The Cryosphere Discuss., 7, 5921–5970, 2013 www.the-cryosphere-discuss.net/7/5921/2013/ doi:10.5194/tcd-7-5921-2013 © Author(s) 2013. CC Attribution 3.0 License.



This discussion paper is/has been under review for the journal The Cryosphere (TC). Please refer to the corresponding final paper in TC if available.

Weekly-gridded Aquarius L-band radiometer/scatterometer observations and salinity retrievals over the polar regions: applications for cryospheric studies

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Received: 25 November 2013 – Accepted: 28 November 2013 – Published: 12 December 2013

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Published by Copernicus Publications on behalf of the European Geosciences Union.





Abstract

Passive and active observations at L band (frequency \sim 1.413 GHz) from the Aquarius/SAC-D mission offer new capabilities to study the polar regions. However, due to the lack of polar-gridded products, applications over the cryosphere are limited.

- ⁵ To allow for an efficient use of the Aquarius data over the polar regions, and to move forward our understanding of the L-band observations of ice sheet, sea ice, permafrost, and polar oceans, we present three weekly-polar-gridded products. Aquarius data at latitudes higher than 50° were averaged and gridded into three weekly products of Brightness Temperature (TB), Normalized Radar Cross Section (NRCS), and Sea Sur-10 face Salinity (SSS). These products are suited for cryospheric studies, and each grid
- cell contains sea ice fraction and the standard deviation of TB, NRCS, and SSS along with the number of footprint observations collected during the seven-day cycle. The data sets are produced on the version 2.0 Equal-Area Scalable Earth (EASE2.0) grid, with a grid cell resolution of 36 km. This paper describes the products, and presents
- ¹⁵ maps and time series of Aquarius weekly-gridded data over the Greenland and Antarctic ice sheets, sea ice in both hemispheres, subarctic land where seasonal snow and frozen soil may exist, and the polar oceans. A brief analysis of L-band observations is given to encourage future use of the products. They can be used for improving our understanding of low microwave frequency observations, and for the development of algorithms. The new weekly-polar-gridded datasets start in August 2011, with the first
- Aquarius observations, and it is anticipated that they will be updated on a monthly basis following the release schedule of the Level 2 data sets.

1 Introduction

The cryosphere has been extensively studied using passive microwave satellite observations for over four decades. The first imaging radiometer, the Electrically Scanning Microwave Radiometer, launched in 1972 provided the earliest maps of the global sea





ice cover (Zwally et al., 1983). In 1978, the longest continuous passive microwave data record started with the launch of the Scanning Multichannel Microwave Radiometer (SMMR). This data record, used for example to monitor sea-ice cover properties, ice sheet melt, and snow on land, continues to the present with the series of Special Sen-

⁵ sor Microwave Imager Sounder (SSM/I, SSMIS), and Advanced Microwave Scanning Radiometers (AMSR-E, AMSR2). The lowest frequency observations available at high latitudes were available at ~ 7 GHz with SMMR and the AMSRs (and only at 18 GHz with the SSM/Is).

In the 1970s, observations at 1.4 GHz were collected by the Skylab S-194 radiometer (Jackson et al., 2004), however the orbit characteristics did not enable the monitoring of the polar regions. Since late 2009, passive L-band observations from the Soil Moisture and Ocean Salinity (SMOS) mission are routinely collected over both polar regions. In addition, the Aquarius/SAC-D mission provides both passive and active L-band observations since August 2011. Similar observations will be collected with the forthcoming

Soil Moisture Active/Passive (SMAP) mission (whose launch is currently planned for late 2014). Despite the fact that these missions have been primarily designed for the monitoring of either Sea Surface Salinity (SSS) or soil moisture, new applications are being developed to study the cryosphere.

In the microwave domain, dry snow and ice are low-loss media and the radiation penetrates into the snow/ice cover. Penetration of microwaves is inversely proportional to losses due to scattering and absorption, and thus increases as frequency decreases. L-band dry snow and ice dielectric losses are very small (e.g., Mätzler, 1987; Warren et al., 2008), and the L-band radiation likely emanates from one to two orders of magnitude deeper than the radiation observed at 6.9 GHz (Surdyk, 2002). Thus, the

L-band radiation can emanate from hundreds to thousands of meters deep (i.e. centuries old ice over the Antarctic plateau for example), which opens new research topics and motivates the development of innovative approaches to study long-term climatic changes over the ice sheets. Moreover, specific regions of the Antarctic ice sheet can be used as calibration and validation sites for L-band radiometers (e.g., Drinkwater





et al., 2004; Macelloni et al., 2013). Additional applications are possible, since investigations showed that the L-band radiometric observations contain information about sea ice thickness (e.g., Kaleschke et al., 2010), and soil freeze/thaw state in subarctic environments (e.g., Rautiainen et al., 2012).

- ⁵ This study specifically deals with the Aquarius/SAC-D mission, developed collaboratively between the US National Aeronautics and Space Administration (NASA) and Argentinian's space agency, Comisión Nacional de Actividades Espaciales (CONAE). The mission was designed to monitor SSS using a combination of passive and active microwave observations. Aquarius has three L-band radiometers, providing Brightness
- Temperature (TB) at 1.413 GHz at vertical (V) and horizontal (H) polarizations. It also has one scatterometer providing Normalized Radar Cross Section (NRCS) observations at 1.26 GHz using the same three feed horns as the radiometers, and thus coincident with the radiometers' observations.

While Aquarius observations are recorded in a large footprint (from 94 km × 76 km to 156 km × 97 km, as incidence angle increases), its observations are acquired with a fine sensitivity (Noise Equivalent Differential Temperature, NEDT) of 0.15 K (Le Vine et al., 2010). In comparison, SMOS has a finer spatial resolution that depends on the viewing angle and the data processing. The 3-dB half power beam width corresponds to a 43 km footprint on average over the field of view, but ranges between 30 km × 30 km

- to 90 km × 33 km depending on the incidence angle (Kerr et al., 2010). SMOS, however, has lower sensitivity 2.0–2.5 K and accuracy of 3.7–4.03 K than Aquarius (Martin-Neira et al., 2010). The forthcoming SMAP mission is designed with an accuracy of 1.3 K, and a 3-dB footprint of the order of 40 km (Entekhabi et al., 2010), placing it between SMOS and Aquarius in spatial resolution and accuracy. SMAP scatterometer will also
- have a finer spatial resolution (~ 3 km) than Aquarius' scatterometer (Entekhabi et al., 2010; Brown et al., 2013).

Aquarius data are distributed by NASA's Physical Oceanography Distributed Active Archive Center (PO.DAAC). The Aquarius Level 2 product consists of the observations and retrievals along track. The Aquarius Level 3 products are 1° gridded SSS, averaged





over different time scales (e.g. daily, weekly, monthly, and seasonal time scales) (Wentz and Le Vine, 2012; Le Vine et al., 2012). Therefore, the Aquarius products available do not include weekly-gridded data sets of TB and NRCS appropriate for studying the Earth's polar regions and the cryosphere. Moreover, the Level 3 SSS products use a projection that is not optimal for studying the high latitudes.

To allow for an efficient use of the Aquarius data over the polar regions, and to move forward our understanding of the L-band observations of ice sheet, sea ice, permafrost, and polar oceans, we present in this manuscript three weekly-polar-gridded products of TB, NRCS, and SSS. These products also include sea ice fraction (ICEF). The gridded products were produced on the version 2.0 Equal-Area Scalable Earth (EASE) grid (Brodzik et al., 2012). Each grid cell is 36 km × 36 km. This grid and resolution were selected because they offer a fair compromise between Aquarius' high sampling along track (an observation every 10 km) and the distance between tracks, and they will be

¹⁵ from the seven-day satellite cycle.

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We describe the Aquarius radiometers and scatterometer in Sect. 2 along with the criteria for valid data. In Sect. 3, we introduce the details of the Aquarius weekly-polar-gridded products. We present the weekly-polar-gridded maps resulting of the TB and NRCS products in Sects. 4 and 5, respectively, with a focus on the Greenland and Antarctic ice sheets, and sea ice in both hemispheres. In Sects. 6 and 7, we present

used by SMAP. Gridded data were produced using all the valid Aquarius observations

²⁰ Antarctic ice sheets, and sea ice in both hemispheres. In Sects. 6 and 7, we present data over land where seasonal snow and frozen soil exist, and maps of polar ocean salinities to discuss the effect of sea ice contamination.

Weekly-polar-gridded products presented in this paper are available on the NASA Cryosphere Science Research Portal http://neptune.gsfc.nasa.gov/csb/index.php?

25 section=273. They cover the entire Aquarius mission period since its start in August 2011. It is anticipated that the products will be updated on a monthly basis following the release schedule of the Level 2 data sets by PO.DAAC.



2 Aquarius data

The passive Aquarius instruments consist of three radiometers providing TB at 1.413 GHz in V and H polarizations. The third Stokes parameter is also measured, but assumed null at the surface. Observations of the third Stokes parameter are used to retrieve properties of the Faraday rotation in the ionosphere (Yueh, 2000), and the 5 antenna cross-polarization coupling (Le Vine et al., 2011; Kim et al., 2011). The three radiometers (1, 2, and 3, at the incidence angles of 28.7°, 37.8°, and 45.6°, respectively) are in a push-broom alignment (i.e. non-scanning) pointing to the right-hand side of the orbit. Most of the time, they point to the night side of the sun-synchronous orbit, but solar contamination in the main beams is possible at high latitudes season-10 ally (Dinnat and Le Vine, 2008). Other radiometers used to the study of the cryosphere typically operate at incidence angles of 53-55°, closer to the Brewster angle for the snow/air interface. The lower Aquarius incidence angles imply that observations are more sensitive to the snow/air interface, and surface snow metamorphism (Brucker et al., 2013b). 15

Table 1 provides the radiometers' footprint dimensions, which are increasing as the incidence angle increases. The spacecraft orbits with a seven-day repeat cycle. These characteristics, and the fact that the radiometers are in a push-broom alignment, imply that the spatial coverage becomes sparser as latitude decreases. According to the orbit inclination and the different radiometers' incidence angles, the areas unobserved at the poles differ for each sensor and hemisphere (Table 1). In the Northern Hemisphere (NH), the largest coverage is obtained using beam 3, whereas in the Southern Hemisphere (SH) beam 3 has the smallest coverage.

As mentioned previously, the accuracy and stability of Aquarius radiometric observations is remarkably good. The radiometer NEDT is 0.15 K, and observations have shown stability within ±0.2 K over the ocean and celestial Sky over several months (Dinnat et al., 2013; Lagerloef et al., 2013).





Aquarius' active instrument consists of one scatterometer observing NRCS at 1.26 GHz at four polarizations: VV, VH, HV, and HH. Despite the fact that antenna NRCS are measured at the four polarizations, the NRCS at the two cross-polarizations are assumed identical at the top of the atmosphere. Their differences at the antenna

- is assumed to be due to Faraday rotation, and antenna cross-polarization coupling. Therefore, only NRCS from one cross-polarization channel is considered in this product. The scatterometer uses the same three feed horns as the radiometers, providing NRCS observations at the same incidence angles as the TB observations. The footprint sizes are slightly smaller due to the two-way path from the spacecraft to the surface
 and back, despite the lower frequency. Scatterometer's sensitivity varies with incidence
 - angle (Table 2), and the seven-day stability is $\pm 0.1 \text{ dB}$.

Both passive and active L-band observations are impacted by Radio Frequency Interference (RFI). Although the current L-band missions operate their radiometers in a protected spectral band, RFI is an issue, especially in the NH. Aquarius has a high

- ¹⁵ radiometric sensitivity and short-time sampling that enhance detection of low level RFI. In the weekly-gridded products presented in the next sections, every data associated with an RFI flag in the Level 2 products is eliminated. An example of weekly RFI occurrence is presented in Fig. 1. RFI contamination is persistent over Northern and Eastern Europe, and along the North Warning System (NWS) at a latitude of ~ 70° N
- in Northern America. The RFI resulting from the NWS impacts Aquarius observations in different regions and with a different intensity whether the spacecraft is in an ascending or descending orbit. Of note, there is RFI contamination over Greenland only during descending orbits when Aquarius sensors are pointing toward the NWS line (Fig. 1, right). In the high latitudes of the SH L-band RFI occurrences are rare. Us-
- ing the 6.8 and 10.7 GHz channels (C and X bands) of the WindSat radiometers, RFI at these frequencies were collocated with research stations in the Antarctic Peninsula (Zhao et al., 2013).

To develop the Aquarius weekly-polar-gridded products for latitudes higher than 50°, we only used information from the Level 2 product (Wentz and Le Vine, 2012). The





Aquarius Level 2 product (version 2.0, September 2013) consists of the footprint observations and retrievals along track (i.e. the swath data). It also consists of ancillary data, flags, converted telemetry, and navigation data. One example of ancillary data is the ICEF. This variable is not retrieved with the satellite observations. The ICEF results

- from a combination of two elements: simulated sea ice concentration, and antenna 5 characteristics. Sea ice concentration is simulated by the US National Centers for Environmental Prediction (NCEP), and distributed as the Global Forecast System (GFS) Global Data Assimilation System (GDAS) sea ice product. Aquarius sea ice fraction is thus simulated, and differs from the passive microwave satellite sea ice concentration
- products (e.g., Comiso et al., 2003). Hereafter, the use of ICEF refers to the simulated 10 sea ice concentration integrated over the sensor's field of view and weighted by the antenna gain patterns. Similarly, land fraction data is integrated over the sensor's field of view and weighted by the antenna gain patterns (Dinnat and Le Vine, 2007; Le Vine et al., 2011).
- Level 2 TB product is computed after empirical calibration of the measured antenna 15 temperatures against a forward radiative transfer model. The antenna temperatures are then corrected for the emission of extraterrestrial sources (Sun, Moon, celestial Sky), the effect of the integration over the antenna gain patterns, the Faraday rotation, and atmospheric effects. This provides the TBs at the Earth's surface. Over the oceans, the
- TB is corrected for the effects of surface roughness using the scatterometer observa-20 tions, and a simulated wind speed. The remaining TB for a smooth surface is converted into SSS using ancillary data for the sea surface temperature, and a model for the sea water dielectric constant. Details can be found in Wentz and Le Vine (2012); Le Vine et al. (2012) and Piepmeier et al. (2013). We use these Aquarius Level 2 data to pro-
- duce weekly-gridded products of TB (at V and H polarizations), NRCS (at VV, VH, and 25 HH polarizations), SSS, and ICEF.



3 Weekly-polar-gridded product details

A gridded product has three main characteristics: projection, spatial resolution, and temporal resolution. These characteristics were defined according to the satellite observations, and to ensure consistency with the forthcoming SMAP mission. The EASE2.0 grid (Brodzik et al., 2012) was chosen, and the grid cell spatial resolution

- EASE2.0 grid (Brodzik et al., 2012) was chosen, and the grid cell spatial resolution set to 36 km × 36 km. These characteristics are appropriate for the Aquarius observations at latitudes higher than 50°. At these high latitudes, there are limited areas not observed. Therefore, the trade-off between spatial resolution and data coverage is fair. Since Aquarius observations are, however, recorded every 10 km along track, it is possible for specific applications to grid the data with a refined spatial resolution better
- than 36 km. The temporal resolution of the gridded product was defined by the time of revisit: seven days.

Gridded data were produced using all the Aquarius observations deemed as RFI free, and recorded during nominal operation of the spacecraft. Non-nominal operation dates include maneuvers and anomaly periods.

For TB, each radiometer was treated independently to produce three weekly-gridded TB products corresponding to observations from the ascending orbit, the descending orbit, and both orbit types combined. The distinction per beam is necessary due to the different incidence angles, which lead to different reflection intensities at the interface created by dielectric constant changes (e.g. snow/air interface) (Brucker et al., 2013b).

For each seven-day cycle, all observations from a given radiometer (RFI free and recorded during nominal operation) whose footprint center was within a given 36 km grid cell were averaged together, and the standard deviation was calculated.

For NRCS, each beam and each orbit type was treated independently. The distinction for each beam is also needed due to the different incidence angles. The distinction between each orbit type is necessary because of the anisotropy of the ice covered surfaces (addressed in Sect. 5.1).





For SSS, two weekly-gridded products are available. The first product provides the SSS corresponding to the weekly-gridded TB: data are distributed separated per beam and per orbit type, and include combined ascending and descending orbits. The second product was generated combining the SSS retrieved from the three beams and both orbit types. In the latter case, the weekly-gridded SSS values are calculated using the maximum amount of retrievals.

SSS retrievals where the radiometer land fraction was ≥ 0.25 were not considered.
If left uncorrected, a land fraction of as little as 0.001 can be detrimental to the SSS retrieval (~ 0.2–0.4 psu error; Dinnat and Le Vine, 2007). However, there is a first order
correction for the effect of land contamination in the Level 2 SSS retrieval that should remove most of the error (Sect. 3.8; Wentz and Le Vine, 2012). In addition, the level of acceptable residual error in coastal studies will depend on specific applications of the product. For these reasons, we set a relatively high threshold for land contamination. Because we provide the information on the land fraction for each grid cell, the user has
the opportunity to select a more stringent threshold when needed.

TB, NRCS, and SSS weekly averages are provided in conjunction with the standard deviations of all footprint observations/retrievals with the given grid cell. Of note, the statistical significance of the standard deviation has to be evaluated against the number of footprint observations (NFP) in the grid cell.

²⁰ Grid cells without data can result from missing observations either because the cells fall in between beam tracks (permanently missing), or because all data gathered were deemed invalid for a given time period. In these situations, a linear interpolation of the weekly-gridded TB and NRCS was applied in grid cells without observations during the cycle. To avoid interpolating over too large a region in case of significant data loss,

the interpolation was not applied (and a NaN element is provided) where six or more contiguous grid cells had no data. Note that because each product contains the number of footprint observations per grid cell, it is easy to identify interpolated values where the number of observations per grid cell is 0. This allows users to either disregard the interpolated data, or apply another interpolation method best suited for their studies.



In any case, one must be cautious with the use of interpolated data, even though the Aquarius radiometer field of views are larger than the 36 km grid cell resolution.

The TB, NRCS, and SSS products are processed for every Aquarius cycle, and distributed in HDF5 format at http://neptune.gsfc.nasa.gov/csb/index.php?section=273.

⁵ Future release dates depend on the release of the Level 2 product, which are currently released monthly in a one month batch. We also provide ancillary files containing the grid cell center latitude and longitude, and the grid cell mean land fraction, as well as a look-up table to convert date to orbit cycle number. Table 3 summarizes the data distributed in each of the three products.

10 3.1 Aquarius 36 km weekly-polar-gridded TB, and SSS

This product containing TB, SSS, and ICEF is composed of six files per cycle: one per beam, and one per hemisphere. Each file contains the weekly-gridded TB, SSS, and ICEF of the ascending orbits, descending orbits, and the two orbit types combined. The data sets consist of weekly mean TB V, TB H, and SSS, along with the standard deviations within the grid cell (TBV_STD, TBH_STD, and SSS_STD). This product also contains the radiometer ice fraction (ICEF_RAD and ICEF_STD_RAD for the mean and standard deviation, respectively). Standard deviation values are never interpolated.

3.2 Aquarius 36 km weekly-polar-gridded NRCS

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This product containing NRCS and ICEF is composed of six files per cycle: one per
 ²⁰ beam, and one per hemisphere. Each file contains the weekly-gridded data of the ascending and descending orbits. NRCS observations from ascending and descending orbits were not combined together due to the anisotropy of the ice covered surfaces. The data consist of weekly mean NRCS at three polarizations (NRCS_VV, NRCS_VH, and NRCS_HH), along with the standard deviations within the grid cell
 ²⁵ (NRCS VV STD, NRCS VH STD, and NRCS HH STD). This product also contains





the scatterometer ice fraction (ICEF_SCA and ICEF_STD_SCA for the mean and standard deviation, respectively).

3.3 Aquarius 36 km weekly-polar-gridded SSS3b

This product combines SSS retrievals from all three beams, as retrievals of SSS should ⁵ be independent of the beam (i.e. of the incidence angle). However, the error on SSS retrievals could vary slightly with the beam as the outer beam averages over a larger field of view and is more sensitive to surface roughness (in horizontal polarization) for example. This product is composed of two files per cycle: one per hemisphere. Each file contains the weekly-gridded SSS of the ascending orbit, descending orbit, and the two orbit types combined, along with the standard deviation within the grid cell (SSS_STD). The distinction between the two orbit types is needed since differences have been identified (Lagerloef et al., 2013).

4 L-band TB product analysis over the ice sheets

Microwave observations are sensitive to snow properties (e.g. liquid water, grain size, density, temperature) (Ulaby et al., 1986; Mätzler, 1987) and to their vertical variations (Brucker et al., 2010, 2011). This sensitivity to snow properties has been used to monitor various geophysical properties over the ice sheets, such as snow melt extent (e.g., Torinesi et al., 2003; Tedesco et al., 2007), snow accumulation (e.g., Abdalati and Steffen, 1998; Winebrenner et al., 2001), and grain size (Brucker et al., 2010; Picard et al.,

20 2012). This section aims at providing for the first time an overall presentation of the Aquarius weekly-polar-gridded TB over the Greenland Ice Sheet (GIS) and Antarctic Ice Sheet (AIS). It is beyond the scope of this study to address the specifics in the data, and to relate quantitatively the remote sensing observations to geophysical properties. The products presented here enable these in-depth analyses for future research.





The GIS and AIS have craggy coastal regions, mountainous areas, and thus present steep slopes in their peripheries, and multiple surface types in the grid cells. Ice sheet roughness is controlled by bedrock topography, and by snow accumulation, erosion, and redistribution (Furukawa and Young, 1997; Long and Drinkwater, 2000). Topo-⁵ graphic variations within the Aquarius 3-dB footprint contribute to the L-band observations, and could, therefore impact the retrievals of geophysical properties; evidences have been identified over North Africa and Australia (Utku and Le Vine, 2013). The interpretation of the L-band observations should thus be made with caution in areas with strong elevation change and heterogeneity.

10 4.1 The Greenland ice sheet

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Interesting spatial and temporal variations are revealed by maps of L-band TB observations over the GIS (Fig. 2). Overall, TBs are high (> 250 K) year round in the northern coastal regions (extending up to ~ 200 km in land). Whereas the other coastal regions show a seasonal cycle with lower TBs in summer than in winter. This cycle directly results from the sea-ice seasonal cycle, and the presence of open water in summer. In North and northeast Greenland, sea ice is present year round, whereas it is not on the other Greenlandic coasts (see ice concentration maps in Brucker et al., 2013a).

Farther inland, TBs increase from \sim 190 K to \sim 230 K as elevation increases, and therefore as temperature, snow accumulation, and days of melting snow decrease.

- The lowest TBs (~ 180 K) are found in the central part of south Greenland at latitudes between ~ 62° N and ~ 68° N. Additionally, at these latitudes there is a clear distinction between the eastern and western sides of the GIS, with systematically lower TBs on the eastern side. This difference comes from the different ice sheet properties. Southeast Greenland has the highest snow accumulation (Burgess et al., 2010), whereas
- 25 southwest Greenland has lower snow accumulation and more melt; a large part of this latter region is in the ablation zone. Therefore, the L-band radiation emanates mainly from ice/firn/snow in southeast Greenland, and from ice/snow in southwest Greenland.





Distributions of the weekly-gridded TB (beam 3) over the GIS for a cycle in January and July are almost similar at vertical polarization, and slightly change at horizontal polarization (Fig. 3). TBs have a range of about 180–260 K and 140–240 K at V and H polarizations, respectively. In summer, lower TBs exist at both polarizations, and they ⁵ correspond to grid cells located in coastal regions where open water is present in the instrument field of view.

Using radiometer 1, the timeseries of TB (Fig. 4) near Summit (72.484° N, 38.246° W, snow accumulation of 0.23 m w.e., Burgess et al., 2010) and South Dome (65.168° N, 43.410° W, snow accumulation of 0.49 m w.e., Burgess et al., 2010) show that there is no annual cycle related to variations in the temperature profile. However, observations in summer are impacted by snow melt. Observations in the South Dome area show a sudden increase of TB associated with the presence of liquid water in the snow cover, followed later in the summer by a slower TB decrease as water refreezes and snow grain size and density increase. This is a TB evolution typical at microwave frequencies when snow is melting.

During the summer of 2012, unusual melt occured over most of the GIS including Summit (Hall et al., 2013). During the period September 2011–September 2013, the mean TB at Summit was 222.9 ± 3.2 K and 210.4 ± 3.9 K at V and H polarization, respectively. TBs started to increase by ~ 8 K in late May 2012, and suddenly dropped by more than 15 K and 20 K at V and H polarizations, respectively. The drop occurred during the Aquarius cycle 47, 12–19 July 2012. Melt occurred at Summit during 4–5 days, ending on 14 July 2012 (Hall et al., 2013). An analysis of the swath data showed that the lowest TB was observed on 11 July. While melting snow has a high microwave emissivity implying high TB, Aquarius did not capture the high TBs at Summit charac-

teristic of melting snow. It is worth noting though that the TB level after the melt event is ~ 5 K lower, revealing the significant impact of the recently formed near-surface refrozen snow layer on TB. Even though the L-band radiation emanates from deep in the ice sheet, surface reflection is an important process of the radiative transfer. It has been demonstrated that regions where melt occur can bias snow accumulation retrievals us-





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ing microwave TB (Magand et al., 2008), Fig. 4 shows that similar bias can be expected using L-band observations.

4.2 The Antarctic ice sheet

- Low TBs (< 200 K) observed in January and February in the continent periphery are due to the presence of the surrounding sea-ice-free ocean (Fig. 5). This effect is only visible on a narrow band of latitudes. Slightly farther inland, there is another narrow band of latitudes with very high TBs (250–265 K) year round. The spatial extent of these high TBs progresses inland as snow is melting, and reaches its maximum in January.
- ¹⁰ The lowest TBs not contaminated by sea-ice-free ocean are located on the high elevation domes of the Antarctic Plateau (eastern part of the continent). The lowest mean-annual TBs of the Plateau are located at both polarizations in grid cells within or adjacent to the Vostok Lake (76.396° S, 105.852° E). This suggest that the radiometrically cold lake and/or the particular ice properties over the lake influence the L-band
- observations. There is also an obvious contrast between the Antarctic Plateau, and the West AIS (WAIS). The Plateau has higher elevations, lower snow accumulations, no melt events, and thus lower TB than Mary Byrd Land (Fig. 5). Several previous studies discuss these spatial distributions of microwave observations at higher frequencies and their correlation with geophysical properties (e.g., Long and Drinkwater, 2000; Schnei-
- ²⁰ der and Steig, 2002; Schneider et al., 2004; Picard et al., 2009; Brucker et al., 2010), which are also evident at L band.

Strong and direct interactions between ice sheet and ocean occur under ice shelves. The Amery ice shelf and ice shelves in the Peninsula (e.g. Larsen C located 67.5° S, 62.5° W) are characterized by low winter TBs (Fig. 5). This suggests that Aquarius observations could be analyzed in these regions to study the relationship between L-

observations could be analyzed in these regions to study the relationship between Lband TBs and ice shelf properties (e.g. thickness). Future studies could also investigate the different processes that can operate at the ice-shelf base (e.g. refreezing, or melting releasing fresher water initiating deep water formation).





Considering one weekly-gridded TB data set in winter (July) and in summer (January), the ranges of TB for the whole observed AIS are 206–270 K and 178–250 K at V and H polarizations, respectively (Fig. 6). At both polarizations, the range of TB is smaller than that observed over the GIS. Distributions slightly change between the two seasons, with more grid cells having higher TBs in summer, but the ranges do not vary much. As noted earlier, the lowest summer TBs are due to the presence of open water along the coast.

The Antarctic ice sheet can be used as a target for sensor calibration and intercalibration at L band. Dome C is a potential candidate (Drinkwater et al., 2004; Macelloni et al. 2012) Over the period Sentember 2011, Sentember 2012, the mean TBe ab

- to et al., 2013). Over the period September 2011–September 2013, the mean TBs observed by Aquarius' radiometer 1 at Dome C (75.1° S, 123.35° E, snow accumulation rates of 7.2 cm yr⁻¹ (Urbini et al., 2008)) are 202.53 \pm 0.26 K, and 190.12 \pm 0.47 K at V and H polarization, respectively (Fig. 7). These TB standard deviations over the two-year period are low, but they are higher than the radiometers' sensitivity of 0.15 K. Over
- ¹⁵ the WAIS, in one grid cell of the Mary Byrd region (79.242° S, 117.718° W), mean TBs for the same period are 229.68 \pm 0.56 K and 222.12 \pm 0.74 K at V and H polarization, respectively (Fig. 7). Temporal variability is higher at H than V polarization, which was already observed at L band (Macelloni et al., 2013; Brucker et al., 2013b), and at higher microwave frequencies (e.g., Lacroix et al., 2009; Brucker et al., 2011).

Aquarius provides the lowest microwave frequency observations to routinely observe the cryosphere. The next lowest frequencies are available on AMSR2 in the frequency range 6.9–89 GHz. Observations from these sensors at Dome C are presented for an entire year in a scatter plot of TB H as a function of TB V (Fig. 8). The amplitude of the variation during the year decreases as frequency decreases (i.e. the width of the

ellipse is reduced going from 89 GHz (gray) to 6.9 GHz (red)). This is explained by the lower frequency channels having larger penetration depths, and being thus less sensitive to variations in the temperature profile resulting from solar heating. Considering data from AMSR2 only, values at V polarization increase as frequency decreases (from gray to red), but values at H polarization decrease as frequency decreases (i.e. each





ellipse of AMSR2 TB observations move toward the bottom right corner of Fig. 8 as frequency decreases). TBs observed by each Aquarius' radiometer decrease as the incidence angle increases (from radiometer 1 to 3; Fig. 8). Aquarius TB V observations are higher than every annual mean AMSR2 TB V. However, Aquarius TB H observations are higher than any of the annual mean AMSR2 TB H. Aquarius radiometers operate at lower incidence angles than AMSR2 (< 45.6° vs. 55°), which likely explains most of the difference. Modeling studies are required to quantify the contributions of the different incidence angles.

4.3 Sea ice

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In contrast to ice sheet, sea ice is a more dynamic cover. Sea ice forms, grows, and melts according to the seasons, and some sea ice remains year round in certain regions of both the Arctic and Antarctic. Microwave TB observations depend on the physical properties of sea ice and its overlying snow cover, both of which evolve with time. They also depend on large scale properties of the sea ice cover, whether the sea ice cover is packed or fractured with the presence of leads (open water or new/thin ice).

To present Aquarius weekly-polar-gridded TB product over sea ice, all data associated with the presence of sea ice is used. This implies that extensive areas with low TBs are also considered (Figs. 9 and 10). For reference, TB V of ice-free ocean is typically less than 125 K for radiometer 3, and less than 107 K for radiometer 1. TBs higher than that are observed with a presence of sea ice.

In the NH (Fig. 9), TBs are the highest when snow on sea ice is melting (typically in May and June). TBs are lower during the summer months likely when the sea ice cover is more fractured, with leads exposing open sea water. Overall, a similar TB evolution occurs during the year over the SH sea ice cover (Fig. 10). In Fig. 10, the Ross polynia (75° S, 178° E) can be observed in Summer (January–March).

Sea ice can extend to low latitudes (especially in the SH down to latitudes of ~ 55° S or less). There, Aquarius spatial coverage is significantly reduced. At higher latitudes, where more observations are available, some noise is visible on the weekly TB maps





(e.g. distinct lines in July; Fig. 9). This is related to the fast changing properties of the sea ice cover during the one-week period used to create the gridded product.

To analyze further Aquarius TBs over sea ice, the relationship between ICEF and TB V is analyzed. To that end, swath TBs and ICEFs were used, so that there is no influence of gridding and averaging. Moreover, data in the Ross Sea were used to avoid land contamination. Figure 11 illustrates two important and distinct points. First, it shows the high sensitivity of TB to the presence of sea ice; there is a continuous increase of TB V (radiometer 1) between ~ 100 K, the level for ice-free water, and ~ 250 K which corresponds to a packed Antarctic sea ice cover in winter. Second, Fig. 11 shows that there are numerous low TB observations (< 107 K) associated with

- ICEFs ranging from 0 to 0.5. From our investigations, we noticed that an ICEF of 0.5 raises TB about 70 K (85 K) above the nominal ocean TB level at V (H) polarization using radiometer 1. ICEF comes from simulated sea ice concentration integrated over the sensor's field of view and weighted by the antenna gain patterns. It is very likely that,
- ¹⁵ for the data with a surprisingly low TB, the NCEP GSF GDAS sea ice concentration overestimated the presence of ice.

Studies carried out with SMOS observations show that the L-band TB contains information about thin sea ice thickness (Kaleschke et al., 2010; Huntemann et al., 2013). Since thin ice is fragile it is associated with a very heterogeneous surface. The Aquar-

²⁰ ius large footprint observations will make the retrievals of thin sea ice thickness very challenging. Though, they may be complementary to CryoSat-2 sea ice thickness as for instance retrieved by Hendricks et al. (2013).

5 Active microwave product analysis

While space-borne L-band synthetic aperture radar (SAR) observations were avail able from Advanced Land Observing Satellite (ALOS) PALSAR, space-borne L-band scatterometer observations were first collected by Aquarius. NRCS analysis over the ice sheets was previously studied at C- and Ku-band (Massom and Lubin, 2006, and





references in there). In this section, NRCS observations are presented at the three different polarizations (VV, VH, and HH) that could be used to infer snow/ice and surface properties over ice sheets, and sea ice.

5.1 Ice sheets

Over the GIS, there is a full spatial coverage of the northern portion of the ice sheet (latitudes higher than 77° N). As with the radiometer (see Sect. 2 and Fig. 1), the amount of valid data considered in this weekly-gridded product is lower for the descending orbits than the ascending orbits due to the RFI contamination.

The different polarizations contain information on the structural characteristics of the ice/firn/snow medium and its surface. NRCS values recorded during the first week of January 2013 over the GIS show similar patterns at the three polarizations (Fig. 12, first row), and similar patterns between the ascending and descending orbit observations (Fig. 12, first row vs. second row), though small differences in intensity exist. At higher frequencies, the spatial distribution of NRCS was correlated with the wet snow zone,

- the percolation zone, and the dry snow zone (Long and Drinkwater, 1994). Seasonal variations are observed on the NRCS, but they are small. Long and Drinkwater (1999) provide a detailed analysis of the spatial and temporal variations of the NRCS over the different zones of the GIS using Ku-band observations, and the relationship with snow properties. Similar processes may be related to the L-band NRCS observations.
- ²⁰ Over the AIS (Fig. 13), the main patterns are also similar at all three polarizations. However, NRCS recorded during ascending and descending orbits are significantly different. Differences between orbit types are visible across the ice sheet, but predominantly in East Antarctica. For instance, differences are easily identified in the Dronning Maud Land sector (20° W–45° E), in the Victoria Land sector (140° E–160° E), and in
- ²⁵ between (e.g. 75° E–105° E). The linear correlation coefficient between NRCS VV observations of the ascending and descending orbits for this first week of July 2013 is 0.52, which illustrates the importance of distinguishing the ascending and descending orbits when analyzing the NRCS observations. At the three polarizations, NRCS obser-





vations reveal interesting differences between the two sides of the WAIS divide, where snow accumulation and surface roughness substantially differ.

5.2 Sea ice

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NRCS observations over sea ice are only presented at VV polarization using beam 3 (Fig. 14). The lack of observations in the Beaufort sea (North of Alaska) results from the RFI contamination (Fig. 1, left for one example with the radiometers). Additionally, since only ascending orbits were considered, the spatial coverage is sparse in all Arctic seas; only the Central Arctic Ocean is fully monitored. There is a strong seasonal cycle of NRCS, and in the NH, NRCS are the lowest (about –20 dB) when the sea ice cover is packed and snow covered. NRCS values increase in summer and reach their maximum (about –12 dB) in September/October before the snow accumulates. In winter, sea ice north of Greenland and the Canadian Archipelago have a higher NRCS (about –18 dB vs. –20 dB or lower elsewhere). These higher values correspond to a region of older and thicker sea ice with less brine, where roughness of each interface within the 3-dB restrict can be large compared to the wavelength (21 cm).

¹⁵ footprint can be large compared to the wavelength (21 cm).

The strong seasonal cycle also exists over Antarctic sea ice (Fig. 15). It is most visible in January–March, when summer sea ice remains. The very low NRCS values far from the Antarctic coast in summer are characteristics of open water. As noted in Sect. 4.2 (and Fig. 11), it is very likely that the NCEP GSF GDAS sea ice concentration overestimated the presence of ice. During the winter months, NRCS values over Antarctic sea ice shows a transition ("bright NRCS ring") from the packed sea ice to lower ICEF (Fig. 15). The transition corresponds roughly to an ICEF contour of 0.3–0.4.

6 L-band remote sensing of land

The detection and mapping of frozen soils can be inferred using passive (e.g., Rauti-²⁵ ainen et al., 2012) and active (e.g., McDonald et al., 2012) L-band observations. To



highlight the potential of the weekly-gridded products for the monitoring of frozen soils, a two-year time series was extracted from the TB and NRCS data sets in three locations over subarctic environments (in northern America, Europe, and Asia). In these regions, seasonal snow is present and frozen soil/permafrost may exist (Brown et al., 2001). The

⁵ ratio TB H over TB V (TB H/V) increases in fall, remains high during the whole winter, and decreases in spring (Fig. 16, top), typical of soil physical state change.

The time series of NRCS in the Asian location also show very distinct values in winter vs. summer (Fig. 16, bottom). At the three polarizations (VV, VH, and HH), NRCS are low during the winter months. These time series show information contained in both the passive and active L-band data sets are related to the soil physical state and its surface, and require additional research.

7 L-band remote sensing of SSS

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As mentioned previously, strong and direct interactions between the cryosphere and ocean exist. They occur primarily with sea ice and ice shelves. The produced weekly-gridded product of SSS enables monitoring SSS changes in the polar regions, and possibly freshening resulting from the melting cryosphere.

The SSS values come from the Level 2 product. The algorithm used to retrieve them performs well in the tropics and mid latitudes (Lagerloef et al., 2013). However, SSS retrievals have not yet been specifically validated in cold water, and should be used with caution.

SSS retrieval in the polar oceans is challenging. Polar ocean waters are cold and Aquarius observations are less sensitive to salinity in cold waters. In addition, salinity retrieval is less accurate for very rough sea surfaces. For instance, in the Southern Ocean there are strong winds and the oceanic circulation is dominated by the Antarctic

²⁵ Circumpolar Current, which reduce the quality of the SSS retrievals. Finally, the presence of sea ice adds complexity to the monitoring of SSS in the high latitudes. As presented in Sect. 4.2, TB is highly sensitive to ICEF, and there is no correction in the





SSS algorithm for the presence of ICEF. The presence of ICEF increases TB (Fig. 11) and, if not identified accurately, is interpreted by the SSS-retrieval algorithm as a decrease of SSS. Therefore, one should be particularly careful when studying SSS in the vicinity of sea ice edge and ice sheet. Put simply, increasing TB due to the presence of ice can appear as erroneous freshening.

When SSS retrievals are possible in the Arctic Ocean (Fig. 17), values are low (< 30 psu). One can also see that blue colors are usually present in the vicinity of the sea ice margin. This may result from sea ice contamination of the Aquarius observations, and does not necessarily reflect a low SSS. These maps also show the significant difference between the Atlantic Ocean (saltier) and the Pacific Ocean (fresher). SSS in the polar regions varies seasonally as ice freezes up (leaving saltier water in the ocean) and melts (adding fresh water in the ocean).

SSS retrievals in the Southern Ocean (Fig. 18) show a seasonal cycle with lower SSS near the sea-ice edge in winter. This again results from sea ice contamination,

because in winter during sea ice formation and growth higher SSS are expected. In addition, unrealistic low SSS retrievals can be associated with the presence of icebergs (e.g. localized SSS lower than 32 psu in January where no land is present nearby). This shows that SSS retrievals should be interpreted with care in polar oceans and polynia until better corrections for sea ice and land contaminations are provided.

20 8 Conclusions

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Recent L-band satellite missions collect passive and active observations which offer new capabilities to study and to monitor the polar regions and its cryosphere. In this paper, we presented three seven-day averaged gridded products of TB, NRCS, and SSS. The product also contains ICEF. For each 36 km grid cell where satellite observations/retrievals are available the standard deviation of TB, NRCS, and SSS is provided

along with the number of footprint observations collected during the seven-day cycle. Where less than six contiguous pixels do not have data, a linear interpolation is ap-





plied. Interpolated values can be tracked by the users using the number of footprint observations. The three products are available in HDF5 format starting from late August 2011. It is anticipated that they will be updated on a monthly basis following the release schedule of the Level 2 data sets.

- L-band observations are sensitive to RFI. Their presence is localized mostly in the NH, and substantially impacts the monitoring of continental land, the GIS, and the Beaufort sea. The magnitude of the RFI contribution depends on the orbit (ascending vs. descending). Over the GIS, observations collected during the ascending orbits are significantly less impacted by RFI.
- ¹⁰ Due to the anisotropy of the ice-covered surface, NRCS observations recorded during the ascending and descending orbits must be studied separately. Substantial difference exists over sea ice and the AIS.

These products enhance the use of combined passive and active L-band observations to study the cryosphere. The spatial distributions of Aquarius TB and NRCS over

- the Greenland and Antarctic ice sheet were presented. We showed that the 2012 melt event over the GIS highest elevations impacted the L-band TB, with implications for surface mass balance monitoring. NRCS values over Antarctic sea ice shows a transition ("bright NRCS ring") from the packed sea ice to lower ICEF of 0.3–0.4. We also showed that both active and passive observations over land present distinct values in
- ²⁰ summer and winter, depending on the soil physical state. Aquarius 3-dB footprints are large and make the monitoring of sea ice complex. In both hemispheres, TB over sea ice was the highest when snow on sea ice melts, and the lowest in summer. Studying the surface ocean freshening resulting from ice sheet or sea ice melting is challenging. The sensitivity of the Aquarius SSS retrievals is degraded in cold waters, with sea ice and icebergs in the field of view, and over rough oceans.

The products described in this manuscript are available on the NASA Cryosphere Science Research Portal http://neptune.gsfc.nasa.gov/csb/index.php?section=273 in HDF5 format for improving our understanding of low microwave frequency observa-





tions, and for the development of new or refined algorithms. Aquarius data shows great promise for new cryospheric studies enabled by these weekly-polar-gridded products.

Acknowledgements. This research was funded by the NASA Scientific Innovation Fund. We acknowledge PO.DAAC for distributing the Level 2 Aquarius data (version 2.0). We also acknowledge Annette Decharon and Lisa Taylor (University of Maine, School of Marine Sciences) for hosting the visualization of the weekly-gridded products of brightness temperature and sea surface salinity on the Aquarius Education and Public Outreach website (http://aguarius.umaine.edu/cgi/gal latitudes tbv.htm and http://aguarius.umaine.edu/cgi/gal latitudes sss.htm, respectively).

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Table 1.	Beam-dependent	characteristics	of the Aquarius	L-band radiometers.
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Room	1	2	2
Dealli	1	2	3
Incidence angle (°)	28.7	37.8	45.6
3 dB footprint size along × across track (km × km)	76 × 94	84 × 120	97 × 156
Northernmost latitude monitored (° N)	84.99	86.07	87.40
Southernmost latitude monitored (°S)	79.01	77.90	76.54

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Beam	1	2	3
3 dB two-way footprint size along × across track (km × km)	58 × 71	65 × 91	74 × 122
NE σ^0 (dB)	-29	-26	-24
Scatterometer sensitivity (dB)	0.04	0.06	0.1
Scatterometer stability for 7 days (dB)	0.13	0.13	0.13





Table 3. Product names with the list of variables per beams and orbits.

Aquarius 36 km weekly-polar-gridded brightne TBV. TBH	ss temperature, and sea surface salinity				
TBV_STD, TBH_STD NFP_RAD SSS, SSS_STD ICEF_RAD, ICEF_STD_RAD	Beams 1, 2, and 3 Orbits ascending, descending, and combined				
Aquarius 36 km weekly-polar-gridded normalized radar cross section					
NRCS_VV, NRCS_VH, NRCS_HH NRCS_VV_STD, NRCS_VH_STD, NRCS_HH_STD NFP_SCA ICEF_SCA, ICEF_STD_SCA	Beams 1, 2, and 3 Orbits ascending, and descending				
Aquarius 36 km weekly-polar-gridded sea surface salinity 3 beams (SSS3b)					
SSS30 SSS3b_STD NFP_SSS3b ICEF_SSS3b, ICEF_STD_SSS3b	Beams 1, 2, and 3 combined Orbits combined				







Fig. 1. Map of RFI occurences (i.e. number of footprint with either a moderate or severe RFI contamination) during cycle 72 (3 January–10 January 2013) in the NH for ascending (left) and descending (right) orbits.





Fig. 2. Maps of TB (beam 3) at V polarization (in K) over the GIS for every other month. Grid cells with land fraction less than 0.25 were masked.





Fig. 3. Distributions of TB V (left) and TB H (right) observed over the GIS by radiometer 3 during the first week of January 2013 (red) and July 2013 (blue).







Fig. 4. Time series of TB (beam 1) at V (+) and H (o) polarizations near Summit (black), and South Dome (gray).



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Fig. 5. Maps of TB (beam 3) at V polarization (in K) over the AIS for every other month. Grid cells with land fraction less than 0.25 were masked.



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Fig. 6. Distributions of TB V (left) and TB H (right) observed over the AIS using radiometer 3 during the first week of January 2013 (red) and July 2013 (blue).







Fig. 7. Time series TB (beam 1) at V (+) and H (o) polarizations over Dome C (black), and Mary Byrd land (gray).





Fig. 8. TB H as a function of TB V at Dome C, Antarctica, using one year of data. 2012 Aquarius TBs and 2010 AMSR2 TBs were used providing 1.413 GHz data at three different incidence angles, and 6.9, 10.7, 18.7, 36.5, and 89 GHz data, respectively.



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Fig. 9. Maps of TB (beam 3) at V polarization over the NH sea ice for every other month. Grid cells where the land fraction was more than 0.25 were masked.



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Fig. 10. Maps of TB (beam 3) at V polarization over the SH sea ice for every other month. Grid cells where the land fraction was more than 0.25 were masked.



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Fig. 11. ICEF as a function of TB V (beam 1) in and north of the Ross Sea (see embedded map with the location of the footprint measurements). Swath data from cycle 85 to 105 (i.e. between 4 April 2013 and 29 August 2013).



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Fig. 12. NRCS (beam 3) at VV (left panels), VH (middle panels), and at HH (right panels) polarizations recorded over the GIS during the cycle 72 (3–10 January 2013) and during ascending orbits (first row), and descending orbits (second row). Grid cells where the land fraction was less than 0.25 were masked.



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Fig. 13. NRCS (beam 3) at VV (left panels), VH (middle panels), and at HH (right panels) polarizations recorded over the AIS during the cycle 98 (4–11 July 2013) and during ascending orbits (first row), and descending orbits (second row). Grid cells where the land fraction was less than 0.25 were masked.





Fig. 14. Maps of beam 3 NRCS VV observed during ascending orbits over the NH sea ice for every other month. Grid cells where the land fraction was more than 0.25 were masked.





Fig. 15. Maps of beam 3 NRCS VV observed during ascending orbits over the SH sea ice for every other month. Grid cells where the land fraction was more than 0.25 were masked.





Fig. 16. (Top) time series of the TB H/V ratio in three locations Northern America (Alaska) 66° N, 150° W (black *), Northern Europe (Finland) 67.5° N, 29° E (blue o), and Asia (Russia) 66° N, 149.5° E (red +). (Bottom) time series NRCS at VV (plus symbols), VH (circles), and HH (downward triangles) for the location in Asia.







Fig. 17. Maps of SSS3b over the NH high latitudes for cycles every other month. Grid cells where the land fraction was more than 0.25 were masked.





Fig. 18. Maps of SSS3b over the SH high latitudes for cycles every other month. Grid cells where the land fraction was more than 0.25 were masked.



