, the LST values hold a positive correlation with dense built up and negative correlation with the vegetation cover.
4	Chennai <sup>S23</sup>	The UHI is assessed through mobile recordings of urban air temperatures covering the major areas of the city for months May 2008 and January 2009. The results indicate UHI with an increasing air temperature in radial fashion from the suburbs towards the city centre. The mean UHI intensity is higher during winter than summer.
5	Delhi and Mumbai <sup>S24</sup>	The UHI is assessed using the Landsat TM image of 5 May 2010 for Delhi and the 17 April 2010 for Mumbai. The UHI is observed to be less prominent in Delhi mainly due to the mixed land use, substantial tree cover along roads; the ridge forests and river cutting across the city. Mumbai, on the other hand, processes a stronger UHI where the heat is trapped within, the built-up zones, whereas the creeks, sea and the lakes act as heat sinks. The stronger negative LST-NDVI correlation is observed over Mumbai than Delhi.
6	Delhi <sup>S25</sup>	Surface meteorological observations were taken using multisite ground based mini weather stations and meteorological towers for 25-28 May 2008. The

		UHI intensity is found to be higher both in the night and afternoon hours, dominating in areas of dense built up with intense human activity.
7	Delhi <sup>S11</sup>	The UHI is assessed with a network of micrometeorological observational stations across the city. Dense commercial areas were observed to have highest UHI. The UHI based on in situ ambient temperature is compared with the satellite-derived LST, a reasonable comparison is observed between this datasets during night time. However, the relation was found to be poor during daytime. The impact of land cover was also reflected with the built-up canopies reported largest gradient between air and skin temperature.
8	Delhi <sup>S26</sup>	The UHI is derived from Landsat data over the years 2000-2014. The study reveals that the UHI intensity is high during summer season compared to monsoon and winter seasons. The formation of heat island is controlled by vegetation density. UHI increases during night time.
9	Delhi <sup>S27</sup>	The monthly day and night time UHI is observed with the help of MODIS satellite data during the period 2007-2010. A prominent night time UHI over central parts of the city is reported almost throughout the year. However, during the day time an urban cool island is observed for the months of May, June, November and December. The study reports strong negative correlation between UHI intensity and mean monthly aerosol optical depth indicating a significant role played by aerosols in governing the urban thermal structure.
10	Hyderabad <sup>S28</sup>	The growth and LST of urban built up areas is observed with Landsat 5 TM data of the years 1989, 1999 and 2009. The study shows a 67% loss in water bodies while the urban built up areas have grown by 270%. The temperature profile shows a dip in temperature while encountering water bodies, gardens and parks.
11	Hyderabad <sup>S29</sup>	The day and night UHI is observed with the help of AATSR satellite data; field campaigns are conducted in synchronous with the satellite over pass for validation purpose. The satellite derived surface temperature values are found to be within 1 <sup>0</sup> C from ground measured values. The night time heat island formation observed with core urban regions showing high temperatures.
12	Kochi <sup>S30</sup>	The UHI intensity is quantified with mobile surveys carried out from January 2011 to March 2013. The pre-dawn UHI is observed to be more intense than early night UHI, also intensity in winter is stronger

		than in summer. The study area is classified into different local climate zones (LCZ), thermal gradient and cooling rates are observed within the zones and validated with the LCZ classification. The maximum UHI intensity is seen in the central part of the city.
13	Nagpur <sup>S31</sup>	Traverse surveys were carried during the summer and winter seasons, for the years 2012-2014, to measure night time mean canopy UHI intensity. Canopy UHI effects were found to be most prevailing in high building and population density areas. The negative impact of vegetation and positive effect of population density is revealed.
14	Pune <sup>S32</sup>	Dry and wet bulb temperature data obtained by a mobile survey conducted in April 1997. The results indicate that at night, the core of the city appears as both heat and moisture islands whereas at the time of sunrise as heat and dry islands. Situated in a basin-like topography, the city experiences stronger influence of winds rather than the thermal circulation systems arising from spatial in homogeneity in thermal and moisture patterns.
15	Thiruvananthapuram <sup>S33</sup>	Air temperature variations across the urban centre were recorded by mobile traverse method on June 29, 2010. Cooling and warming rates in the urban centre and suburban area were derived from stationary air temperature recorders. The study observes significant difference in the urban and rural cooling rates with the UHI intensity reported as 2.4°C.
16	Bopal <sup>S34</sup>	The study estimates UHI with the help of Landsat TM data of the year 2006. The study reports a prominent UHI effect over the city. The UHI intensity is observed to be higher over the roads and industrial zones and lower over the vegetated areas as well as the residential regions with lighter roofing.
17	Ahmedabad <sup>S35</sup>	In UHI effect is studied using Landsat ETM satellite data along with field measurements using Infra Red Gun in various zones of the city. The surface temperature near industrial areas and dense urban areas is reported to be higher as compared to other suburban areas in the city.
18	Mumbai <sup>S36</sup>	The UHI is estimated for the time period 1976-2007 of based on the meteorologically observed surface air temperature over the two urban stations of the urban region and two peripheral non urban stations. The study reveals prominent UHI during winter season for both day and night time than the summer season.

19	Noida <sup>S37</sup>	The UHI is assessed with the help of field data, meteorological observations and Landsat thermal dataset for year 2000 and 2013. Estimated UHI showed a negative correlation between NDVI, Emissivity and temperature whereas, NDBI, Albedo and temperature showed a positive correlation. The change in temperature is reported mainly due to increase in impervious areas.
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207 **Supplementary References:**

208 S1. Loveland, T. R., & Belward, A. S. The IGBP-DIS global 1km land cover data set,  
209 DISCover: first results. *Int. J. of Remote Sensing* **18.15**, 3289-3295 (1997).

210 S2. Scean, J. Thematic validation of high-resolution global land-cover data sets. *Photo.*  
211 *Engg. and remote sensing* **65.9**, 1051-1060 (1999).

212 S3. Wang, K., Wang, J., Wang, P. Sparrow, M. Yang, J., Chen, H. Influences of urbanization  
213 on surface characteristics as derived from the Moderate-Resolution Imaging  
214 Spectroradiometer: A case study for the Beijing metropolitan area. *J. Geophys. Res.* **112**,  
215 (2007).

216 S4. Coll, C., Wan, Z., & Joan M. G. Temperature based and radiance based validations of the  
217 V5 MODIS land surface temperature product. *J. of Geo. Res. Atm.* **114**. D20 (2009).

218 S5. Bosilovich, M. G. A comparison of MODIS land surface temperature with in situ  
219 observations. *Geo. Res. Let.* **33.20** (2006).

220 S6. Wan, Z., & Li, Z. L. Radiance based validation of the V5 MODIS land surface  
221 temperature product. *Int. J. of Remote Sensing* **29.17-18**, 5373-5395 (2008).

222 S7. Dash, P., et al. Land surface temperature and emissivity estimation from passive sensor  
223 data: Theory and practice-current trends. *Int. Jour. of remote sensing* **23.13** (2002): 2563-  
224 2594.

225 S8. Wan, Z., Zhang, Y., Zhang, Q., Li, Z. L., Quality assessment and validation of the  
226 MODIS global land surface temperature. *Int. J. Remote Sens.*, **25** (1), 261-274 (2004).

227 S9. Fung W. Y., Lam K. S., Nichol J., Wong M. S. Derivation of night time urban air  
228 temperatures using a satellite thermal image. *J. Appl. Meteorol. Climatol.* **48**, 863–872  
229 (2009).

230 S10. Hung T., Uchihama D, Ochi S., Yasuoka Y. Assessment with satellite data of the urban  
231 heat island effects in Asian mega cities. *Int J. Appl. Earth Obs. Geoinf.* **8**, 34–48 (2006).

232 S11. Mohan, M. et. al., Assessment of urban heat island effect for different land use–land  
233 cover from micrometeorological measurements and remote sensing data for megacity Delhi.  
234 *Theor. and applied climatology* **112.3-4**, 647-658 (2013).

235 S12. Jensen, J. R., Univ. of South Carolina, Columbus, 1986.

236 S13. Richards, J. A., and J. A. Richards. Vol. 3. Berlin et al.: Springer, 1999.

237 S14. Park, S., Cheol, M. Kyu P., & Moon G. K. Super-resolution image reconstruction: a  
238 technical overview. *IEEE signal processing magazine* **20.3**, 21-36 (2003).

239 S15. Sadavarte, P., & Venkataraman, C. Trends in multi-pollutant emissions from a  
240 technology-linked inventory for India: I. Industry & transport sectors, *Atmos. Env.*, **99**, 353-  
241 364 (2014).

242 S16. Pandey, A. et al. Trends in multi-pollutant emissions from a technology-linked  
243 inventory for India: II. Residential, agricultural & informal industry sectors, *Atmos. Env.*, **99**,  
244 341-352 (2014).

245 S17. Kuhlmann, J. & Quaas, J. How can aerosols affect the Asian summer monsoon?  
246 Assessment during three consecutive pre-monsoon seasons from CALIPSO satellite data,  
247 *Atmos. Chem. Phys.*, **10**, 4673-4688 (2010).

248 S18. Dey, S., & Girolamo, L. D. A decade of change in aerosol properties over the Indian  
249 subcontinent, *Geophys. Res. Lett.*, **38**, L14811 (2011).

250 S19. Tripathi, S. N. et. al. Measurements of atmospheric parameters during Indian Space  
251 Research Organization Geosphere Biosphere Programme Land Campaign II at a typical  
252 location in the Ganga basin: 1. Physical & optical properties, *J. Geophys. Res.*, **111**, D23209  
253 (2006).

254 S20. Ramachandra, T. V., and Uttam K. Greater Bangalore: Emerging urban heat island. *GIS*  
255 *Development* **14.1**, 86-104 (2010).

256 S21. Ambinakudige, S. Remote sensing of land cover's effect on surface temperatures: a case  
257 study of the urban heat island in Bangalore, India. *Applied GIS* **7.1** (2011).

258 S22. Faris, A. A., & Reddy Y. S., Estimation of urban heat Island using Landsat-7 ETM+  
259 imagery at Chennai city- a case study. *Int. J. Earth Sci. Engg.* **3.3**, 332-340 (2010).

260 S23. Devadas, M. D., & Rose, A. L. Urban factors and the intensity of heat island in the city  
261 of Chennai. *Proceedings of the Seventh International Conference on Urban Climate.*  
262 *Yokohama, Japan* (2009).

263 S24. Grover, A., & Ram B. S. Analysis of Urban Heat Island (UHI) in Relation to  
264 Normalized Difference Vegetation Index (NDVI): A Comparative Study of Delhi and  
265 Mumbai. *Environments* **2.2**, 125-138 (2015).

266 S25. Mohan, M. et. al. Assessment of urban heat island intensities over Delhi. *Seventh*  
267 *International Conference on Urban Climate, Yokohama, Japan, 2009.*

268 S26. Babazadeh, M., & Parvendra K. Estimation of the urban heat island in local climate  
269 change and vulnerability assessment for air quality in Delhi. *Euro. Sci. J., Sp. Ed.*, 1857-  
270 7431 (2015).

271 S27. Pandey, P., et. al. A study of urban heat island and its association with particulate matter  
272 during winter months over Delhi. *Science of the Total Environment* **414**, 494-507 (2012).

273 S28. Franco, Sainu, et al. Study of temperature profile on various land use and land cover for  
274 emerging heat island. *Journal of Urban and Environmental Engineering (JUEE)* **9.1**, 32-37  
275 (2015).

276 S29. Badarinath, K. V. S., et al. Studies on urban heat islands using ENVISAT AATSR data.  
277 *Journal of the Indian Society of Remote Sensing* **33.4**, 495-501 (2005).

278 S30. Thomas, G., et al. Analysis of urban heat island in Kochi, India, using a modified local  
279 climate zone classification. *Procedia Environmental Sciences* **21**, 3-13 (2014).

280 S31. Kotharkar, R., and Meenal S. Land Use, Land Cover, and Population Density Impact on  
281 the Formation of Canopy Urban Heat Islands through Traverse Survey in the Nagpur Urban  
282 Area, India. *Journal of Urban Planning and Dev.* **04015003** (2015).

283 S32. Deosthali, V., Impact of rapid urban growth on heat and moisture islands in Pune City,  
284 India. *Atmospheric Environment* **34.17**, 2745-2754 (2000).

285 S33. Ansar, S., et. al. A study of urban/rural cooling rates in Thiruvananthapuram, Kerala. *J.*  
286 *Ind. Geophys. Union* **16.1**, 29-36 (2012).

287 S34. Gupta, A., Dey, V., & Goel. A., Mitigation of Thermal Pollution to enhance urban air  
288 quality through Remote Sensing and GIS. 10th ESRI India User Conf. 2009 Noida (NCR)  
289 (2009).

290 S35. Joshi, Ram, et al. Urban Heat Island Characterization and Isotherm Mapping Using  
291 Geo-Informatics Technology in Ahmedabad City, Gujarat State, India. *Int. J. of Geosci.* **6.3**,  
292 274 (2015).

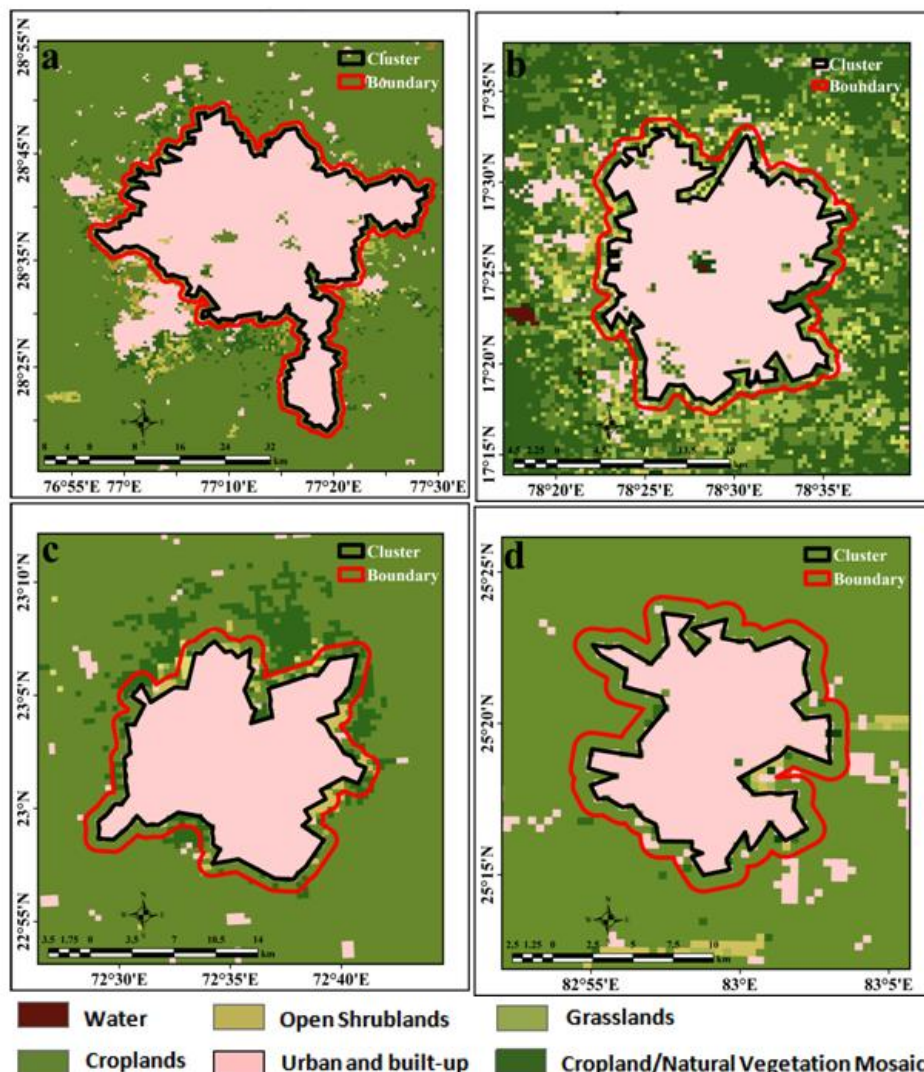
293 S36. Maral, S. G., & Mukhopadhyay, T. Signal of urban heat island (UHI) effect: A case  
294 study of Mumbai metropolitan region. *Signal* **551.522**: 540-53.

295 S37. Kikon, N., et al. Assessment of urban heat islands (UHI) of Noida City, India using  
296 multi-temporal satellite data. *Sustainable Cities and Society* **22**, 19-28 (2016).

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## Supplementary Figures



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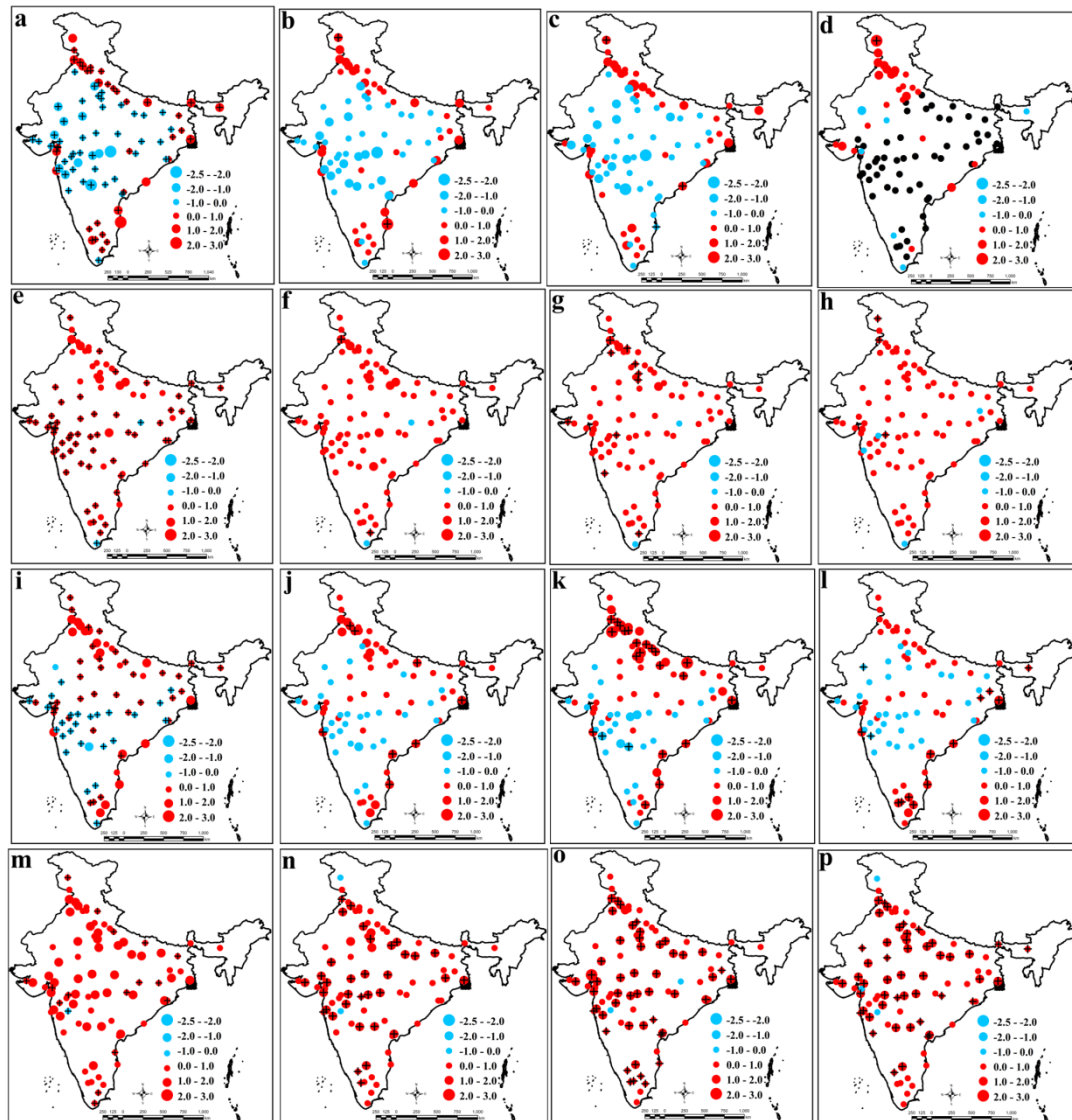
### Supplementary Figure S1.

The selection of urban and surrounding non-urban region is presented

(a) Identified urban cluster of New Delhi, the pink area represents the cluster for urban centre, while its boundary is outlined with black line. The green areas are other surrounding region near an urban centre, where the land cover is characterized by a large share of crop areas. The boundary of nearby non-urban region is outlined with red line. (b,c,d) same as (a) for the urban centres of Hyderabad, Ahmedabad and Varansi respectively. The maps are generated with Arc Map Ver 10.2 (<http://www.esri.com/software/arcgis/arcgis-for-desktop>)

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315 **Supplementary Figure S2.**

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317 **Monthly Surface Urban Heat Island Intensity (SUHII) in Indian cities**

318 SUHII during summer day (a) with the monthly SUHII over the month of March (b) April

319 (c) and May (d); (e-h) same as (a-d) but for summer night time SUHII. Similarly (i-l) and

320 (m-p), are presented for the winter season day and night time monthly SUHII

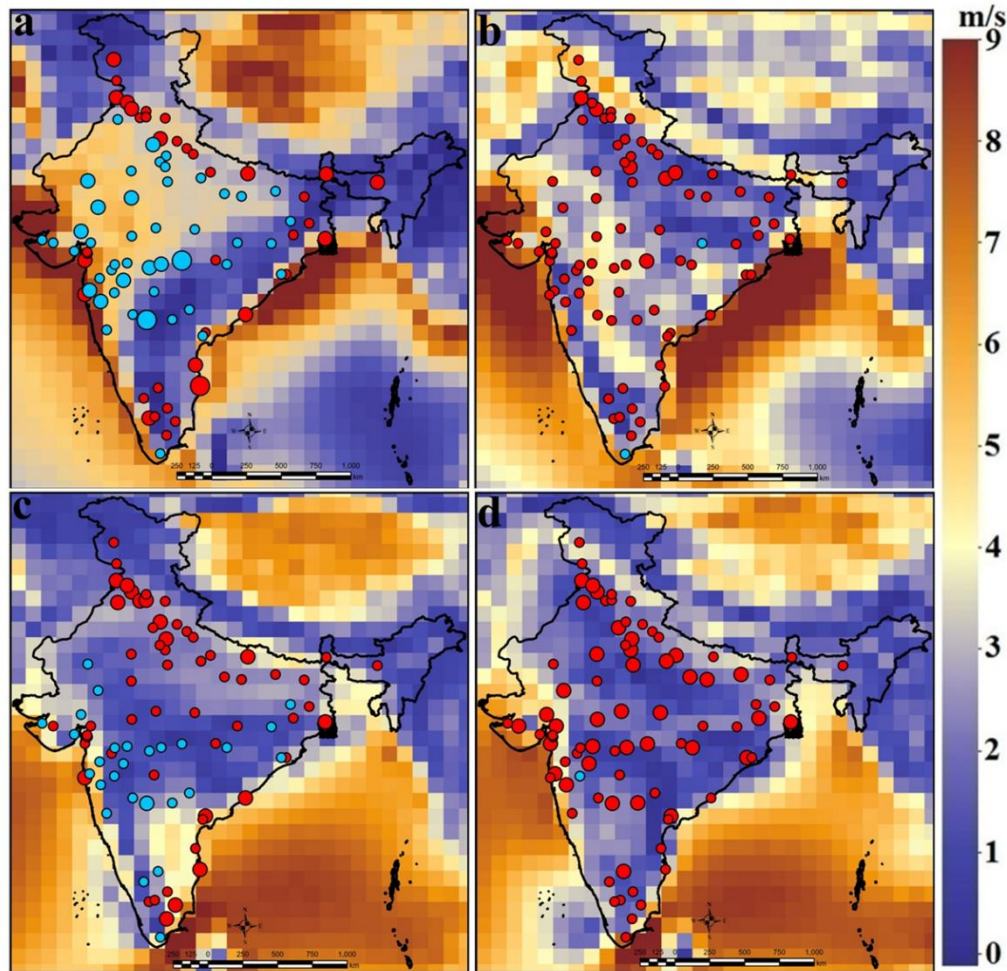
321 respectively. The cities, for which land surface temperature differences between urban

322 and surrounding non-urban areas are statistically significant, are shown with “+”. The

323 maps are generated with Arc Map Ver 10.2 (<http://www.esri.com/software/arcgis/arcgis->  
324 [for-desktop](http://www.esri.com/software/arcgis/arcgis-for-desktop))

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**Supplementary Figure S3.**

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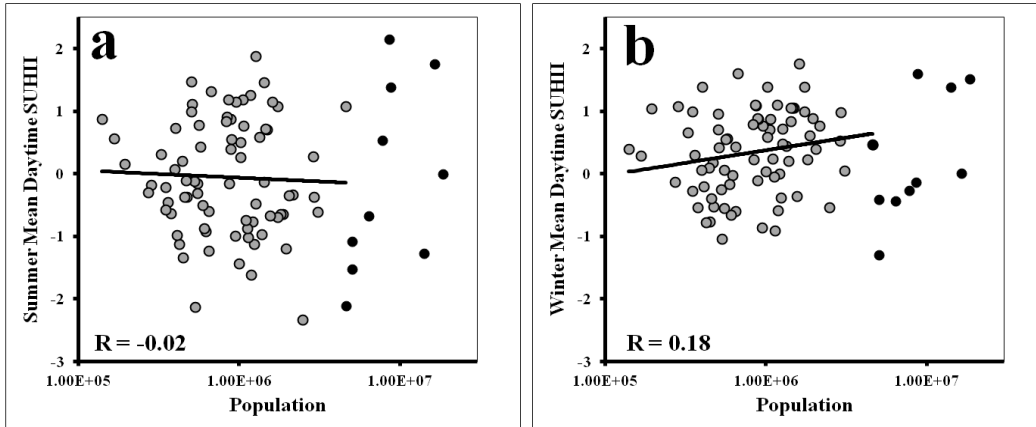
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Association of SUHII in India to the surface wind. The SUHII for summer day, summer night, winter day and winter night are overlaid on the surface wind map ((a) to (d) respectively) as obtained from reanalysis data. The maps are generated with Arc Map Ver 10.2 (<http://www.esri.com/software/arcgis/arcgis-for-desktop>)



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337 **Supplementary Figure S4.**

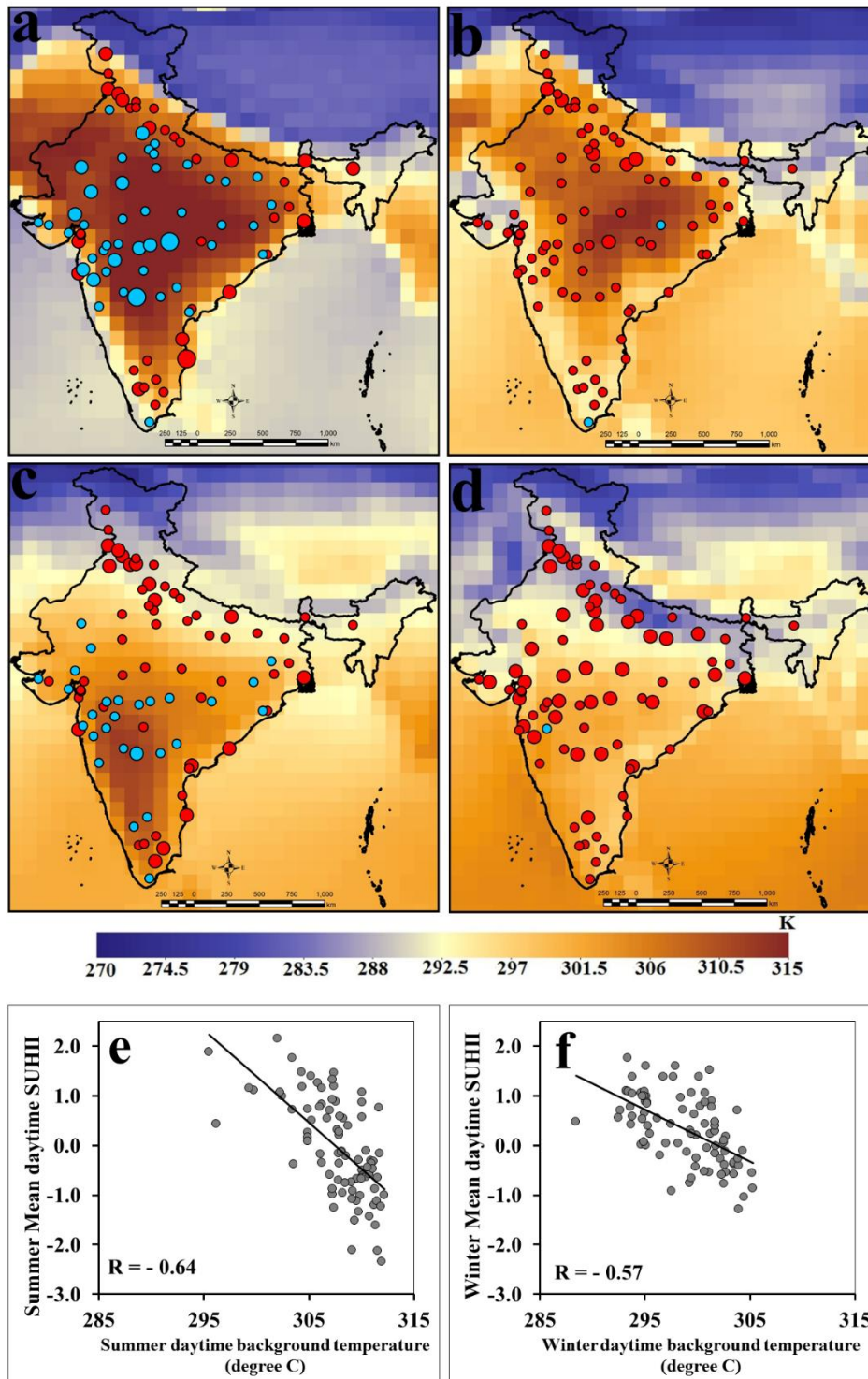
338 Sensitivity of summer (a) and winter (b) day time SUHII to the population.

339 ([http://in.mathworks.com/products/new\\_products/release2012b.html](http://in.mathworks.com/products/new_products/release2012b.html))

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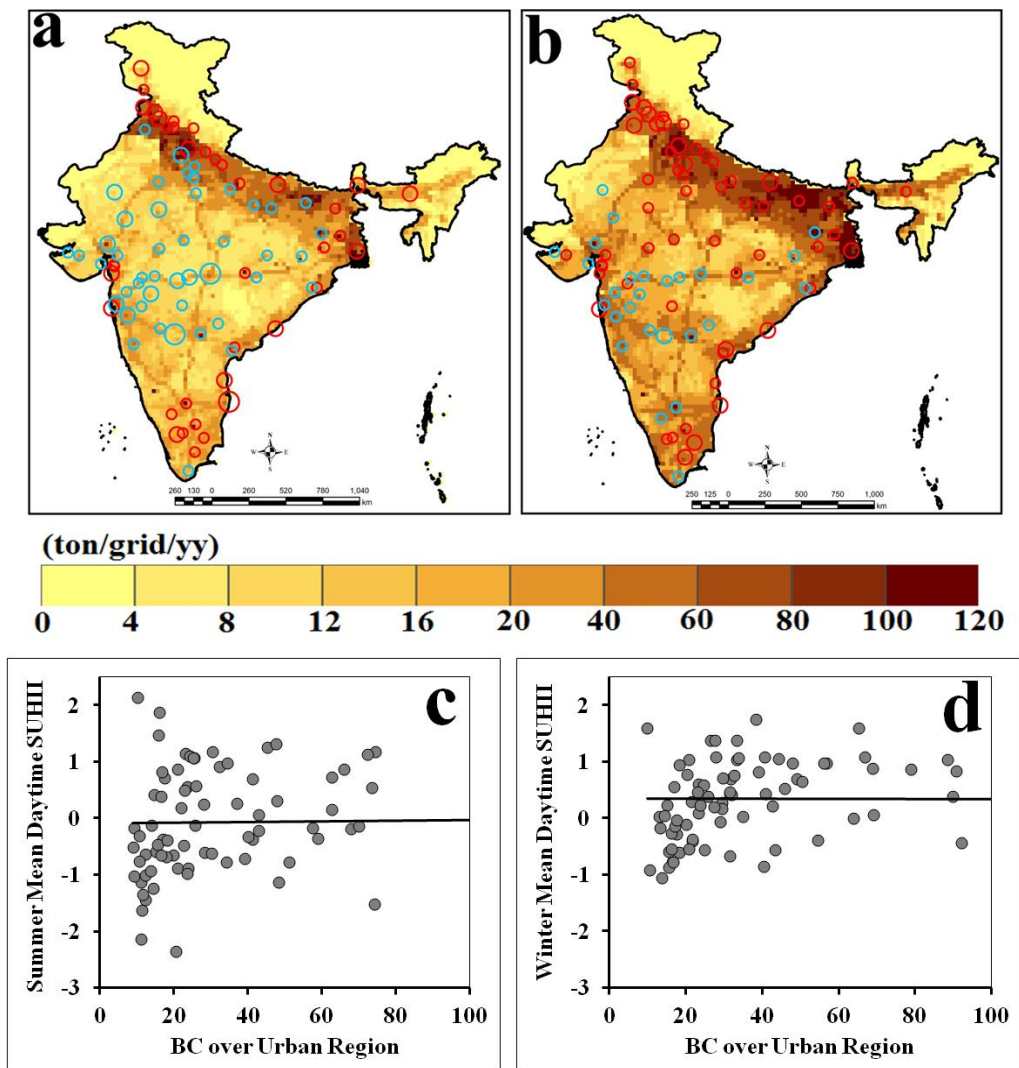
344 **Supplementary Figure S5.**

345 Association of SUHII in India with background air temperature. The SUHII for summer

346 day, summer night, winter day and winter night are overlaid on the temperature ((a) to (d)

347 respectively) map as obtained from reanalysis data. Summer and winter daytime SUHII is

348 negatively associated with background temperature ((e) and (f) respectively). The red and  
349 blue circles denote the SUHII, similar to Figure 1. The maps are generated with Arc Map  
350 Ver 10.2 (<http://www.esri.com/software/arcgis/arcgis-for-desktop>) and, MATLAB  
351 R2012b ([http://in.mathworks.com/products/new\\_products/release2012b.html](http://in.mathworks.com/products/new_products/release2012b.html))  
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### 354 **Supplementary Figure S6**

#### 355 **The effect of seasonality in Black Carbon (BC) emissions on SUHII.**

356 The SUHII during summer and winter day-time are overlaid on the BC emission colour  
357 map ((a) and (b) respectively). The red and blue circles denote the same as main figure  
358 Sensitivity of day time SUHII to the BC emissions. Summer and winter daytime SUHII

359 show no association with the BC emissions ((a) and (b) respectively). The maps are  
360 generated with Arc Map Ver 10.2.

361 (<http://www.esri.com/software/arcgis/arcgis-for-desktop>) and, MATLAB R2012b

362 ([http://in.mathworks.com/products/new\\_products/release2012b.html](http://in.mathworks.com/products/new_products/release2012b.html))

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